

**SURVEY AND LOGISTICS REPORT
ON A HELICOPTER BORNE
VERSATILE TIME DOMAIN
ELECTROMAGNETIC (VTEM)
SURVEY**

on the

**CARANBINI AREA
AUSTRALIA**

for

BRUMBY RESOURCES LIMITED

by



GEOTECH AIRBORNE PTY LTD

*Unit 1, 29 Mulgool Road
Malaga, WA, 6090,
Australia
Tel: +61 8 9249 8814
Fax: +61 8 9249 8894
www.geotechairborne.com
e-mail: info@geotechairborne.com*

**Project AA911
December, 2010**

TABLE OF CONTENTS

1. SURVEY SPECIFICATIONS	3
1.1. General	3
1.2. VTEM flight plan on Google EARTH™ Background	3
1.3. Survey block coordinates.....	4
1.4. Survey block specifications.....	4
1.5. Survey schedule.....	4
2. SYSTEM SPECIFICATIONS	5
2.1. Instrumentation	5
2.2. VTEM Configuration	6
2.3. VTEM decay sampling scheme	6
2.4. VTEM Transmitter Waveform over one half-period (November 2010).....	7
3. PROCESSING	8
3.1. Processing parameters.....	8
3.2. Flight Path.....	8
3.3. Electromagnetic Data	8
3.4. Magnetic Data	8
3.5. Digital Terrain Model	9
4. DELIVERABLES	10
5. PERSONNEL.....	12

APPENDICES

A. Modeling VTEM data	13
B. VTEM X-Component data	19
C. Geophysical Maps	28



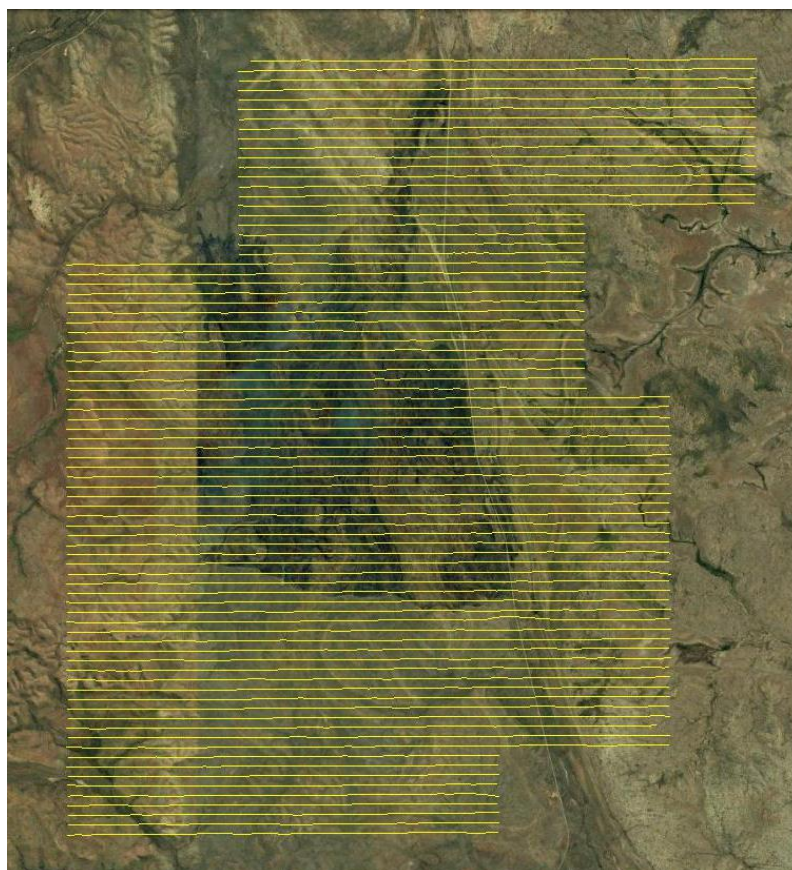
SURVEY AND LOGISTICS REPORT ON A HELICOPTER-BORNE VTEM SURVEY

1. SURVEY SPECIFICATIONS

1.1. General

Job Number	AA911
Client	Brumby Resources Limited.
Project Area	Caranbini Area
Location	Australia
Number of Blocks	1
Total line kilometres	904km
Survey date	18 - 21 November, 2010
Client Representative	John Ikstrums Tel: +61 8 9486 8333 Fax: +61 8 9322 5123 Email: john@brumbyresources.com.au
Client address	Unit 3, 49 Ord Street West Perth, WA 6005, Australia

1.2. VTEM flight plan on Google EARTH™ Background



1.3. Survey block coordinates.

<i>Easting UTM Z 53S</i>	<i>Northing UTM Z 53S</i>
<i>Caranbini Area</i>	
610440.83	8206365.81
621129.80	8206311.23
621111.92	8203079.02
617549.73	8203094.91
617529.57	8199407.09
619310.52	8199397.44
619270.13	8192021.73
615709.46	8192040.97
615699.64	8190197.06
606798.86	8190242.66
606857.25	8202152.05
610419.48	8202134.35
610440.83	8206365.81

1.4. Survey block specifications

<i>Survey block</i>	<i>Line spacing (m)</i>	<i>Line-km (contractual)</i>	<i>Line-km (delivered)</i>	<i>Flight direction</i>	<i>Line number</i>
<i>Caranbini</i>	200	898	904	090-270	L10010 – L10810

1.5. Survey schedule

<i>Date</i>	<i>Flight #</i>	<i>Block</i>	<i>Nominal Production Km flown</i>	<i>Comments</i>
18-Nov-10	1-4	Caranbini	220	Production
19-Nov-10	5-7	Caranbini	178	Production
20-Nov-10	8-12	Caranbini	352	Production
21-Nov-10	13-15	Caranbini	156	Production



2. SYSTEM SPECIFICATIONS

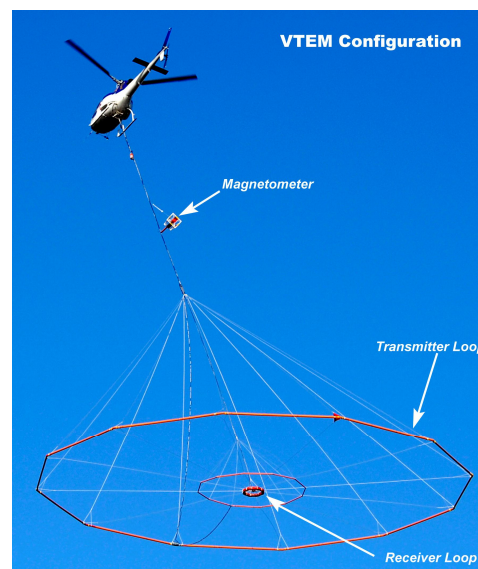
2.1. Instrumentation

Survey Helicopter	
Model	Augusta AW119-KE
Registration	VH-KEJ
Nominal survey speed	80 km/h
Nominal terrain clearance	88 m
VTEM Transmitter	
Coil diameter	35 m
Number of turns	4
Pulse repetition rate	25 Hz
Peak current	226 Amp
Duty cycle	28.67%
Peak dipole moment	870,099 NIA
Pulse width	5.74 ms
Nominal terrain clearance	34 m
VTEM Receiver	
Coil diameter	1.2 metre
Number of turns	100
Effective area	113.1 m ²
Sampling interval	0.1 s
Nominal terrain clearance	42 m
Magnetometer	
Type	Geometrics
Model	Optically pumped cesium vapour
Sensitivity	0.02 nT
Sampling interval	0.1 s
Cable length	13 m
Nominal terrain clearance	78 m
Radar Altimeter	
Type	Terra TRA 3000/TRI 40
Position	Beneath cockpit
Sampling interval	0.2 s
GPS navigation system	
Type	NovAtel
Model	WAAS enabled OEM4-G2-3151W
Antenna position	Helicopter tail
Sampling interval	0.2 s
Base Station Magnetometer/GPS	
Type	Geometrics
Model	Cesium vapour
Sensitivity	0.001 nT
Sampling interval	1 s



2.2. VTEM Configuration

Configuration	
Cable angle with vertical	35 °
Cable length (EM receiver)	57 m
Cable length (Magnetometer)	13 m

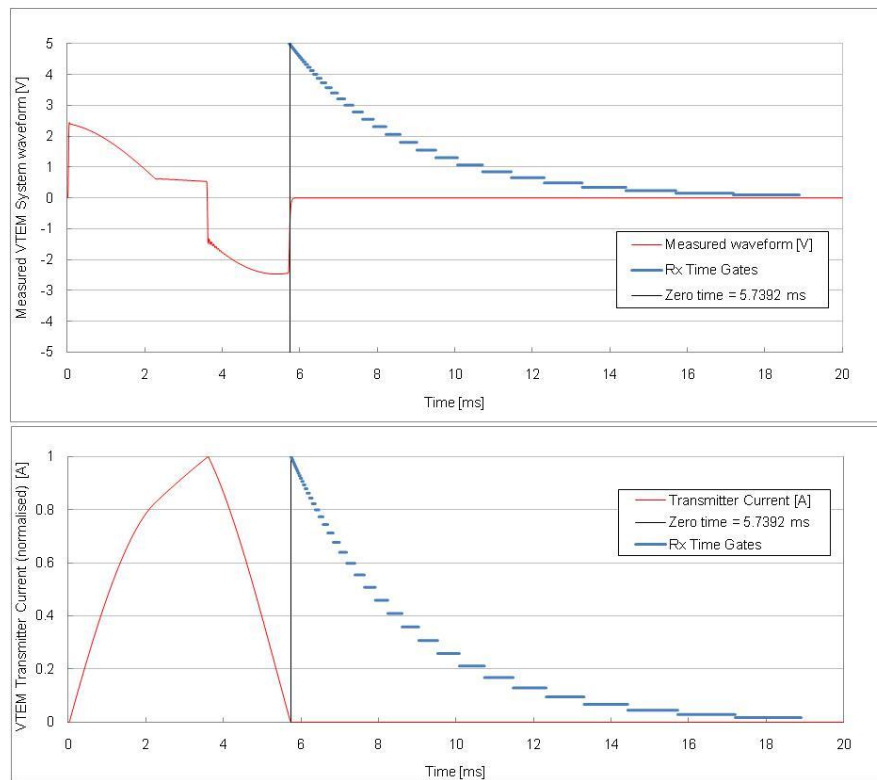


2.3. VTEM decay sampling scheme

B-field VTEM Decay Sampling scheme				
Array	Microseconds			
Index	Middle	Start	End	Width
13	83	78	90	12
14	96	90	103	13
15	110	103	118	15
16	126	118	136	18
17	145	136	156	20
18	167	156	179	23
19	192	179	206	27
20	220	206	236	30
21	253	236	271	35
22	290	271	312	40
23	333	312	358	46
24	383	358	411	53
25	440	411	472	61
26	505	472	543	70
27	580	543	623	81
28	667	623	716	93
29	766	716	823	107
30	880	823	945	122
31	1010	945	1086	141
32	1161	1086	1247	161
33	1333	1247	1432	185
34	1531	1432	1646	214
35	1760	1646	1891	245
36	2021	1891	2172	281
37	2323	2172	2495	323
38	2667	2495	2865	370
39	3063	2865	3292	427
40	3521	3292	3781	490
41	4042	3781	4341	560
42	4641	4341	4987	646
43	5333	4987	5729	742
44	6125	5729	6581	852
45	7036	6581	7560	979
46	8083	7560	8685	1125
47	9286	8685	9977	1292
48	10667	9977	11458	1482



2.4. VTEM Transmitter Waveform over one half-period (November 2010)



3. PROCESSING

3.1. Processing parameters

Coordinates	
Projection	MAP GRID AUS ZONE 53
Datum	GDA 94
Spherics rejection (EM and Magnetic data)	
Non-linear filter	4 point
Non-linear filter sensitivity	0.00001
Low-pass filter wavelength	20 fids
Lag correction of other sensors to EM receiver position	
GPS	25 m
Radar	35 m
Magnetometer	25.5 m

3.2. Flight Path

The flight path, recorded by the acquisition program as WGS 84 latitude/longitude, was converted into the UTM coordinate system in Oasis Montaj. The flight path was drawn using linear interpolation between x,y positions from the navigation system. Positions are updated every second and expressed as UTM eastings (x) and UTM northings (y).

3.3. Electromagnetic Data

A three stage digital filtering process was used to reject major sferic events and to reduce system noise. Local sferic activity can produce sharp, large amplitude events that cannot be removed by conventional filtering procedures. Smoothing or stacking will reduce their amplitude but leave a broader residual response that can be confused with geological phenomena. To avoid this possibility, a computer algorithm searches out and rejects the major sferic events.

The signal to noise ratio was further improved by the application of a low pass linear digital filter. This filter has zero phase shift which prevents any lag or peak displacement from occurring, and it suppresses only variations with a wavelength less than the specified filter wavelength.

3.4. Magnetic Data

The processing of the magnetic data involved the correction for diurnal variations by using the digitally recorded ground base station magnetic values. The base station magnetometer data was edited and merged into the Geosoft GDB database on a daily basis. The aeromagnetic data was corrected for diurnal variations by subtracting the observed magnetic base station deviations.

A micro-levelling procedure was then applied. This technique is designed to remove persistent low-amplitude components of flight-line noise.

The corrected magnetic data was interpolated between survey lines using a random point gridding method to yield x-y grid values for a standard grid cell size of a quarter of the line spacing. The Minimum Curvature algorithm was used to interpolate values onto a rectangular regular spaced grid.



3.5. Digital Terrain Model

Subtracting the radar altimeter data from the GPS elevation data creates a digital elevation model.



4. DELIVERABLES

VTEM Survey and logistics report		
Format	PDF	
Database		
Format	Digital Geosoft (.GDB)	
Channels	Name	Description
	X_UTM	X positional data (UTM Z53S / WGS84)
	Y_UTM	Y positional data (UTM Z53S / WGS84)
	X_MGA	X positional data (MGA Z53 / GDA94)
	Y_MGA	Y positional data (MGA Z53 / GDA94)
	Lon	Longitude data
	Lat	Latitude data
	Z	GPS antenna elevation (metres above sea level)
	Radar	Helicopter terrain clearance from radar altimeter (metres above ground level)
	RxAlt	EM Receiver and Transmitter terrain clearance (metres above ground level)
	DTM	Digital terrain model (metres)
	Gtime	UTC time (seconds of the day)
	MagTF	Raw Total Magnetic field data (nT)
	MagBase	Magnetic diurnal variation data (nT)
	MagDiu	Total Magnetic field diurnal variation and lag corrected data (nT)
	MagMicL	Microleveled Total Magnetic field data (nT)
	dBdtZ[13] to dBdtZ[48]	dB/dtZ, Time Gates 83 μ s to 10667 μ s (pV/A/m ⁴)
	BfieldZ[13] to BfieldZ[48]	B-fieldZ, Time Gates 83 μ s to 10667 μ s (pV.ms/A/m ⁴)
	dBdtX[20] to dBdtX[48]	dB/dtX, Time Gates 220 μ s to 10667 μ s (pV/A/m ⁴)
	BfieldX[20] to BfieldX[48]	B-fieldX, Time Gates 220 μ s to 10667 μ s (pV.ms/A/m ⁴)
	dBdtX_FF[20] to dBdtX_FF[48]	Fraser Filter dB/dtX, Time Gates 220 μ s to 10667 μ s (pV/A/m ⁴)
	BfieldX_FF[20] to BfieldX_FF[48]	Fraser Filtered B-fieldX, Time Gates 220 μ s to 10667 μ s (pV.ms/A/m ⁴)
	dBdtX_SFF_20_30	Stacked Fraser Filtered data from channel 20 to 30 (pV/A/m ⁴)
	BfieldX_SFF_20_30	Stacked Fraser Filtered data from channel 20 to 30 (pV.ms/A/m ⁴)
	dBdtY[20] to dBdtY[48]	dB/dtY, Time Gates 220 μ s to 10667 μ s (pV/A/m ⁴)
	BfieldY[20] to BfieldY[48]	B-fieldY, Time Gates 220 μ s to 10667 μ s (pV.ms/A/m ⁴)
	dBdtY_FF[20] to dBdtY_FF[48]	Fraser Filter dB/dtY, Time Gates 220 μ s to 10667 μ s (pV/A/m ⁴)
	BfieldY_FF[20] to BfieldY_FF[48]	Fraser Filtered B-fieldY, Time Gates 220 μ s to 10667 μ s (pV.ms/A/m ⁴)
	dBdtY_SFF_20_30	Stacked Fraser Filtered data from channel 20 to 30 (pV/A/m ⁴)
	BfieldY_SFF_20_30	Stacked Fraser Filtered data from channel 20 to 30 (pV.ms/A/m ⁴)
	PLM	Power line monitor



Grids		
Format	Digital Geosoft (.GRD and .GI) ¹	
Grids	Name	Description
	AA911_Mag	Total Magnetic field (nT)
	AA911_dBdtX_SFF	dBdtX Stacked Fraser Filtered data
	AA911_BfieldX_SFF	BfieldX Stacked Fraser Filtered data
	AA911_dBdtY_SFF	dBdtY Stacked Fraser Filtered data
	AA911_BfieldY_SFF	BfieldY Stacked Fraser Filtered data

Maps		
Format	Digital Geosoft (.MAP)	
Scale	1:25 000	
Maps	Name	Description
	AA911_Mag	Total Magnetic field colour contours
	AA911_dBdtZ_Log	VTEM dB/dt profiles, Time Gates 0.667 – 10.667 ms in linear - logarithmic scale
	AA911_BfieldZ_Log	VTEM B-field profiles, Time Gates 0.667 – 10.667 ms in linear - logarithmic scale
	AA911_dBdtX_SFF	dBdtX Stacked Fraser Filtered data
	AA911_BfieldX_SFF	BfieldX Stacked Fraser Filtered data
	AA911_dBdtY_SFF	dBdtY Stacked Fraser Filtered data
	AA911_BfieldY_SFF	BfieldY Stacked Fraser Filtered data

Waveform		
Format	Digital Excel Spreadsheet (AA911_VTEM_Waveform.xls)	
Columns	Name	Description
	Time	Sampling rate interval, 10.416 μ s
	Volt	Output voltage of the receiver coil (volt)
	Current	Transmitter current (normalised to 1A peak)

Google Earth Flight Path file	
Format	Google Earth AA911_FlightPath.kml
	Free version of Google Earth software can be downloaded from, http://earth.google.com/download-earth.html

¹ A Geosoft .GRD file has a .GI metadata file associated with it, containing grid projection information.



5. PERSONNEL

<i>Geotech Airborne Limited Personnel</i>	
<i>Operator / Crew chief</i>	<i>Gregory Downs</i>
<i>Data Processing (Preliminary)</i>	<i>Pete Holbrook</i>
<i>Data Processing (Final) /Reporting</i>	<i>Matt Holbrook</i>
<i>Final data supervision</i>	<i>Malcolm Moreton</i> <i>Data Processing Manager</i> <i>(malcolm@geotechairborne.com)</i>
<i>Overall project management</i>	<i>Keith Fisk</i> <i>Managing Partner and Director</i> <i>(keith@geotechairborne.com)</i>



APPENDIX A

GENERALIZED MODELING RESULTS OF THE VTEM SYSTEM (by Roger Barlow)

Introduction

The VTEM system is based on a concentric or central loop design, whereby, the receiver is positioned at the centre of a 35 metres diameter transmitter loop that produces a dipole moment up to 870,099 NIA at peak current. The wave form is a bi-polar, modified square wave with a turn-on and turn-off at each end. With a base frequency of 25 Hz, the duration of each pulse is approximately 7.5 milliseconds followed by an off time where no primary field is present.

During turn-on and turn-off, a time varying field is produced (dB/dt) and an electro-motive force (emf) is created as a finite impulse response. A current ring around the transmitter loop moves outward and downward as time progresses. When conductive rocks and mineralization are encountered, a secondary field is created by mutual induction and measured by the receiver at the centre of the transmitter loop.

Measurements are made during the off-time, when only the secondary field (representing the conductive targets encountered in the ground) is present.

Efficient modeling of the results can be carried out on regularly shaped geometries, thus yielding close approximations to the parameters of the measured targets. The following is a description of a series of common models made for the purpose of promoting a general understanding of the measured results.

Variation of Plate Depth

Geometries represented by plates of different strike length, depth extent, dip, plunge and depth below surface can be varied with characteristic parameters like conductance of the target, conductance of the host and conductivity/thickness and thickness of the overburden layer.

Diagrammatic models for a vertical plate are shown in figures A and G at two different depths, all other parameters remaining constant. With this transmitter-receiver geometry, the classic M shaped response is generated. Figure A shows a plate where the top is near surface. Here, amplitudes of the dual peaks are higher and symmetrical with the zero centre positioned directly above the plate. Most important is the separation distance of the peaks. This distance is small when the plate is near surface and widens with a linear relationship as the plate (depth to top) increases. Figure G shows a much deeper plate where the separation distance of the peaks is much wider and the amplitudes of the channels have decreased.

Variation of Plate Dip

As the plate dips and departs from the vertical position, the peaks become asymmetrical. Figure B shows a near surface plate dipping 80°. Note that the direction of dip is toward the high shoulder of the response and the top of the plate remains under the centre minimum.

As the dip increases, the aspect ratio (Min/Max) decreases and this aspect ratio can be used as an empirical guide to dip angles from near 90° to about 30°. The method is not sensitive enough where dips are less than about 30°. Figure E shows a plate dipping 45° and, at this angle, the minimum shoulder starts to vanish. In Figure D, a



flat lying plate is shown, relatively near surface. Note that the twin peak anomaly has been replaced by a symmetrical shape with large, bell shaped, channel amplitudes which decay relative to the conductance of the plate.

Figure H shows a special case where two plates are positioned to represent a synclinal structure. Note that the main characteristic to remember is the centre amplitudes are higher (approximately double) compared to the high shoulder of a single plate. This model is very representative of tightly folded formations where the conductors were once flat lying.

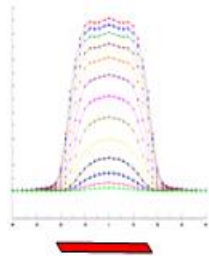
Variation of Prism Depth

Finally, with prism models, another algorithm is required to represent current on the plate. A plate model is considered to be infinitely thin with respect to thickness and incapable of representing the current in the thickness dimension. A prism model is constructed to deal with this problem, thereby, representing the thickness of the body more accurately.

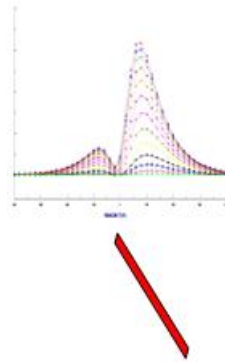
Figures C, F and I show the same prism at increasing depths. Aside from an expected decrease in amplitude, the side lobes of the anomaly show a widening with deeper prism depths of the bell shaped early time channels.



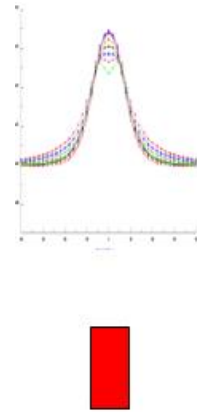
A



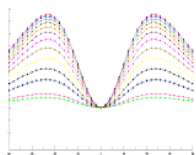
B



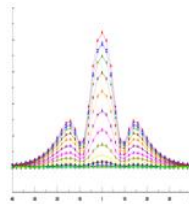
C



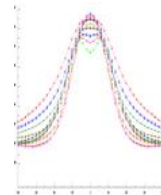
D



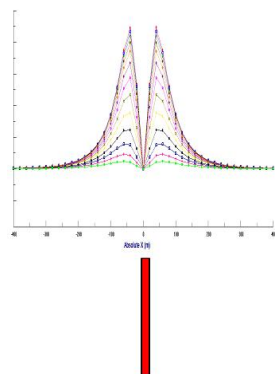
E



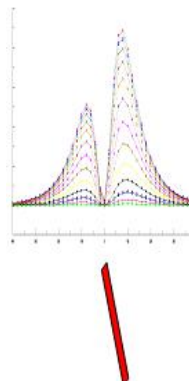
F



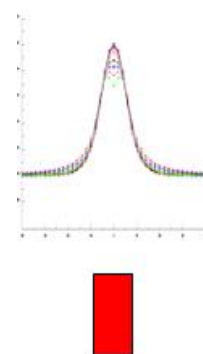
G



H



I



General Modeling Concepts

A set of models has been produced for the Geotech VTEM[®] system with explanation notes (see models A to I above). The reader is encouraged to review these models, so as to get a general understanding of the responses as they apply to survey results. While these models do not begin to cover all possibilities, they give a general perspective on the simple and most commonly encountered anomalies.

When producing these models, a few key points were observed and are worth noting as follows:

- *For near vertical and vertical plate models, the top of the conductor is always located directly under the centre low point between the two shoulders in the classic **M** shaped response.*
- *As the plate is positioned at an increasing depth to the top, the shoulders of the **M** shaped response, have a greater separation distance.*
- *When faced with choosing between a flat lying plate and a prism model to represent the target (broad response) some ambiguity is present and caution should be exercised.*
- *With the concentric loop system and Z-component receiver coil, virtually all types of conductors and most geometries are most always well coupled and a response is generated (see model H). Only concentric loop systems can map this type of target.*

The modelling program used to generate the responses was prepared by PetRos Eikon Inc. and is one of a very few that can model a wide range of targets in a conductive half space.

General Interpretation Principals

Magnetics

The total magnetic intensity responses reflect major changes in the magnetite and/or other magnetic minerals content in the underlying rocks and unconsolidated overburden. Precambrian rocks have often been subjected to intense heat and pressure during structural and metamorphic events in their history. Original signatures imprinted on these rocks at the time of formation have, in most cases, been modified, resulting in low magnetic susceptibility values.

The amplitude of magnetic anomalies, relative to the regional background, helps to assist in identifying specific magnetic and non-magnetic rock units (and conductors) related to, for example, mafic flows, mafic to ultramafic intrusives, felsic intrusives, felsic volcanics and/or sediments etc. Obviously, several geological sources can produce the same magnetic response. These ambiguities can be reduced considerably if basic geological information on the area is available to the geophysical interpreter.



In addition to simple amplitude variations, the shape of the response expressed in the wave length and the symmetry or asymmetry, is used to estimate the depth, geometric parameters and magnetization of the anomaly. For example, long narrow magnetic linears usually reflect mafic flows or intrusive dyke features. Large areas with complex magnetic patterns may be produced by intrusive bodies with significant magnetization, flat lying magnetic sills or sedimentary iron formation. Local isolated circular magnetic patterns often represent plug-like igneous intrusives such as kimberlites, pegmatites or volcanic vent areas.

Because the total magnetic intensity (TMI) responses may represent two or more closely spaced bodies within a response, the second derivative of the TMI response may be helpful for distinguishing these complexities. The second derivative is most useful in mapping near surface linears and other subtle magnetic structures that are partially masked by nearby higher amplitude magnetic features. The broad zones of higher magnetic amplitude, however, are severely attenuated in the vertical derivative results. These higher amplitude zones reflect rock units having strong magnetic susceptibility signatures. For this reason, both the TMI and the second derivative maps should be evaluated together.

Theoretically, the second derivative, zero contour or colour delineates the contacts or limits of large sources with near vertical dip and shallow depth to the top. The vertical gradient map also aids in determining contact zones between rocks with a susceptibility contrast, however, different, more complicated rules of thumb apply.

Concentric Loop EM Systems

Concentric systems with horizontal transmitter and receiver antennae produce much larger responses for flat lying conductors as contrasted with vertical plate-like conductors. The amount of current developing on the flat upper surface of targets having a substantial area in this dimension, are the direct result of the effective coupling angle, between the primary magnetic field and the flat surface area. One therefore, must not compare the amplitude/conductance of responses generated from flat lying bodies with those derived from near vertical plates; their ratios will be quite different for similar conductances.

Determining dip angle is very accurate for plates with dip angles greater than 30°. For angles less than 30° to 0°, the sensitivity is low and dips can not be distinguished accurately in the presence of normal survey noise levels.

*A plate like body that has near vertical position will display a two shoulder, classic **M** shaped response with a distinctive separation distance between peaks for a given depth to top.*

It is sometimes difficult to distinguish between responses associated with the edge effects of flat lying conductors and poorly conductive bedrock conductors. Poorly conductive bedrock conductors having low dip angles will also exhibit responses that may be interpreted as surficial overburden conductors. In some situations, the conductive response has line to line continuity and some magnetic correlation providing possible evidence that the response is related to an actual bedrock source.

The EM interpretation process used, places considerable emphasis on determining an understanding of the general conductive patterns in the area of interest. Each area has different characteristics and these can effectively guide the detailed process used.



The first stage is to determine which time gates are most descriptive of the overall conductance patterns. Maps of the time gates that represent the range of responses can be very informative.

Next, stacking the relevant channels as profiles on the flight path together with the second vertical derivative of the TMI is very helpful in revealing correlations between the EM and Magnetics.

Next, key lines can be profiled as single lines to emphasize specific characteristics of a conductor or the relationship of one conductor to another on the same line. Resistivity Depth sections can be constructed to show the relationship of conductive overburden or conductive bedrock with the conductive anomaly.



APPENDIX B

VTEM X-COMPONENT DATA



6. X COIL DATA

6.1. Sign convention

VTEM's X component data produces crossover type anomalies. This is unlike the Z component of maxima or minima above conductors. During acquisition the convention is for X coil data to be positive in the direction of flight. In the processing phase the polarity is adjusted to follow the right hand rule for multi-component transient electromagnetic methods.

For N-S lines the sign convention for the X in-line component crossover is positive-negative pointing south to north for vertical plate conductors perpendicular to the profile. For E-W lines the sign convention for the X in-line component crossover is positive-negative pointing west to east for vertical plate conductors perpendicular to the profile. X component data for alternating/opposite flight directions are reversed (multiplied by negative one) in the final database to account for this polarity convention.

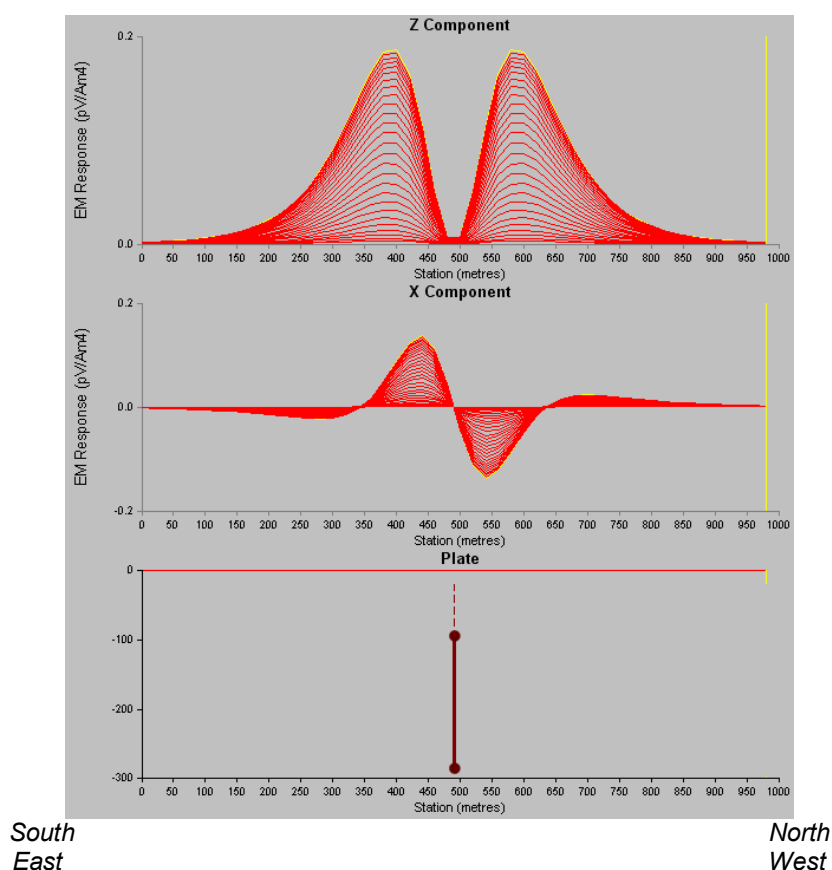


Figure 1: Z- and X-component responses over vertical plate conductor indicating sign convention for X component data.



6.2. Fraser Filter and Stacked Fraser Filter

The Fraser Filter converts crossovers of the correct polarity into peak responses of X component by differencing successive values. It is calculated as $(f_1+f_2)-(f_3+f_4)$ where f_i are data from four consecutive stations. This is a derivative filter and likely to increase any noise in data.

A useful presentation of X-component data is the Staked Fraser Filter. The Stacked Fraser Filter data are calculated as the average value of 11 channels (15 to 25) of Fraser Filtered X-component data. The signal to noise ratio is improved and information from 11 channels are combined into one, which allows easier presentation in grid or map format.

6.3. Effect of loop tilt

Whenever the X coil is not aligned exactly vertical, it also measures a part of the Z-component response. When the Z-component response is much larger in amplitude than the X-component response it can dominate the measured X-component data. This becomes especially evident when line polarities are reversed; true X-component responses would be coherent from line to line after polarity correction whereas the Z-component becomes alternating negative and positive responses. An example of this is shown in Fig. 2.

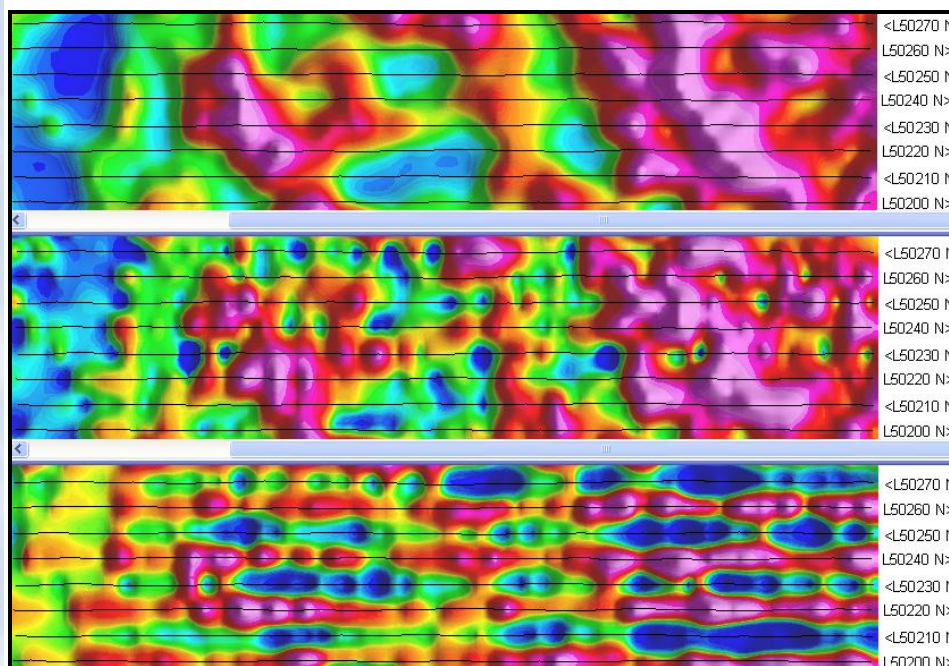


Figure 2: Channel 15 Z-component (top), X-component (middle) and polarity corrected X-component (bottom). The flight line direction is indicated by > or < next to the line number. It is clear that the measured X-component is dominated by Z-component response.



Provided the tilt angle of the coil is known, a correction can be applied to the data. X coil tilt is not measured in with the VTEM system and an approximate procedure was developed to calculate tilt angles in the X - Z plane and remove the Z -component influence on the X -component data.

1.3.1 Quantifying the effect of loop tilt

With Z and X the real vector components and $\underline{Z}$ and $\underline{X}$ the measured components and α the tilt angle of the loop in the line direction, we have:

$$\underline{Z} = Z \cos(\alpha) + X \sin(\alpha)$$

$$\underline{X} = X \cos(\alpha) - Z \sin(\alpha)$$

Over half-space or layered earth environments $X=0$.

So that:

$$\underline{Z} = Z \cos(\alpha)$$

$$\underline{X} = -Z \sin(\alpha)$$

and

$$\alpha = \text{atan}(-\underline{X} / \underline{Z}).$$

Any value of $\underline{X}$ is therefore ascribed to the term $-Z \sin(\alpha)$. This is sometimes observed in VTEM data with measured X values being negative in general and following inverse trends from the Z -component data (Fig. 3).

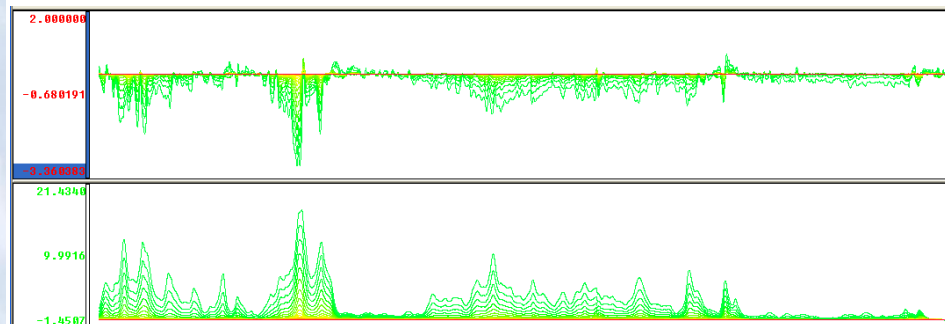


Figure 3: X-component (top) and Z-component (bottom) VTEM data.



1.3.2 Approximate calculation of loop tilt

The equation $\alpha = \text{atan}(-\underline{X} / \underline{Z})$ is used to calculate approximate tilt-angles at each position for all channels. An average tilt angle is calculated from channels 15 to 20 and filtered with:

Non-Linear filter: 100 fids; 0.2 sensitivity

LP filter: 500 fids.

(Parameters are determined empirically for individual data sets)

The filtering ensures that anomalous $\underline{X}$ responses are excluded from subsequent calculations, but also causes small errors in the correction factor. As a final step and average tilt angle (α) for each line is calculated.

1.3.3 Correcting X-component data

A new channel for X-component data is calculated as:

$$\text{SLx_Angle_Corrected} = \underline{X} * \cos(\alpha) + \underline{Z} * \sin(\alpha).$$

The results are to centre responses round zero and visually enhance the occurrences of X-component anomalies (Fig. 4).

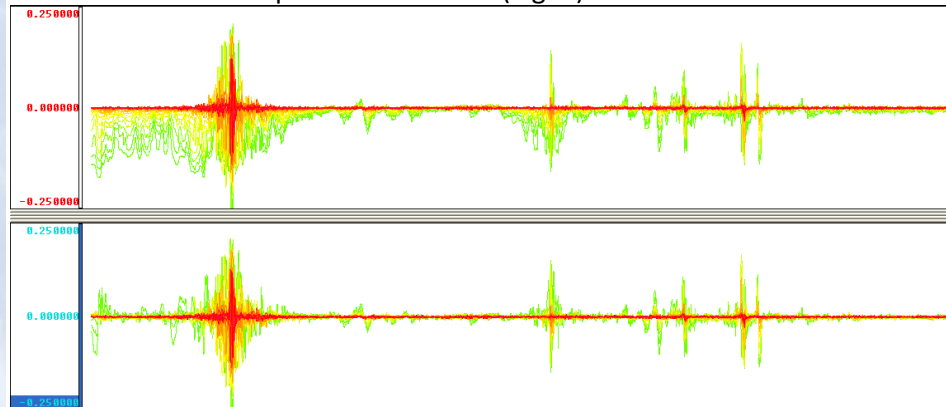


Figure 4: X-component data before (top) and after (bottom) approximate tilt correction.

As this is an approximate correction only, the corrected data might not give accurate results in quantitative modeling or inversions.



The results of the tilt correction and subsequent polarity corrections on gridded data are shown in Fig. 5. After the tilt correction the long wavelength similarities between X- and Z-coil data is no longer evident. The polarity correction now improves line-to-line continuity of as expected.

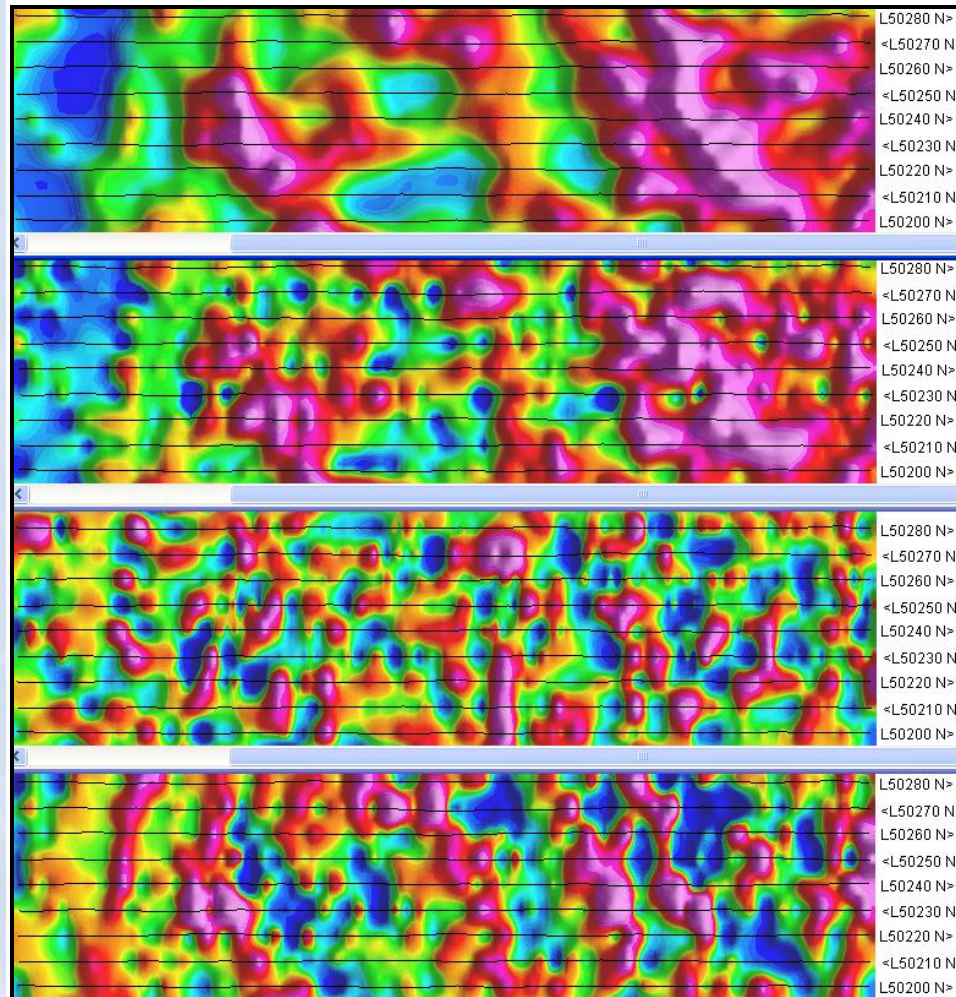
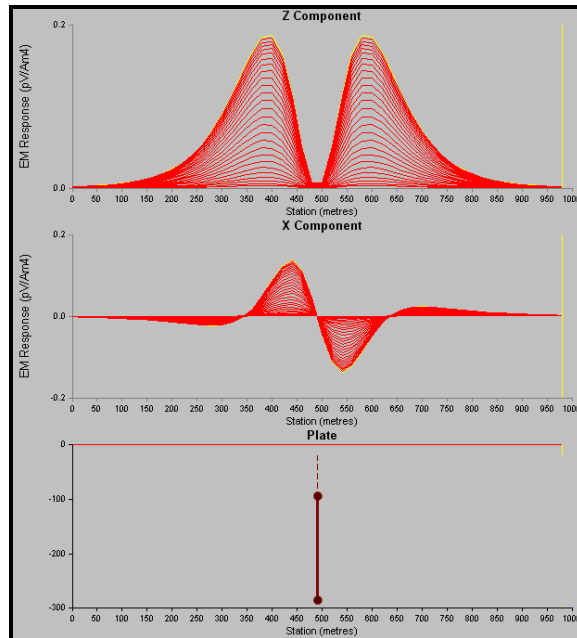


Figure 5: Channel 15 Z-component (top), X-component (second from top), X-component corrected for tilt (second from bottom) and polarity corrected X-component after tilt correction (bottom).

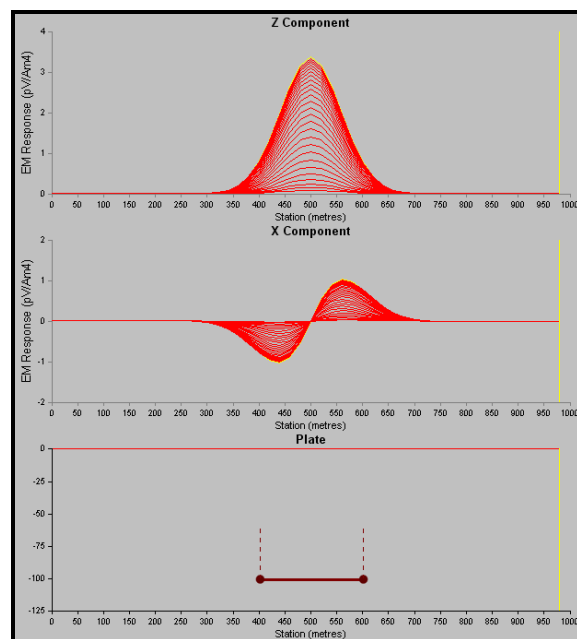


6.4. Typical X-component responses

Forward modeled X- and Z-component responses for some typical plate models in free air are provided below as reference. These were calculated with Maxwell software.

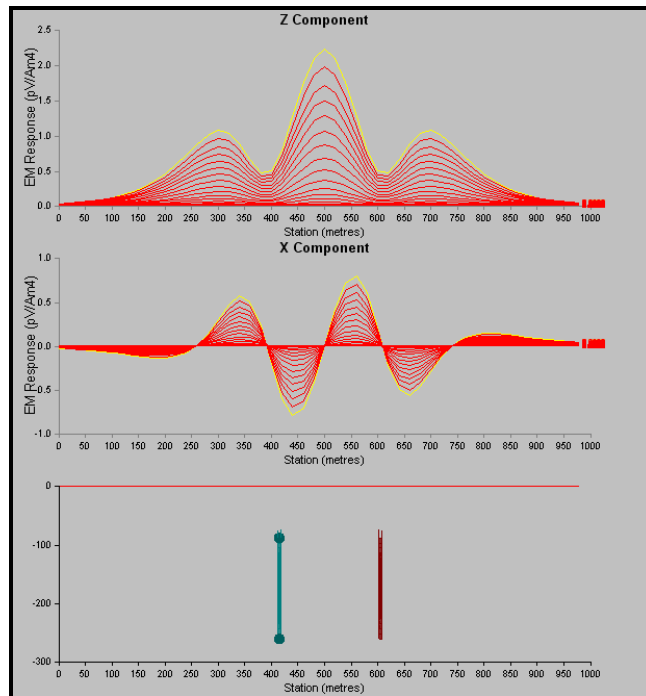


VERTICAL PLATE

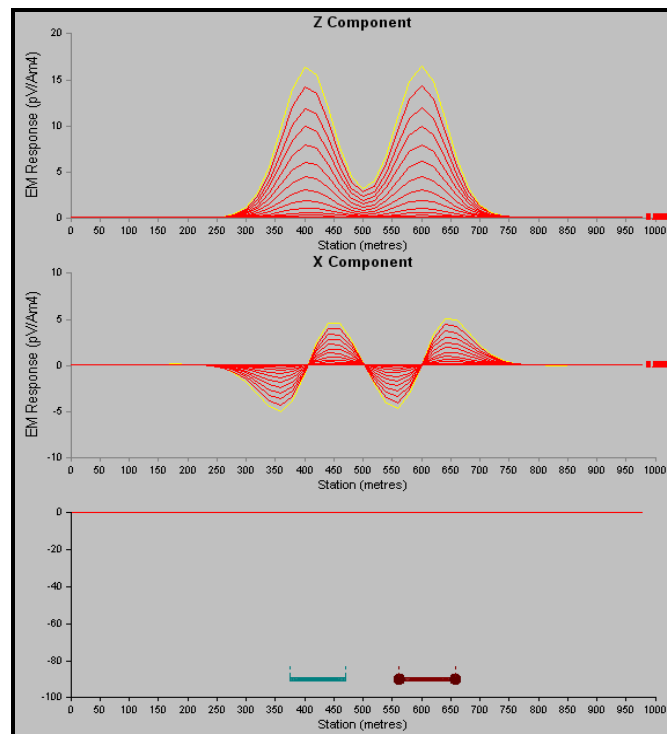


HORIZONTAL PLATE



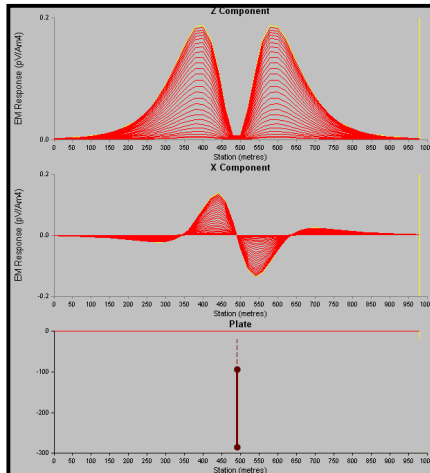


TWO VERTICAL PLATES 200m APART

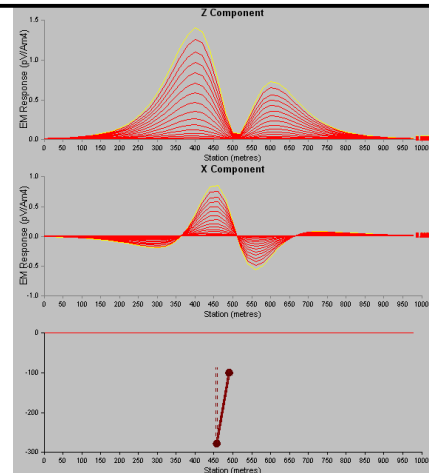


TWO HORIZONTAL PLATES; EDGES 100m APART

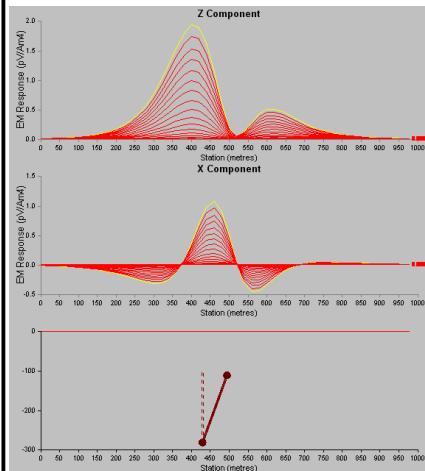




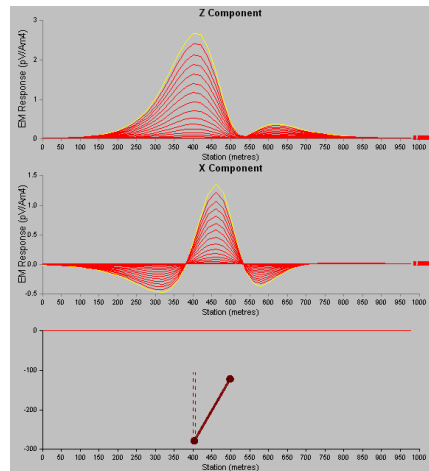
90 degrees



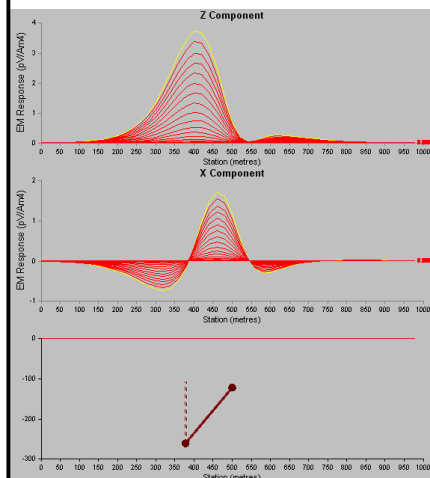
80 degrees



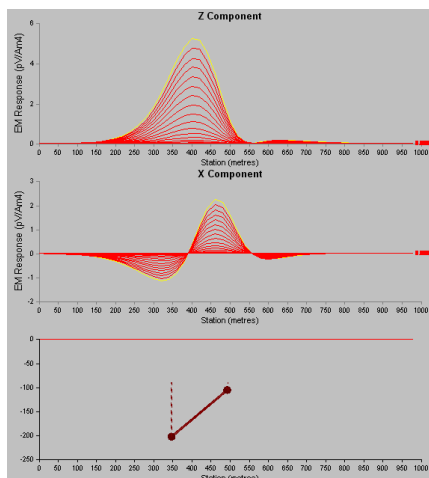
70 degrees



60 degrees



50 degrees



40 degrees

VARIATION WITH DIP



APPENDIX C
GEOPHYSICAL MAP IMAGES
(not to scale)



