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THE 12 MONTHS TO JUNE 17, 1983

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ANNUAL REPORT FOR EL 3478 "WOOLNER" FOR THE 12 MONTHS

TO JUNE 17, 1983.

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SUMMARY

Exploration Licence 3478 of 494.7 square kilometres was granted to Mobil Energy Minerals Australia (M.E.M.A.) on June 18, 1982, with a first year expenditure commitment of \$38,000. The area is about 60 kilometres east of Darwin, and wholly within the Woolner pastoral lease. The licence was obtained to seek Alligator Rivers type uranium-gold deposits, which occur in Lower Proterozoic metasediments near the margins of an Archean granite.

Previous exploration in EL 1642 by CRAE, and a CRAE/Peko-EZ joint venture, had established the existence of the Woolner Granite, of probable Archean age. This was shown to be overlain by graphitic and chloritic schists (the Fish Creek Schists), correlated with the Cahill Formation, and by the Coomalie Dolomite - both of Lower Proterozoic age. All the Precambrian rocks are obscured by an average of 57 metres of Recent and Cretaceous soft sediments. This prior work included drillhole traverses to elucidate the regional geology, and detailed drilling of an area of Fish Creek Schists, as first located by their strong Airmagnetic response. However, the granite margins elsewhere were largely not prospected.

The first year M.E.M.A. exploration programme was planned to confirm the favourable geological setting, and to define the geophysical signature of the schists so that they could be located (under the cover sequence) by ground geophysical surveys. Orientation magnetic and electromagnetic (E.M.) surveys were run on two traverses across the known subcrop of the Fish Creek Schists, and these techniques were then used as a mapping tool over previously untested granite areas, which were then examined by drilling.

The first 12 months work, to July 17, 1983, is reported in three volumes (text, plates and appendices) and includes:

- An evaluation of prior exploration
- . Enhancement of the EL 1642 Airmagnetic Survey data

- . 17 lines of ground magnetic and Gem 8 E.M. surveys, of total length 29.5 kilometres
- . One seismic line of length 4 kilometres
- Drilling 5 rotary/percussion holes, each cored for a short length below the cover sequence, for a total of 491.2 metres.
- . The determination of the age of the Woolner Granite, by uranium:lead ratio's in zircons, as 2675±14 million years.

It was found that the known graphitic beds were weakly indicated by the lowest frequency (32 Hz) radio waves available with the Gem 8 (E.M.) equipment. Reconnaissance magnetic and E.M. traverses located weak anomalies near the Granite rim, which were not found (where drilled) to be due to metasediments.

Further regional geophysical mapping along the Granite margins is planned for the second licence year, using ground magnetic, seismic, gravity, E.M. and downhole E.M. methods.

Expenditure during the first 12 month term, to 17th. June 1983, was \$137,788.

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INTRODUCTION

GENERAL

Exploration Licence 3478 is located approximately 60 kilometres east of Darwin, lying within latitudes 12°15' and 12°27' S and longitudes 131°17' and 131°34' E. Figure 1 shows the location of the exploration licence and its boundaries.

Access to the area is by the sealed Arnhem Highway, then by a graded road to Woolner Station homestead, or by air to the station airstrip.

Roughly 70 percent of the area is black soil plains, swamps and coastal salt pans, subject to seasonal flooding, and to tidal inundation in parts, the remainder being open savannah woodland with some areas of dense scrub. Field work on the black soil plains section is limited to a few months each year.

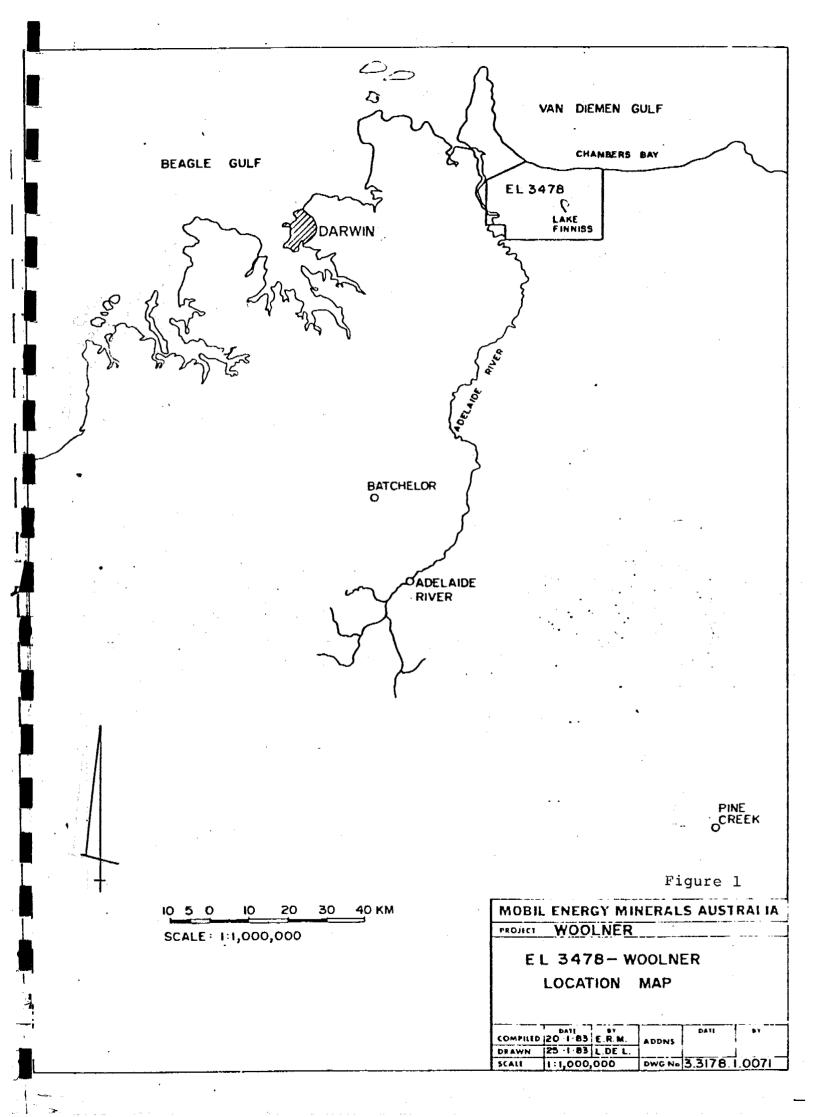
The whole of the licence area falls within PL 793, and it includes a small area of freehold land along the north eastern margin of Lake Finniss. The only prior Mining Title is ML 1298B, which covers a deposit of shell grit.

TENEMENT

EL 3478 was granted to Mobil Energy Minerals Australia Incorporated on June 18, 1982 for a period of twelve months. The area covered is 494.7 square kilometres (191 square miles) with an expenditure commitment of \$38,000.

OBJECTIVES

The licence area was thought to be broadly similar to the geological setting of the Alligator Rivers Uranium Field having a granite complex which was inferred to be Archean, flanked by graphitic and chloritic metasediments with a strong magnetic signature.



Previous exploration by CRAE and a CRAE/Peko-EZ joint venture (in EL 1642) left large gaps, particularly at the base of the metasediments which have not been prospected. A methodical exploration programme was therefore proposed, with three definitive tests to be carried out in the first year.

- (a) Determination of the age and fertility of the granite domes.
- (b) Definition of the magnetic stratigraphy throughout the area, but particularly in the Lower Proterozoic rocks near their base.
- (c) An appraisal of E.M. systems, by which graphitic beds may be detected using conductivity measurements, through the Phanerozoic cover.

Successful completion of these three tests would then lead on to exploration of the granite margins by E.M. surveys followed by drilling to test selected conductive horizons.

PREVIOUS EXPLORATION

Previous exploration by CRAE and a CRAE/Peko-EZ joint venture (in EL 1642) included a detailed airmagnetic survey, close grid ground magnetic and gravity measurements, an unsuccessful trial of ground EM methods, and a major drilling programme, comprising 134 holes for a total of 10059.64 metres.

The drilling defined two granite domes similar to the Nanambu Complex, possibly with a rim of gneisses equivalent to the Kakadu Group. A zone of graphitic and chloritic schists, containing disseminated magnetite and BIF bands, on the western rim of the western dome was thoroughly tested by drilling.

No anomalous radioactivity was found in the 54 holes drilled here, but low grade gold mineralisation (maximum 1.67 parts per million gold over 60 centimetres) was identified in the BIF horizons. These schists are correlated on mineralogy and stratigraphic position with the Cahill Formation. Regional drillhole traverses show that most of the non-granite area is foliated dolomite, probably equivalent to the Coomalie Dolomite. No anomalous radioactivity or metalliferous mineralisation was found in any of these holes.

GEOLOGICAL SETTING

Since the area has a cover, averaging 57 metres thick, of Cretaceous Bathurst Island Formation unconsolidated clayey sands and Recent silt, information on the Lower Proterozoic rocks has been derived from gravity and magnetic characteristics, and on (generally) widely spaced drill holes (see Plate 1).

STRATIGRAPHY

Granites

Three domes of granitoid rocks were recognised by CRAE-Peko-EZ. Subsequent examination of drill core (held at the Department of Mines and Energy) by Dr. John McAndrew of the CSIRO led to the following observations.

- (i) Drill core from the Northwest Dome does not contain any granitic rock, only psammitic schist.
- (ii) Both the West Dome and the East Dome contain foliated to gneissic granites in their central northern parts, a weakly gneissic dark red granite in their adjoining central parts and younger unfoliated two mica granites in their southern parts. These unfoliated granites are pink in the East Dome but grey in the West Dome.
- (iii) Pink pegmatites and associated feldspathization in the older granites may be genetically associated with the younger granites. However, albitisation observed in core of DDH Pll/l (West Dome) may be a separate event.

A copy of Dr. McAndrew's report on this work can be found in the section on Petrology.

The West Dome appears to have a rim of interbanded quartz-feldsparmica gneisses (meta-arkoses) and chlorite-calcite-quartz schists, which may be transitional, with diminishing metamorphic rank, into the overlying Fish Creek Schists and probably correlate with the Kakadu Group in the Alligator Rivers Uranium Field.

A breccia unit, containing clasts of all of the Lower Proterozoic metasediments and the granites, has been described from hole P4/21 on the southern tip of this dome. Examination of cored sections indicates a conglomerate rather than a breccia, with similarities to the Crater Formation of the Rum Jungle area.

Fish Creek Schists

These occur as an isoclinal syncline (see Plate 1) on the western flank of the West Dome, and contain several magnetite rich horizons - thus they have a strong magnetic signature.

The Fish Creek sequence is generally quartz-mica schists, with muscovite and biotite, the latter often altered (by retrograde metamorphism) to chlorite. Plagioclase-rich and calcite-rich bands occur. Some members have fine disseminated graphite, others disseminated magnetite; pyrite is fairly common, while pyrrhotite and chalocopyrite occur occasionally. The sequence includes major amphibolite (hastingsite-albite-epidote-quartz-carbonate) bands, formed by metamorphism of an original carbonate rich sediment.

There are at least six bands of iron formation (BIF) comprising layered magnetite-calcite metaquartzites, in varying thicknesses from 15 centimetres to $1\frac{1}{2}$ metres - these, with the zones of disseminated magnetite, are the source of the prominent and complex magnetic anomaly.

Based on the presence of graphite, disseminated and banded magnetite, carbonate rich bands, and their general pelitic nature, the Fish Creek Schists are correlated with the Cahill Formation.

However, Johnston (1981) has listed the following differences between the Fish Creek Schists and the Cahill Formation:

- (i) There is no basal carbonate unit in the Schist sequence.
- (ii) The Schists contain calcite bands, while the Cahill contains dolomite as the major carbonate.
- (iii) Graphite is disseminated through major thicknesses in the Schists, and not present as finely banded graphite rich and graphite poor beds as in the Cahill.
- (iv) Widespread chlorite in the Schists is the result of retrograde metamorphism of biotite, and not the product of magnesium metasomatism (as in the Cahill ore host).

These differences may explain (at least in part) the lack of uranium mineralisation in the Fish Creek beds.

Coomalie Dolomite

The regional drilling intersected sub-horizontal dolomite in most of the fenceline traverses. A typical section comprises a silicified and cavernous regolith (with highly ferruginous sections and numerous vugs and cavities) in which the effects of weathering diminish with depth to fresh dolomite. This is brecciated in part, but otherwise has a prominent foliation depicted by impure or chloritic dolomite bands. It is correlated with the Coomalie Dolomite. The underlying contact between the dolomite and the Fish Creek Schists is shown by seismic work and two drill holes to be generally horizontal, but locally dipping away from the domes.

The dolomite is generally pale grey to cream coloured, with blobs of fuchsite scattered throughout. It is well foliated, with minor interbeds to about 20 centimetres thickness of talc-chlorite schist. Silicified bands or zones are common throughout; these are not

consistently conformable with the foliation, and are all that remains of the dolomite in the weathered zone.

Core samples of dolomite from three MEMA holes were analysed at Australian Laboratory Services Pty. Ltd. in Brisbane, results are:

<u>Hole No</u> .	% Cao	% Mgo	% S102	% L01	TOTAL
WD1	27.0	19.9	7.99	42.5	97.39
WD2	27.3	20.3	6.59	42.7	96.89
WD2	28.3	20.8	4.54	43.8	97.44

These compare closely with the stoichiometric formula for dolomite (which has 19 to 22 percent Mgo), and indicate that magnesium metasomation has not occurred.

Wildman Siltstone

Previously drilled holes in the far southeast of EL 1642 (and outside the area of EL 3478), intersected colour banded phyllite correlated with the Wildman Siltstone. Cretaceous/Recent cover thickness here is about 100 metres, or more.

Phanerozoic Cover

Bathurst Island Formation forms a capping on all of the older rocks in EL 3478, and is generally unconsolidated clayey sands, which become feldspathic sands at depth. The "porcellanite" claystone bed typical of this Formation is thin — a metre or two — or missing. Correlation with the (Cretaceous) Bathurst Island Formation seems unequivocal. The base of this unit frequently contains a pebble bed, in which the pebbles can be recognised as silicified dolomite.

The unit is variable in thickness - from nil to 138 metres, with average 41 metres.

Recent alluvium comprising a layer of black mud, laterite and dune sands containing shells and plant remains, up to 44 metres thick, is the uppermost bed in this area. The average thickness is 16 metres.

The total cover thickness ranges from 15 to 140 metres, with an average of 57 metres. Weathering in the preCambrian rocks averages about 10 metres. Plate 2 shows rough isopachs of cover thickness, and indicates the granites are topographic highs, with a narrow valley along the granite-dolomite margins, somewhat similar to the present surface topography in the Rum Jungle district.

STRUCTURE

Little precise information can be derived for the pre-Coomalie Dolomite Lower Proterozoic beds. The layering of the BIF units, the formation contacts, and the schistosity all appear to dip steeper than 70°, and all are probably isoclinally folded.

METAMORPHISM

The gneisses of the Woolner Domes are of epidote - amphibolite facies, and commonly show two deformations: the first producing the mineral assemblage, and the second the gneissic fabric. On the margins of the Domes, in the transitional area, the gneisses are of upper greenschist rank.

The Fish Creek Schists and Coomalie Dolomite are mildly metamorphosed sediments, initially chemical sediments and siltstones. Biotite is incompletely altered to chlorite in some cases, and there is no evidence of magnesian metasomatism in the schists.

GRANITE STUDIES

Small sections of drill core from the EL 1642 (CRAE and CRAE/Peko-EZ) holes were collected (from the N.T.G.S. core reference store in Darwin) by Dr. John McAndrew of the CSIRO Institute of Energy and Earth Resources. This material was then used for studies on the petrology, fertility (i.e. the potential of the granite as a source of uranium for redistribituion or enrichment into adjacent ore deposits), and the age of the granite. The results of these studies are reported below.

PETROLOGY

These cores consist of a short length (1/2 to 2m) of granitic rock from the subcrop of a dome. A few longer cores through the metamorphic rocks were briefly examined. Overall, core was examined from 28 drill holes, and small samples taken for further study from 22 of these, 19 being of granitoids and pegmatites and 3 of metamorphic rocks (Table 1). The granite samples are adequate for study of the composition of accessory zircon, and to search for accessory uraninite in an assessment οf their potential as source rocks for mineralization. A number are of sufficient size for meaningful chemical analysis.

None of the core examined contained more than 2kg of granitoid. With half the core being available for study, no more than 1kg is available from any single core, and for most cores less than 3/4kg.

Granites

The drill core from the "Northwestern Woolner Dome" does not contain any granitic rock. The grey psammitic schist of this dome is notable for an extensive length with red-brown hematitic staining in DDH P10/8 on the NW of this dome.

The Woolner West Dome and Woolner East Dome drill cores contain a number of types of granite with pegmatite. These may include granites of two ages; however, age dates (this study) indicate no distinguishable age differences in the limited sampling.

The central northern parts of both domes contain foliated to gneissic granites. A connection between the two domes is provided by a distinctive weakly gneissic dark red brown granite, which forms the central western margin of the Woolner East Dome, and on the adjoining central eastern margin of the Woolner West Dome.

The southern part of each of these domes contains granites with little foliation. These, by inference from this feature, are somewhat younger than the northern granites. In both domes they are two-mica granitoids, with the grey granitoid of the souther Woolner West Dome being distinctive from the pink granitoid of the southern Woolner East Dome.

Pink pegmatites and associated feldspathization in the "older" foliated granites may be genetically associated with "younger" granites. However, the albitization observed in core of DDH P11/1 may be a separate event.

There is appreciable sericitization in some core along shears or parallel fractures. While such muscovite is unambiguously secondary, it is not yet clear if the two-mica granites in the southern part of the two domes contain primary as well as secondary muscovite.

Woolner West Dome

Drill cores from the northern part consist of foliated to gneissic grey granite, intruded by pink potash feldspar pegmatite with associated development of potash feldspar in the adjoining granite. Distinct to the latter is the irregular albitization of foliated grey biotite granite in the northern drill hole Pl1/1.

While the south central drill hole WORD 1 penetrated pink feldspar pegmatite, drill holes south-west and east of this are in grey two-mica granite. The lack of foliation in this granite suggests it is younger than gneissic granites in the northern part of the dome.

The most easterly drill hole Pl3/6 is in deformed pink granite correlated with that in the adjoining part of the Woolner East Dome.

Woolner East Dome

The most northerly drill core examined, Pl4/7, is in pink feldspar pegmatite intruding pelitic schist.

Two cores in the eastern part consist of biotite-muscovite schistose granite or granite-like psammitic schist. Disseminated pyrite crystals are present. A third drill core (P14/6) is sheared granite with abundant sericite along parallel shear planes.

The western part is also an "older" granite, being more melanocratic, relatively dark red brown, and having a slightly gneissic texture. As noted above, this may correlate with granite in the adjoining part of the Main Dome.

In contrast, the southern part of the Woolner East Dome is apparently a younger granite, having but slight foliation. It is a two-mica granite, generally with a substantial proportion of pink feldspar.

Description of Drill Core

TABLE 1 (from Woolner domes, N.T., examined at Darwin core store of NTGS, 19th-20th July 1982, and core samples obtained).

DDH	Location in dome		nples n) CSIRO	No.	Comments
NORTHWE	ESTERN WOOL	NER DOME			
P10/8	N	57.9	70496		hematite-stained psammitic schist
		59.4	70497		hematite-stained psammitic schist
P10/9	NE	-			two-mica psammitic schist
P10/18	E	56.8	70498		grey psammitic schist with small white albite porphyroblasts
P10/7	S	•-			grey psammitic schist
WOOLNER	R WEST DOME				
P11/1	NC	70.0	70478		irregularity albitized foliated biotite granite, with sericitic shears
P6/3	NW	-			feldsparthized biotite gneiss
P7/1	WNW	66.7	70495		dark grey psammitic schist intruded by quartz pink Kspar pegmatite with Kspar augen adjoining the pegmatite
P1/3	W	55.8	70474		dark grey feldspathized gneiss (? deformed granite) cf WP2/S1
WORD 9	W	58.3	70475a)	medium grey sheared gneissic granit
		58.8	70475ъ)	with quartz pink microcline pegmatite and local irregular pink feldspathization of
		59.3	70475c	5	granite
P13/6	E	67.7	70477		melanocratic pink biotite granite, deformed and ? porphyritic.
WP2/S1	WSW	293	70490)	dark grey quartz white-feldspar
·		299.8	70491))	meta-arkose to psammitic schist wit irregular bands and patches of metasomatic pink feldspar
WORD 1	sc	51.0	70473		pink feldspar pegmatite with quartz veins

Description of Drill Core

TABLE 1 (Cont.)

DDH	Locati		Samples	
	in don	ie Dept	h (m) CSIRO	No. Comments
WB2	SE	50.0	70476	grey two-mica granitoid with but slight foliation or deformation
P4/5	S	49.0	70472	weathered schist and weathered grey sheared granite
P4/2	S		-	light grey weathered granite
P4/1	S	37.5	70471	weathered light grey two-mica granite
WOOLNER	EAST D	OME		
P14/7	N	77.6	70493	quartz pink-felspar pegmatite intruding pelitic schist
		80.6	70494	quartz pink-albite pegmatite
P14/1B	NW	76.50	70492	gneissic melanocratic granite similar to P13/6 with some pink felspar adjoining psammitic schist
P14/3	WNW	64.80	70485	light grey psammitic muscovite-biotite schist (cf. 70478) which could be schistose granite. Has euhedral pyrite (cube + octahedron)
		69.30	70486	do
P14/4	Ene	59.0	70487	well foliated to schistose relatively fine grained granite schist with more biotite than 70485
P14/6	E	54.6	70484	sheared pink leycocratic granite with sericitized shear planes
p14/11	SW	72.7	70488	red-brown two-mica granite, slightly gneissic, cf 70492
		73.6	70489	similar gneissic dark red-brown granite with occasional white feldspar
P14/8	SE	52.20	70482	weakly foliated off white two-mica granite with some pinkish brown feldspar (cf WB/2 as well as 12/11)
		53.50	70481	weakly foliated relatively fine grained granite

Description of Drill Core

TABLE 1 (Cont.)

DDH	Locat in do		Samples (m) CSIRO	No. Comments
P12/1	SW	55.5	70480	very leucocratic pink sheared granite with sericitic shears
P12/7	S		·	fractured sheared slightly weathered pink leucogranite
P12/10	S	_		fractured quartzite
P12/11	s	52.6	70489a	unfoliated pink two-mica granite
•		53.1	70479ъ	do

FERTILITY INVESTIGATION

The Woolner domes were thought to be possible sources of uranium which may have been subsequently transported and concentrated in Carbonaceous Lower Proterozoic hosts. The potential fertility of these granitic domes has been evaluated by Dr. John McAndrew; the following interpretations and comments are those of MEMA utilizing the results of Dr. McAndrew, and do not necessarily reflect his thoughts.

The northern granites, types A and B, are most metamorphosed to schistose granite. The central red granite, type C, has thorough alteration of potash feldspar to granular albite and quartz. This albite is red from hematite dust and lacks muscovite to saussurite alteration of magmatic plagioclase.

The Southern granites, types D and E, are microgneiss and lack the stronger dynamic deformation of northern granites. There are occasional residuals of more calcic unsaussuritzed plagioclase within some of this magmatic feldspar, otherwise they are now albite with muscovite with minor eipdote.

Zircon is mainly very fine grained, prismatic to elongate prismatic (early zircon habit) with strike zoning throughout. They are magmatic, and included in felsic minerals without any concentration into biotite. This resembles the occurrence in the Nanambu Complex, and are different to those in the Litchfield Complex granitoids. Most specimens are low in the phosphate accessory minerals, apatite and monazite.

The Woolner granites were originally peraluminous ilmenite-series granitoids, very low in opaque accessory minerals. There is occasional pyrite. Ilmenite is absent except for residuals in the southern granites which have less metamorphism and metasomatism. However, leucoxene pseudomorphs, with biotite, representing altered ilmenite, is present.

The relatively high uranium in zircons suggest some granites may have crystallized as fertile source rocks. The low overall whole rock uranium (except for Word 9, and to a lesser extent in P4/1. P14/1) may also suggest potentially labile uranium may have been mobilized during metamorphism and metasomatism. Schemes are plausible:

- 1. The postulated mobilising fluids may have been oxidising, a judgement possible from the oxidation of ilmenite to leucoxene.
- 2. An alternative possibility suggested by pyrite and leucoxene after ilmenite is an extended "pronto reaction" under reducing conditions. Unfortunately there is no mention of trace brannerite in the petrographic reports. The suggested reaction (not balanced) is:

$$S + Fe Ti O_3 = Fe S + TiO_2$$

 $UO_2 + 3 Ti O_2 = U Ti O_8$ (brannerite)

The corollary to the latter reducing style alteration is the uranium should have been retained in the system.

- 3. One may also question if the high uranium content of some zircons is an inherited trait, and bears no relation to the moderately low uranium content of host granitic rocks.
- 4. Finally, there may have been little uranium lost from the system. The estimated partition coefficient of uranium into zircon, using bulk rock uranium as an approximation of the magma uranium content, is roughly 250-1000. This is not exceptionally high, particularly in the absence of other concentrating accessory minerals such as apatite and monazite.

The Woolner Granite may differ from the Nanambu Complex, where metamorphism under reducing conditions is evident by overall retention of ilmenite, with only slight development of rutile together with retention of accessory uraninite through the metamorphism.

Table 2:	Uranium, thorium and the Woolner domes (content of gran	ites from
<u>Type</u>	Location (DDH)	<u>u</u>	<u>Th</u>	Zr
A	P 11/1 a P 11/1 b	3 4	4 3	72 64
A	Word 9	11	13	149
. В	P 14/2	2	5	102
C C	P 13/6 P 14/1	2 5	3 5	106 64
C	P 14/11	2	4	193
D	WB 2	3	. 8	101
D	P 4/1	6	6	90
E	P 14/8 a P 14/8 b	2 2	2 6	32 120
E	P 12/11	2	4	55

Table 3:	Uranium content	of zircons,	Woolner	granitoid A	
	mmer m11/1				

	DDH P	11/1	Word 9		
Crystal		Content (%)	Uranium Com		
	<u>interior</u>	outer part	interior or	uter part	
1	.030	.246	.305	.240	
2	.190	.078	.254	.155	
3	.078	.062	.439	.407	
4	.205	.045	.102	.104	
5	.0331	.069	1.021	.164	
6	.154	.143	.554	.164	
7	.035	.148	.279	.225	
8	.148	.153	.333	.324	
9	.200	.206	.406	.270	
10	.151	.097	.485	.203	
11	.109	.094	.146	.171	
12			.912	.910	
			.179	.237	
			.181	.270	
1,		001			

¹Intermediate part

.024

Table 4:	Uranium co	ontent of	zircons,	Woolner	granitoid C	
----------	------------	-----------	----------	---------	-------------	--

	DDH P			P 14/1B	DDH P 1	4/11
Crystal		Content (%) outer part		Content (%) outer part	Uranium Con interior ou	
						F
1	.030	.104	.073	.044	.146	.306
2	.082	.121	.036	.063	.075	.034
3	.117	.126	.045	.059	.041	.034
4	.178	.305	.103	.030	.076	.196
5	.211		.059 ¹	.066	.047	.038
6	.109	.144	.083	.065	.383	.302
7	.149	.125	.051	.030	.110	.095
8	.120		.259	.188	.214	.097
9	.104	.790	.052	.030	.204	.182
10	.088	.133	.067	.086	.181	.053
11	.056	.106	.049	.116	.082	.110
12	.246	.129	.031	.030	.072	.078
13			.049	.069	.166	.271
14					.068	.162
15					.102	.098

¹ Intermediate part .030

Table 5:	Uranium Content of zircons, Woolner granitoid D						
Crystal	DDH V Uranium Co interior		DDH P4/l Uranium Content (%) interior outer part				
1	.089	.096	.206	.190			
2	.203	.080	.249	.246			
3	.063	.117	.234	.208			
4	.131	.092	.262	.277			
5	.230	.110	.142	.208			
6	.093	.119	.172	.332			
7	.161	.205	.030	.065			
8	.055	.030	.365	.193			
9	.082	.095	.155	.243			
10	.087	.079	.277	.678			
11	.055	.030	.180	.235			
12	.198	.261	.194	.209			
13	.084	.114	.189	.141			
14	.369	.392	.184	.133			

Table 6: <u>Uranium content of zircons</u>, Woolner granitoid E

Crystal		ontent (%) outer part	DDH P12/11 Uranium Content (%) interior outer part
1	.0331	.153	.063 .207
2	.3212	.139	.116 .282
3	.1943	.213	.103 .111
4	.042	.041	.030 .032
5	.217	.072	.035 .084
6	.179	.336	.129 .030
7	.453	.366	.079 .320
8	.314	.240	.296 .657
9	.356	.240	.439 .270
10	.202	.250	.122 .175
11	.143	.077	
12	.364	.366	
13	.046	.178	
1 Interme	diate nart	030	

¹ Intermediate part .030

² Intermediate part .193

³ Intermediate part .129

AGE DETERMINATION

Individual zircon grains were separate from the granite matrix from (drillcore holes P4/1. P11/1, P 12/11 and P14/1) by Dr. McAndrew. The isotopic composition of the lead, and the lead/uranium ratio in the zircons, were then measured by Dr.'s I.S. William and W. Compston of the Australian National University, using the "SHRIMP" ion microscope. Their report is attached as appendix 2, the results of this work are summarised below.

Two of the samples, P12/11 and P14/1. contained zircon with U-Pb compositions suitable for precise age measurements. These gave ages of 2660 ± 20 and 2690 ± 20 Ma, respectively. The compositions of the zircons from the other two samples were not suitable for measuring an age independently but were wholly consistent with these ages. As there is no evidence for a significant difference in age between the samples, the best estimate of the age of all is the mean of the above, 2675 ± 14 Ma.

AIRMAGNETIC DATA

The airborne magnetic survey has been enhanced to give added knowledge of the Woolner Dome. A special analysis of the data was carried out and matched filter designed to display the regional and residual components. These maps were then reduced to the pole to remove lows produced by the low angle of the earths magnetic field.

OBSERVED MAGNETIC TOTAL INTENSITY CONTOUR PLAN (Plate 3)

The most pronounced magnetic feature of the map is a north to north easterly trending magnetic high of 20 to 60 nanoteslas extending from the SW corner of the data sheet MO403 to offshore.

Less pronounced features are two north easterly trending magnetic highs situated on map sheets MO601 and MO603.

The least pronounced feature of the plan is the change of magnetic character on the northwest portions of maps sheet MO405 and MO304. This pattern extends northeast out into Van Dieman Gulf.

The southern portion of the plan is dominated by a regional gradient increasing by 6 nanotesla (nT) per kilometre to the southwest.

RESIDUAL MAGNETIC TOTAL INTENSITY CONTOUR PLAN (Plate 5)

The Residual Magnetic Total Intensity Contour Plan has four anomaly zones similar to the observed magnetic contour plan.

The first zone is located along the western boundary of map sheet MO403. The readings range from a background of -10 to -80 nanoteslas (nT) away from the feature to 0 to 140 nanoteslas over the feature.

The second and third zones are located on map sheets M0601 and M603. These anomalies trend north east and have limited lateral extent.

The fourth zone is positive magnetic anomaly of 60-120 nT located in the north west corner of the sheet. The anomaly of large surface extent extends out into Van Dieman Gulf.

REGIONAL MAGNETIC TOTAL INTENSITY CONTOUR PLAN (Plate 4)

The regional map has outlined a central magnetic low (769E 8638N) surrounded on the southern and western edges by magnetic positive values. A pronounced regional magnetic high is located at the junction of the two trends (e.g. 758E 8620N).

A distinguishable deviation of the magnetic contour can be seen over the residual magnetic anomaly zone on the western boundary of map sheet M0403.

REGIONAL MAGNETIC TOTAL INTENSITY POLE REDUCED TO THE POLE (RTP) CONTOUR PLAN (Plate 6)

This contour plan is the reduction to the Pole of the regional magnetic data using the Pole formula (Syberg, 1972).

The transformation moves the main magnetic feature south to 769E 8634N. The elongation to the east on the original map is not apparent on the RTP map with the feature becoming more circular with a positive value of 6 nanoteslas. The presence of the limb to the south east is noticable on this plan. The small anomaly on the west side of the dome has been suppressed.

The magnetic rim features are also moved south with a value minima of -8 nanoteslas.

REGIONAL MAGNETIC DIPOLE REDUCED TO THE DIPOLE CONTOUR PLAN (Plate 7)

This transformation has moved the central magnetic feature further south to 760N 8630E. The intensity of the anomaly is -7 nanoteslas.

A small deep seated magnetic feature is located on the western edge of the anomaly. The rim features are not clearly outlined. The strongest anomalies of +11 nanoteslas in the north west corner and 9 nanoteslas in the south west corner which indicate change of basement magnetic susceptibility.

GROUND GRAVITY SURVEY

OBSERVED BOUGUER GRAVITY CONTOUR MAP (Plate 9)

The Observed Bouguer Contour Map has been constructed from 131°15'E to 131°36'E from 12°19'S to 12°34'S from data in open file.

The dominant feature on this plan is a gravity low consisting of two minima. The major minima is centred on 764E 8628N with the smaller feature centred at 771E 8624N. The steepest gradient of its fracture is located on the western edge.

The gravity pattern around the central gravity low of -11 milligals has different characteristics to the east and to the west. The eastern portion is a plateau gravity high with a number of small gravity closures. The western boundary gradually decreases in gravity value and has a small number of noticable gravity features.

REGIONAL BOUGUER GRAVITY CONTOUR PLAN (Plate 10)

The principal feature of the Regional Bouguer Gravity Contour Plan is the central gravity low which has partly removed the distortion produced by a second minimum to the south west of the major minimum. This map indicates that the feature has its deepest section below the western gravity minimum and plunges south east. This pattern is very similar to the Regional Dipole RTP Total Intensity Contour map with the magneitc pattern displaced to the north east by 3.5 kilometres.

RESIDUAL BOUGUER GRAVITY CONTOUR PLAN (Plate 11)

A residual plan was constructed by subtracting the regional value at the station from the observed gravity value.

The most significant features are two residual gravity lows of 3.0 and 3.5 milligals centred at 8623N 762E and 8622N 772E.

These two features are separated by and surrounded by gravity highs ranging in intensity from 1 to 4.5 milligals. The strongest intensity gravity highs are located along the southern boundary of the central gravity lows.

The regions to the north east corner and western boundary of the sheet exhibit small positive and negative gravity residuals.

Low order north south trending zones are located to the south of the ring gravity highs on the southern boundary of the area.

GROUND ELECTROMAGNETIC (E.M.) SURVEYS

These surveys are covered by Appendix 1. The Gem 8 E.M. method proved capable of penetrating the conductive overburden with the lowest frequencies 41 HZ and 82HZ. The Fish Creek schists graphitic areas and pyritic units on the north western edge of the dome could be located.

SEISMIC SURVEY

A limited seismic refraction/reflection survey of 4 kilometres in length was conducted along the northern edge of Lake Finniss. The purpose of this survey was to provide an estimate of the thickness of carbonate rocks lying between the Woolner East and Woolner West granite domes.

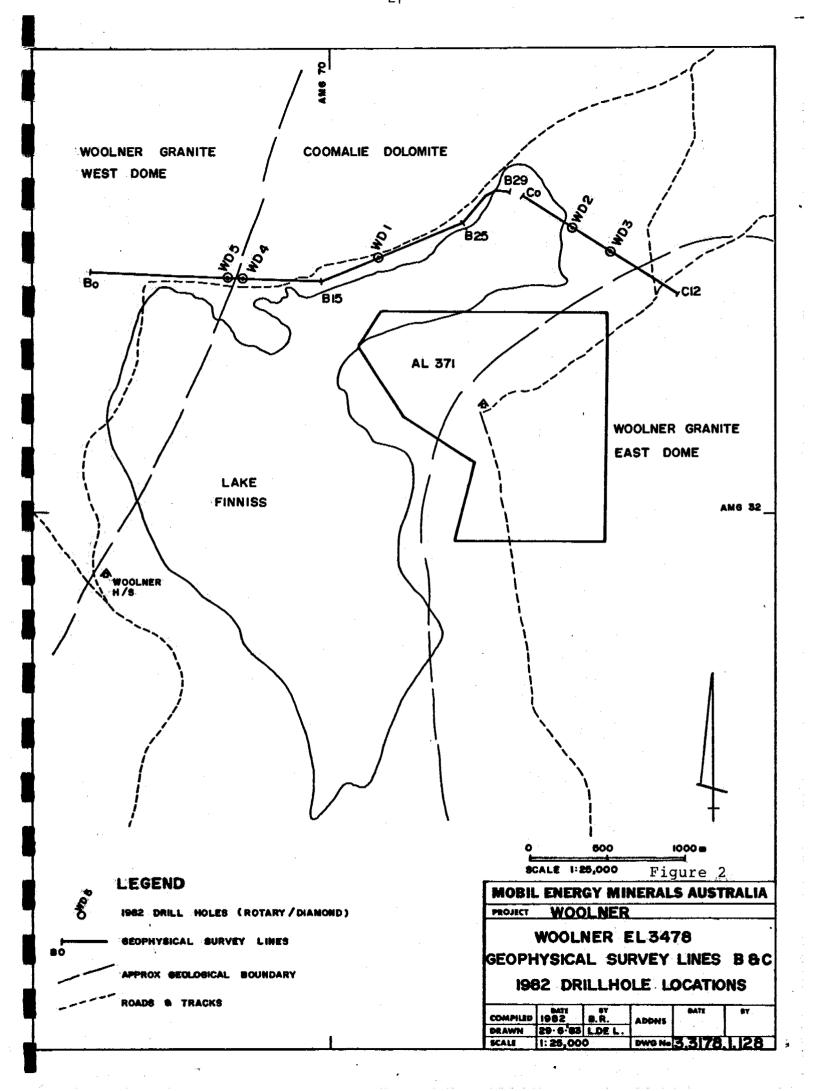
Carbonate sequences are not good transmitters of energy and have high velocity which shield structure below the sequence from being mapped by this method.

The sequence at Woolner has proved to be the normal case and clear evidence of reflections and deep refraction layers has not yet be achieved.

The Seismic data has been filtered with 100 cps band pass filter and normal move out conversions have been applied. A band of reflections with a two way time of up to approximately 200 milliseconds has been located.

It is considered that this reflector is the basement/Proterozoic contact. If a velocity of 3 kilometers per second is assumed for the section. A depth of 350 metres is estimated in the sedimentary section of the Embayment.

The results of the seismic survey are presented in Volume II.



DRILLING

The drilling programme was designed to test below the Phanerozoic cover for the prospective Fish Creek Schists in the corridor of metasediments between the East and West granite domes. Five rotary/diamond holes were drilled on geophysical or stratigraphic targets, but only Coomalie Dolomite and Woolner Granite were intersected (Figure 36). Each hole was precollared through the Phanerozoic sediments (40 metres) and the four holes intersecting dolomite were cored (83 metres) (Figure 3 and Table 7). The core was tested every 50 centimetres using an URTEC spectrometer and no anomalous radioactivity was recorded.

A ground magnetic and EM survey was conducted along drill lines B and C and the holes WD2, 3, 4 were sited on resulting EM anomalies, WD 1 on a weak magnetic anomaly and WD5 positioned between WD4 and the then known eastern boundary of the Woolner Granite.

The Coomalie Dolomite is generally pale grey to cream coloured, weakly foliated and contains thin schist interbeds. The foliation dips between 10° and 30° to the axis of the core and the dolomite is partially silicified. In WD5 over 10 metres of weathered granite was intersected.

The thickness of the Phanerozoic cover varies between 56 and 83 metres and consists of the Recent dune, beach and lake deposits with underlying kaolinitic sands, clays and minor carbonates of the Bathurst Island Formation. The full sequence of coastal sands, estuarine organic muds and shell fragments, the laterite profile and the Bathurst Island Formation is best shown in hole WD4 (see drill log in Appendix 3).

Table 7	Summary	of	Rotary/Diamond	Drilling	-	Woolner	EŁ	3478

Commencement Date : November 10, 1982 Completion Date : November 24, 1982

Depth Drilled (m)

	Local		-		
Hole No	Co-ord	AMG Co-ord	Rotary	Diamond(NQ)	Lithological Summary
WD1	B19	33.68N/70.30E	0-62	62-82.2	Phanerozoic sediments over-lying Coomalie Dolomite
WD2	C4	33.85N/71.60E	0-57	57-81.3	Phanerozoic sediments overlying Coomalie Dolomite
WD3	С7	33.72N/71.82E	0-56	56-75.3	Phanerozoic sediments overlying Coomalie Dolomite
WD4	B10	33.55N/69.45E	0-83.2	83.2-102.4	Phanerozoic sediments overlying Coomalie Dolomite
WD5	В9	33.55N/69.35E	0-150	-	Phanerozoic sediments overlying Woolner Granite
			408.2	<u>83</u>	

CONCLUSIONS

The age of the Woolner granites has been evaluated at 2675 ± 14 Ma. This compares favorably with the age of the Nanambu Complex. The fertility study is inconclusive as to whether labile uranium was present or was remobilized during the events of metamorphism and metasomatism.

Study of the drill core available at the Department of Mines and Energy in Darwin has established a band of interbanded quartz-feldspar -mica gneiss and chlorite-calcite-quartz schist which may be equivalent of the Kakadu Group of the Alligator Rivers Uranium Field.

No indication of magnesium alteration has been located by the MEMA drilling and the differences between the Fish Creek Schists and the Cahill Formation (Johnson 1981) still apply.

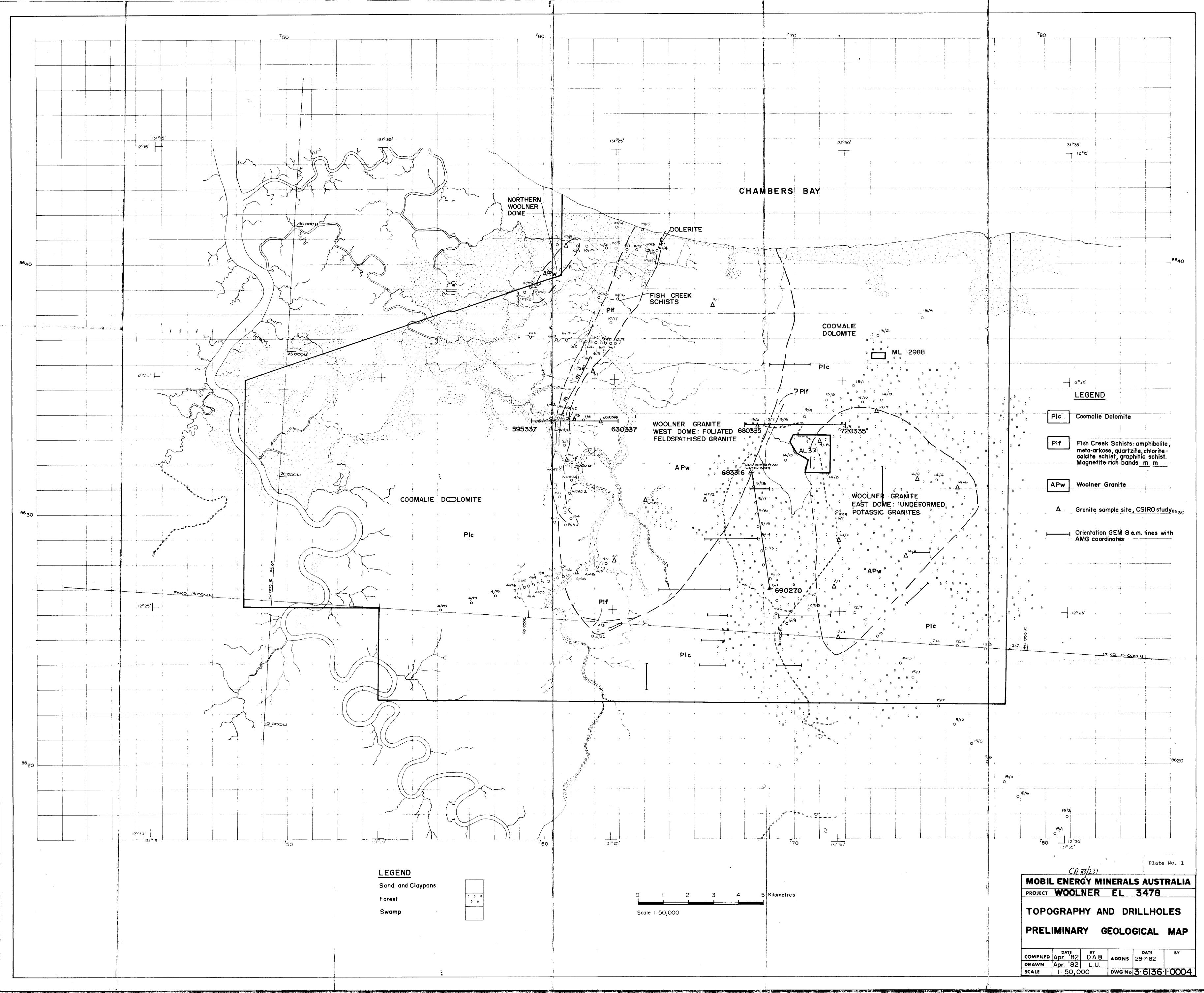
The enhanced treatment of the airborne magnetic survey has indicated only small discontinuous magnetic features associated with the embayment area between the Woolner West Dome and the Woolner East Dome. It is therefore considered unlikely that any large area of B.I.F. sediments occurs around the contact zone. However the magnetic treatment does not rule out the possiblity of a non-magnetic equivalent of the Fish Creek schist being present.

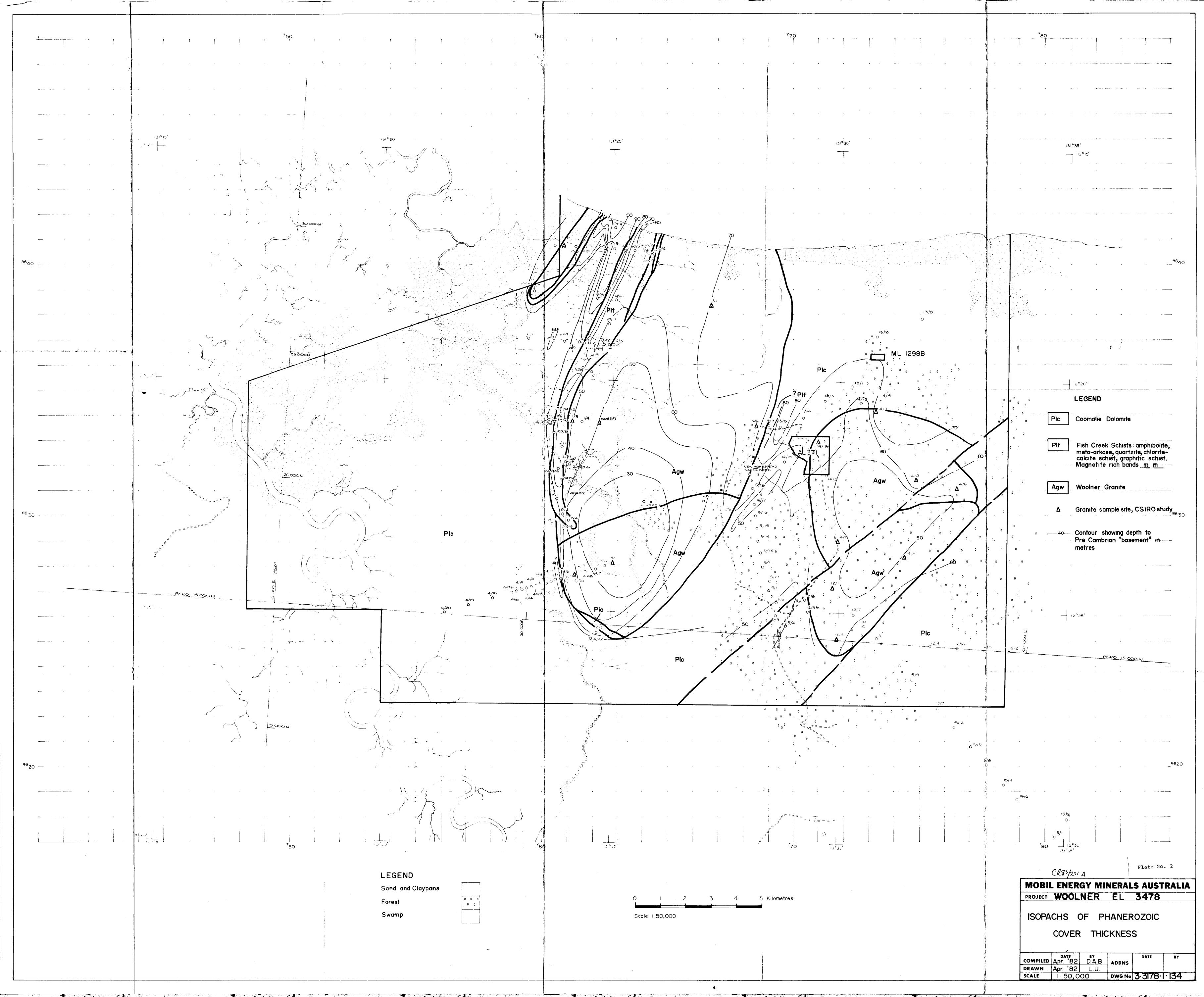
The ground EM surveys have been located on the eastern edge of the Woolner West Dome. However no major EM conductor has been delineated. The EM conducted over the known geology indicated that the granite schist contact could be mapped. However no major anomaly was associated with a graphite zone located by drilling. This indicates that the Fish Creek Schists are possibly different facies to the Cahill Formation where the EM conductors are more clearly defined.

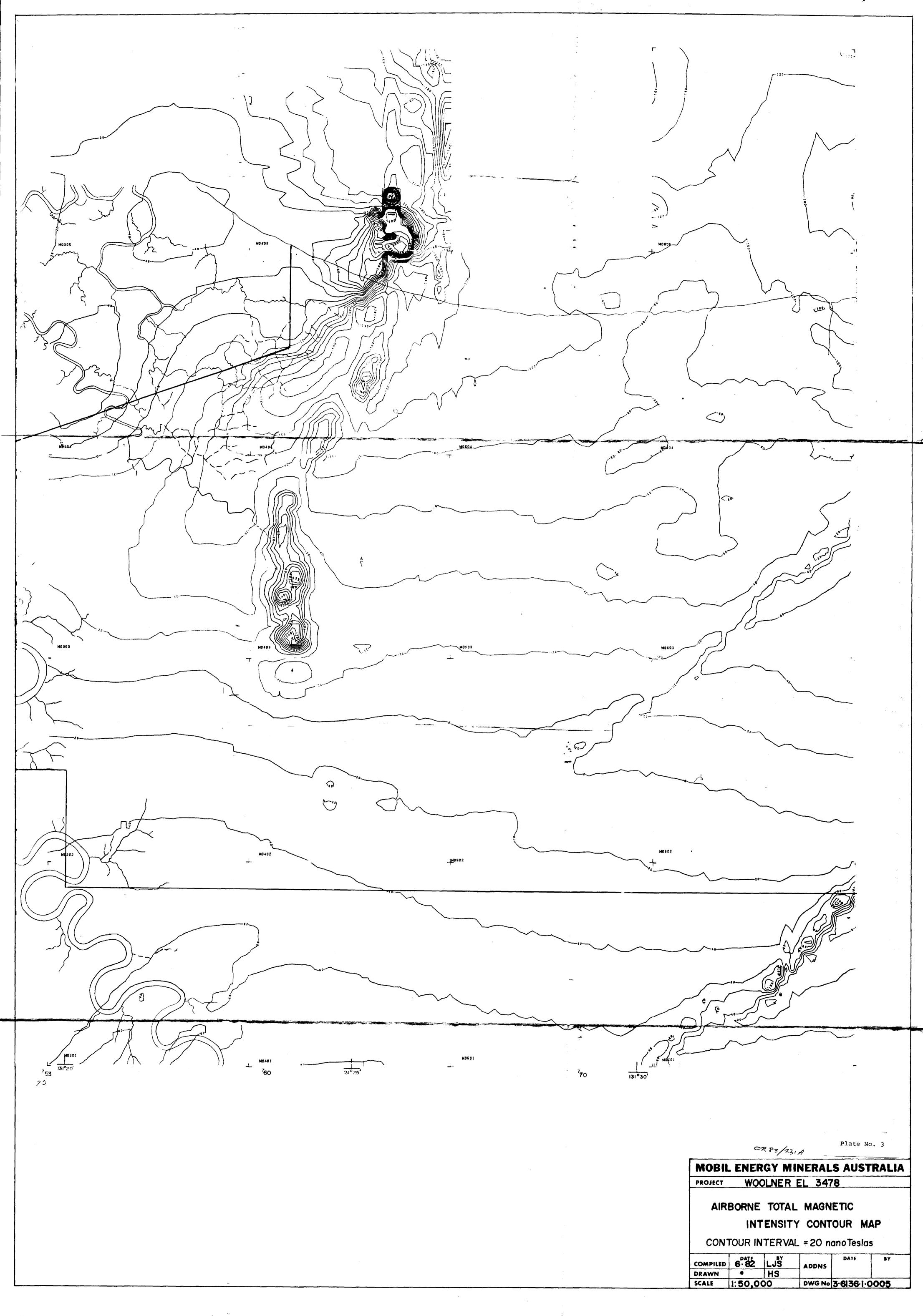
It is recommended that no further investigation be carried in the Embayment area as EM anomalies are poorly defined and there was no

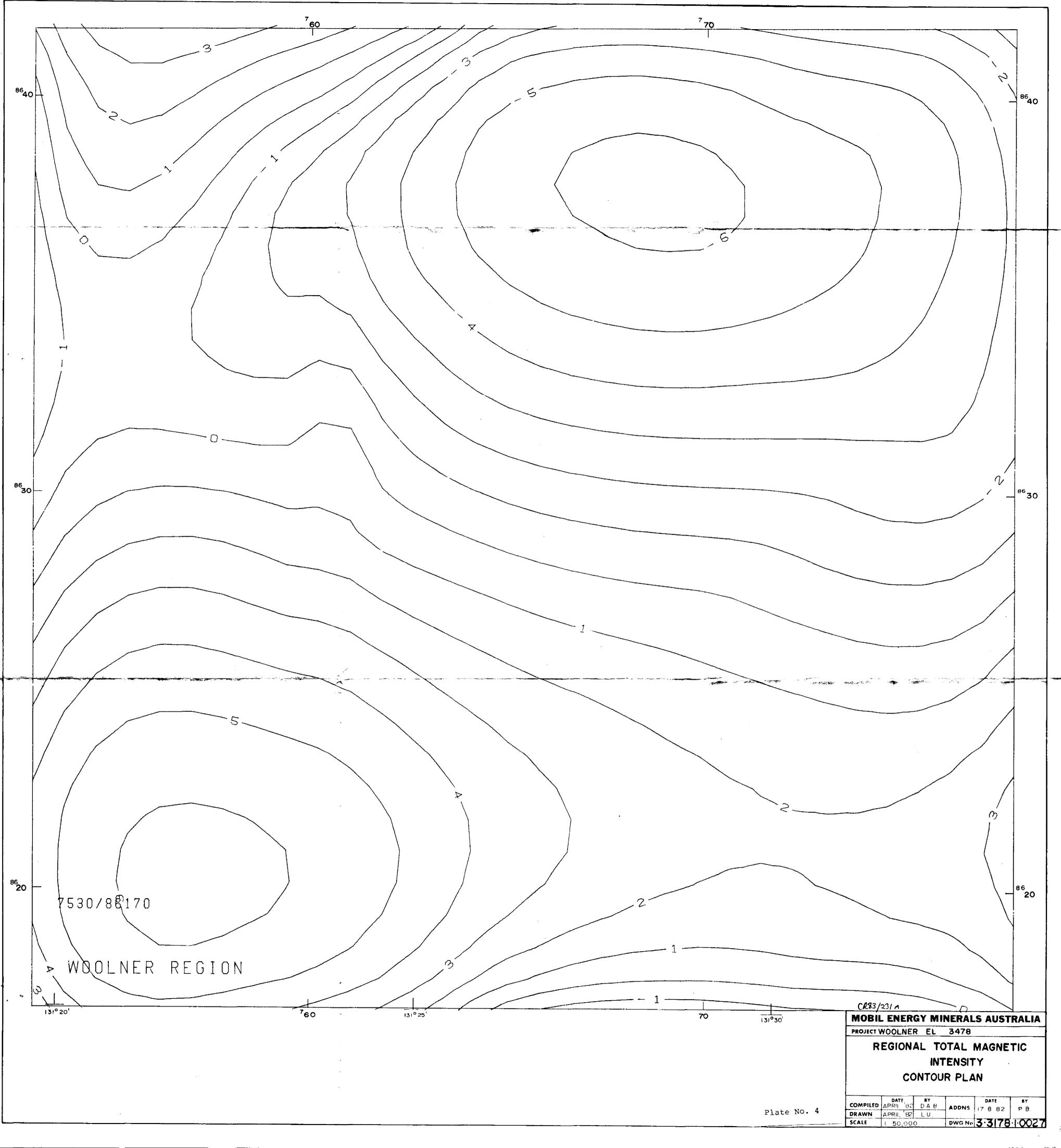
continuous magnetic anomaly. The seismic survey also indicates a sedimentary sequence of possible 350 metres above the granite which could consist completely of Coomalie Dolomite. There is no suggestion from the seismic survey of a section of Fish Creek Schists located below the dolomite.

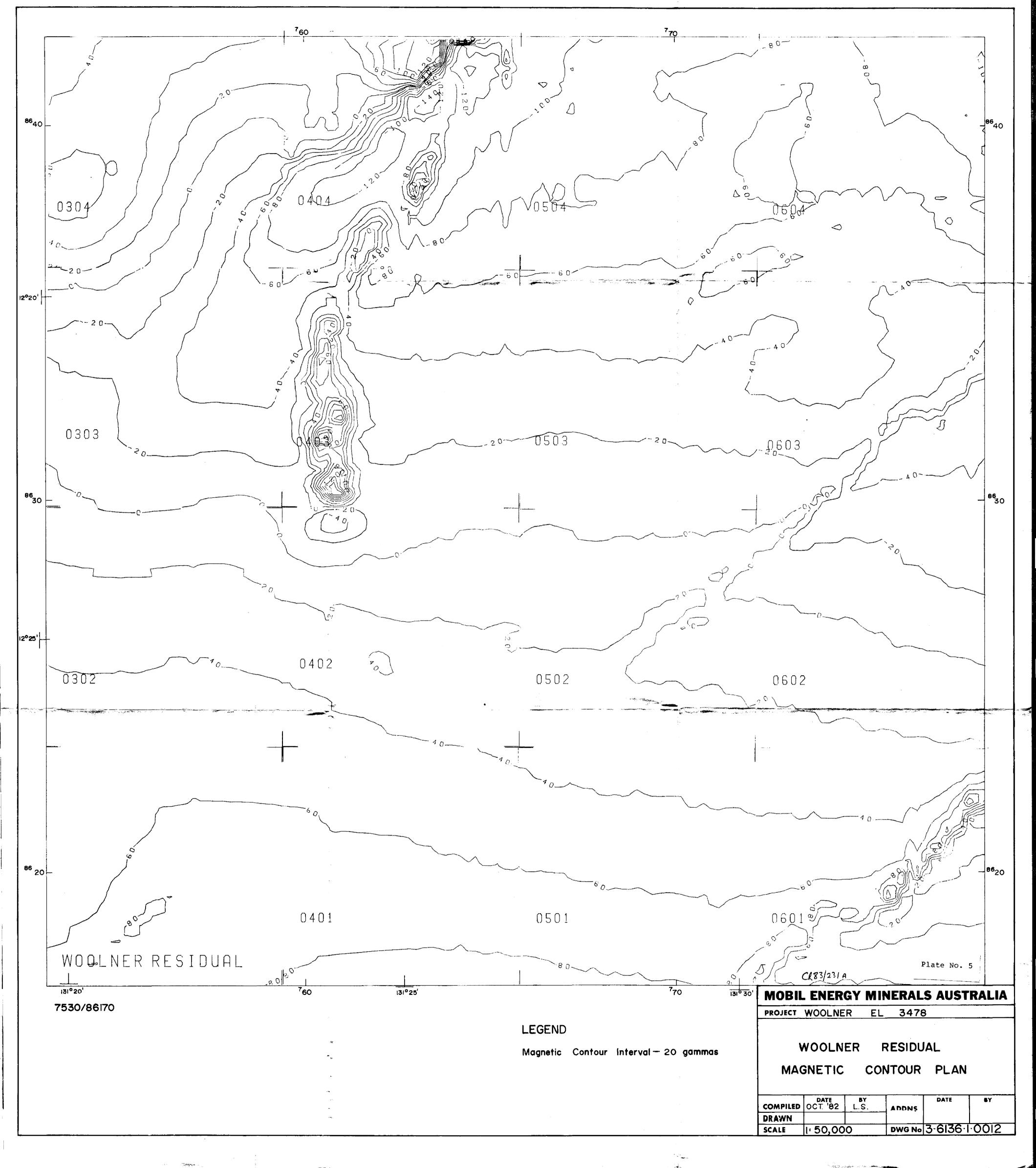
Additional exploration should consider testing the relationship of the proposed Crater Formation at the southern tip of Woolner West Dome to the basement by detailed magnetic and gravity traversing. It would be of benefit to carry out complex resistivity and CSMAT methods to obtain a clearer definition of the Fish Creek Schists and discern how they are related to the basement complex and how they compare to the Manton Group in the Rum Jungle area.

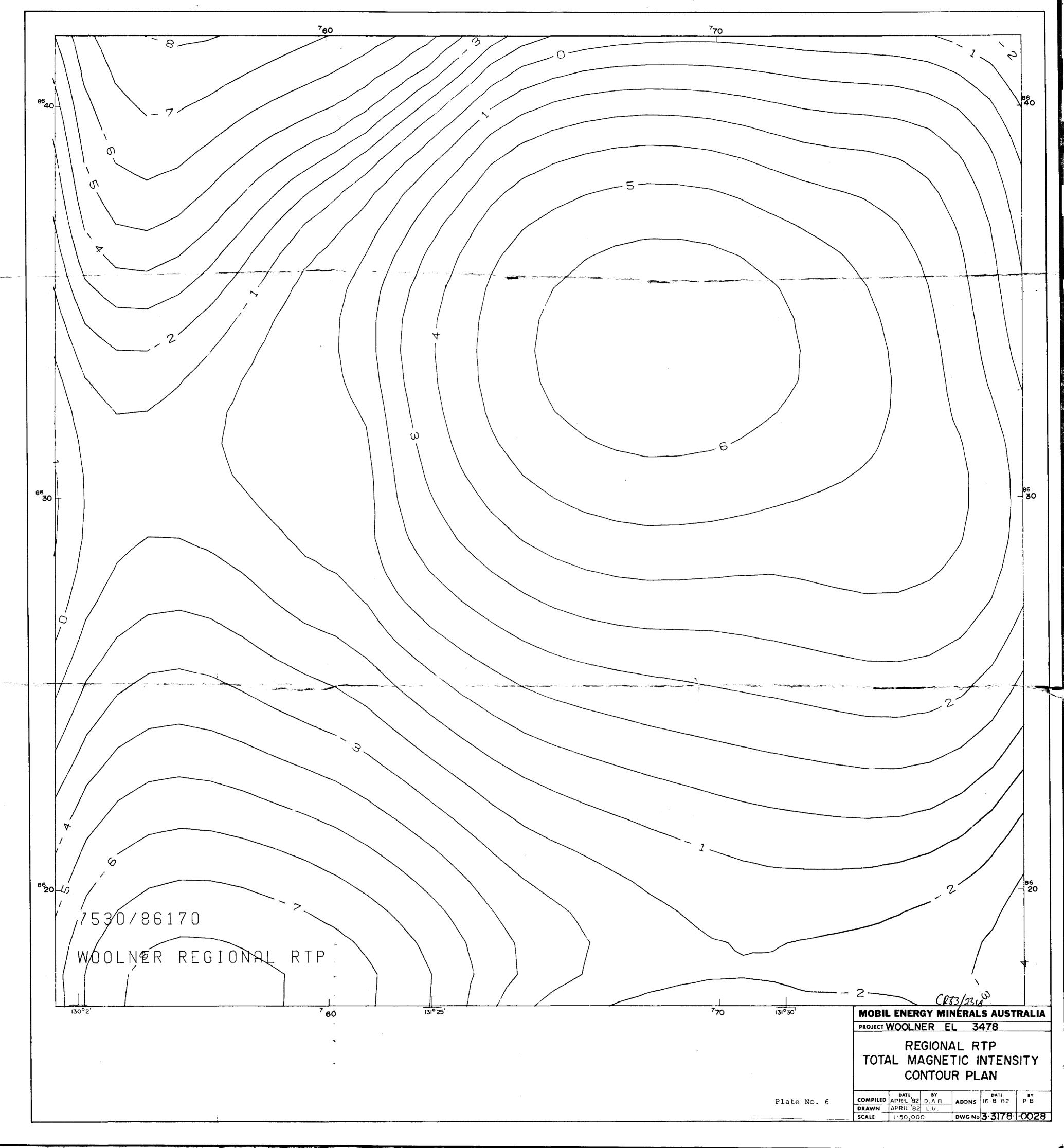


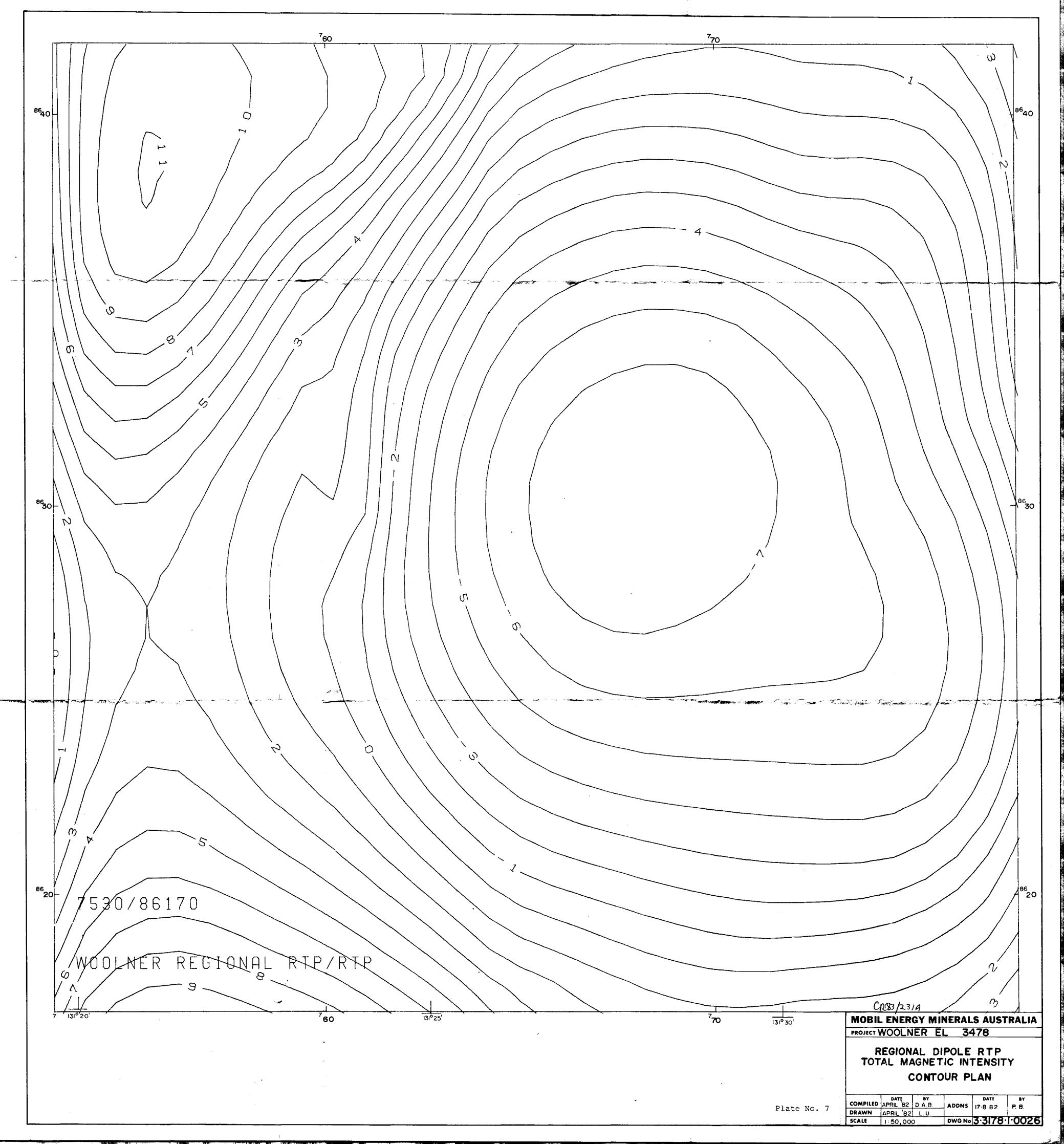


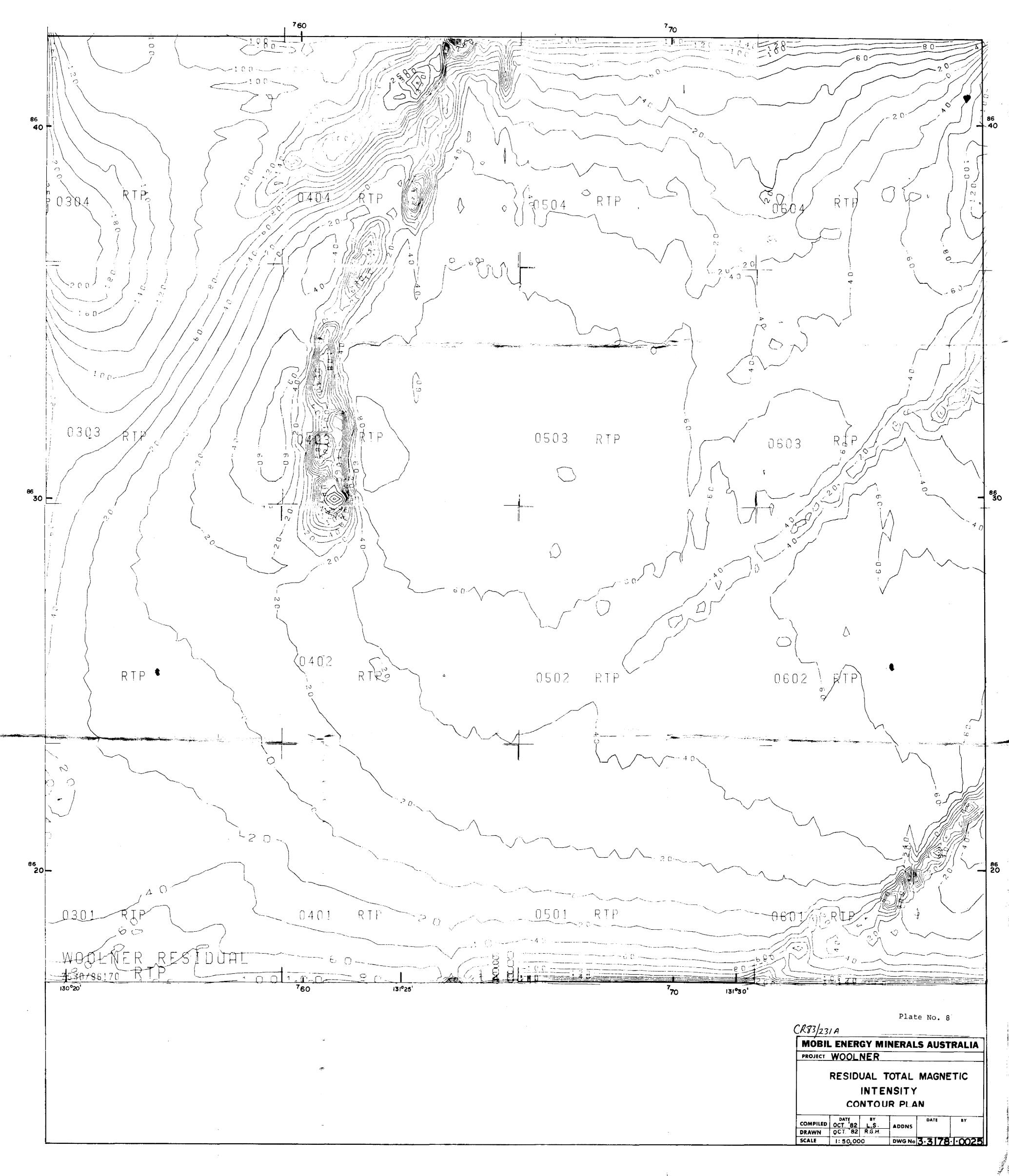


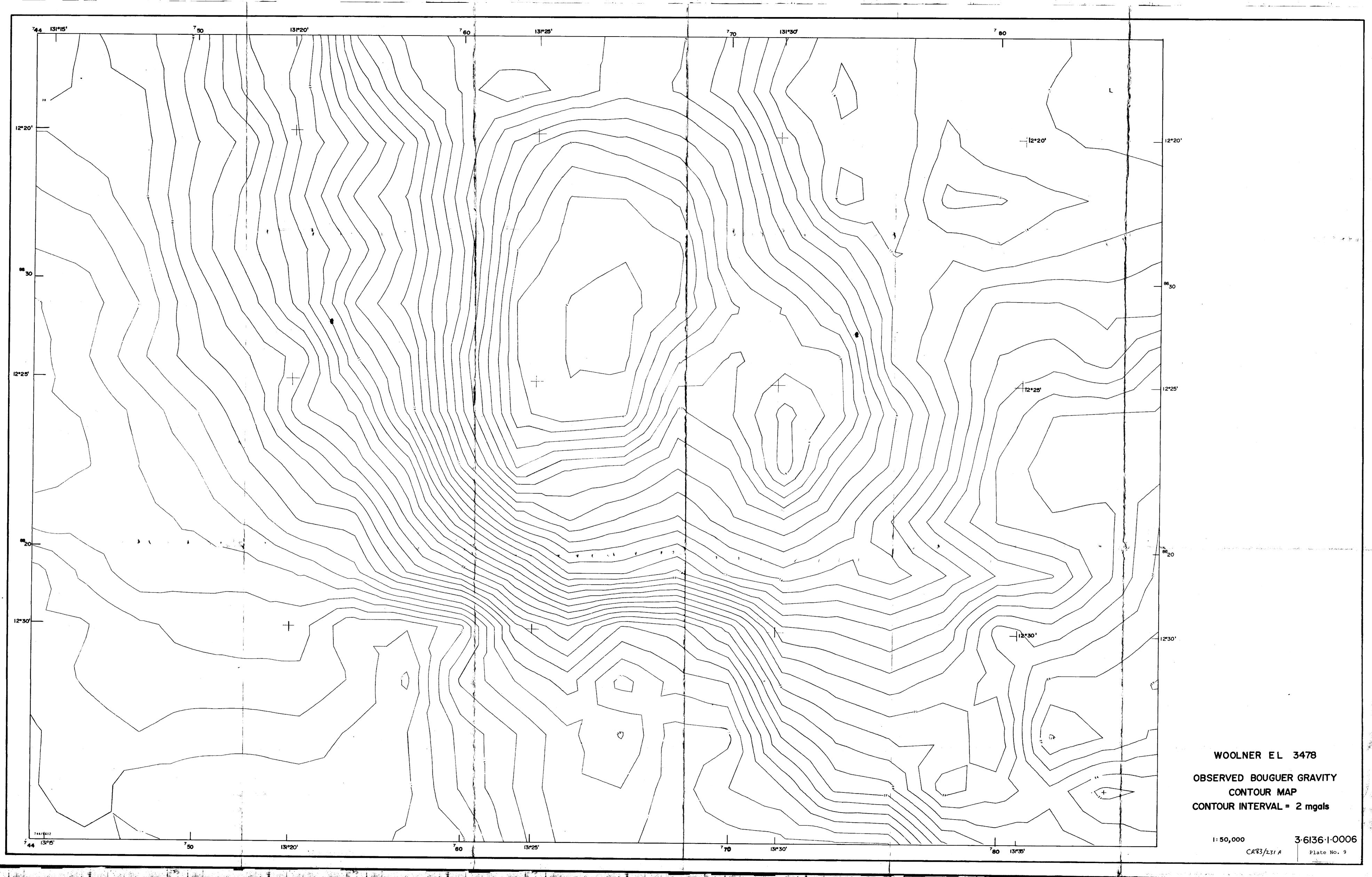


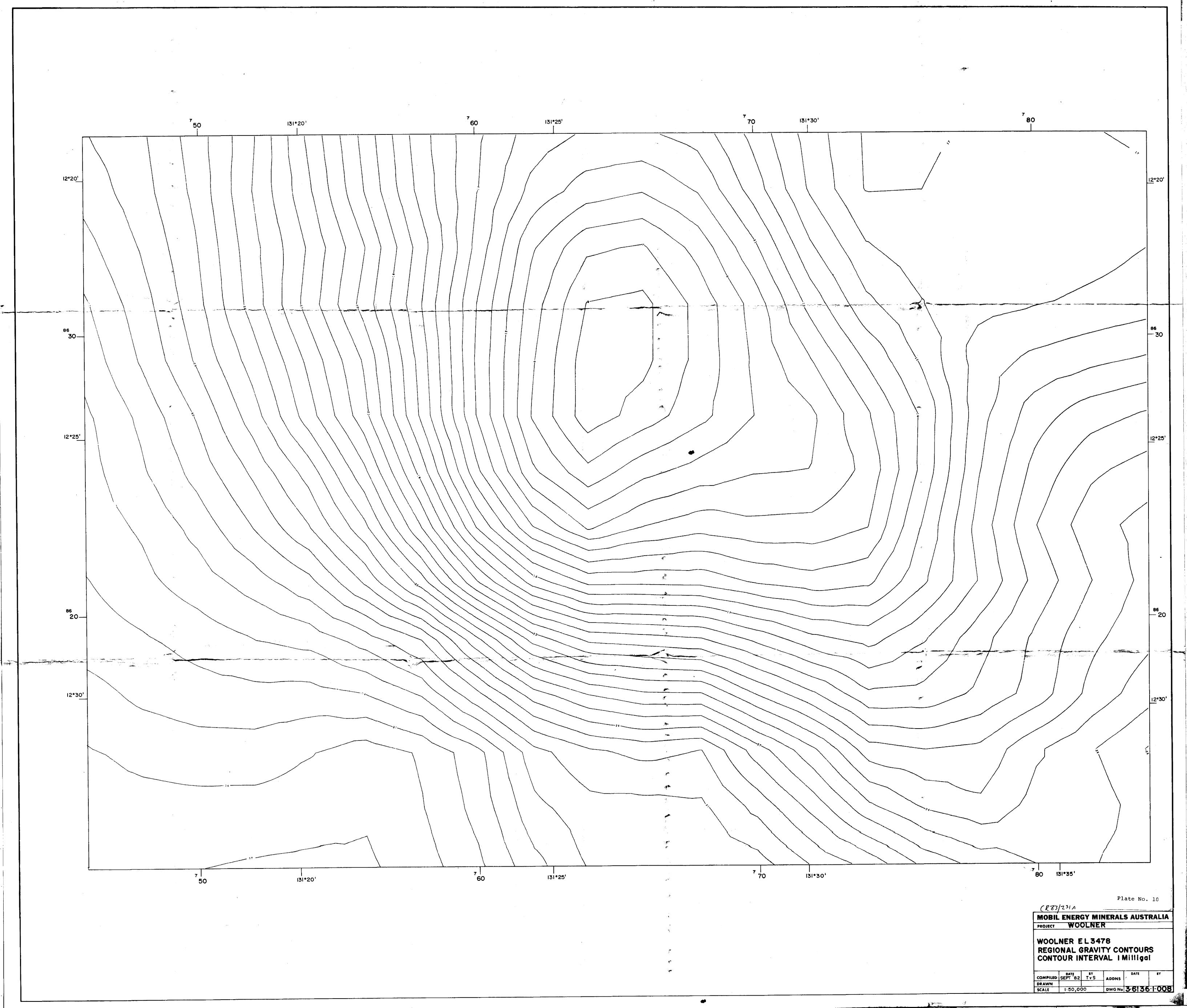


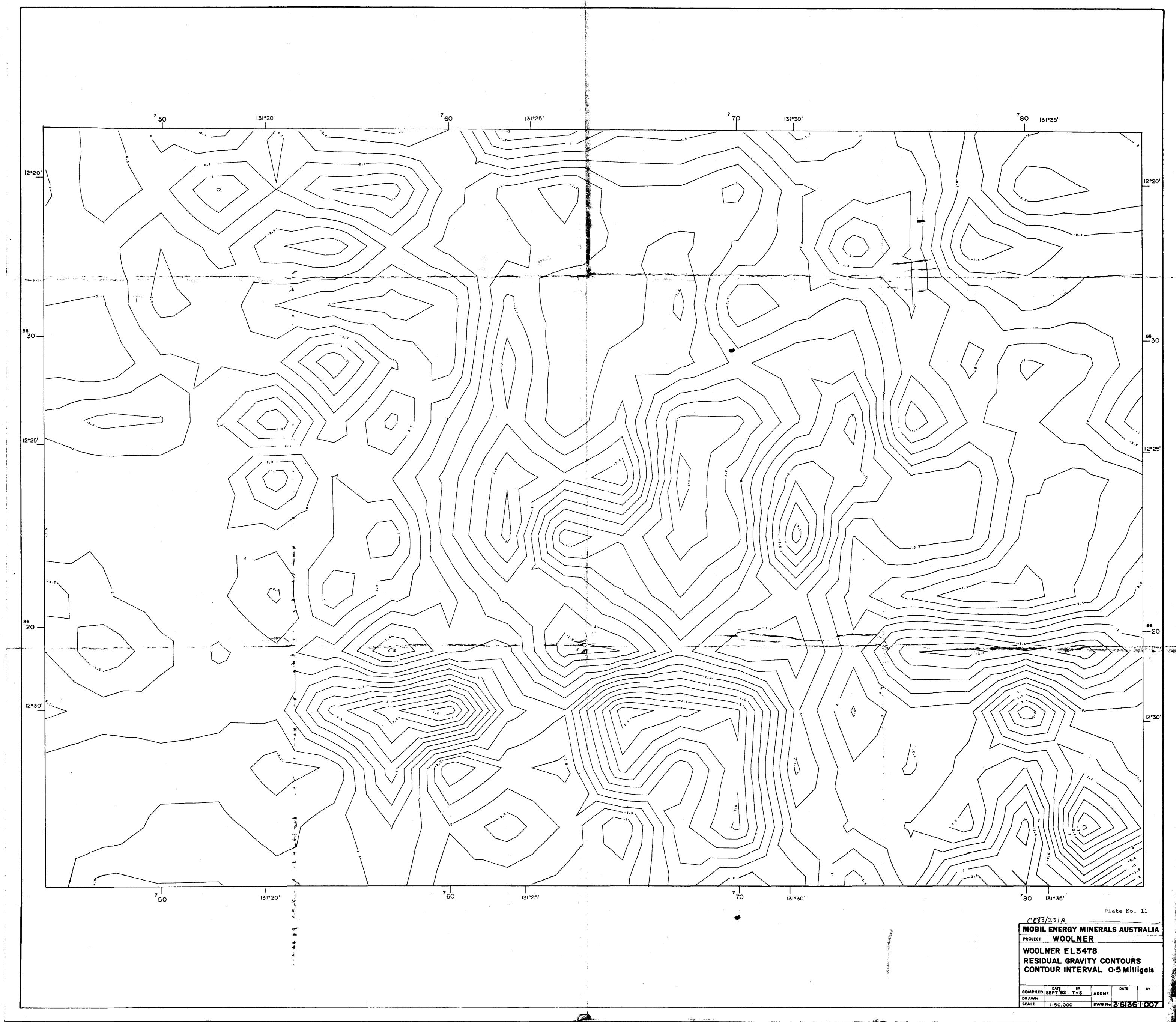


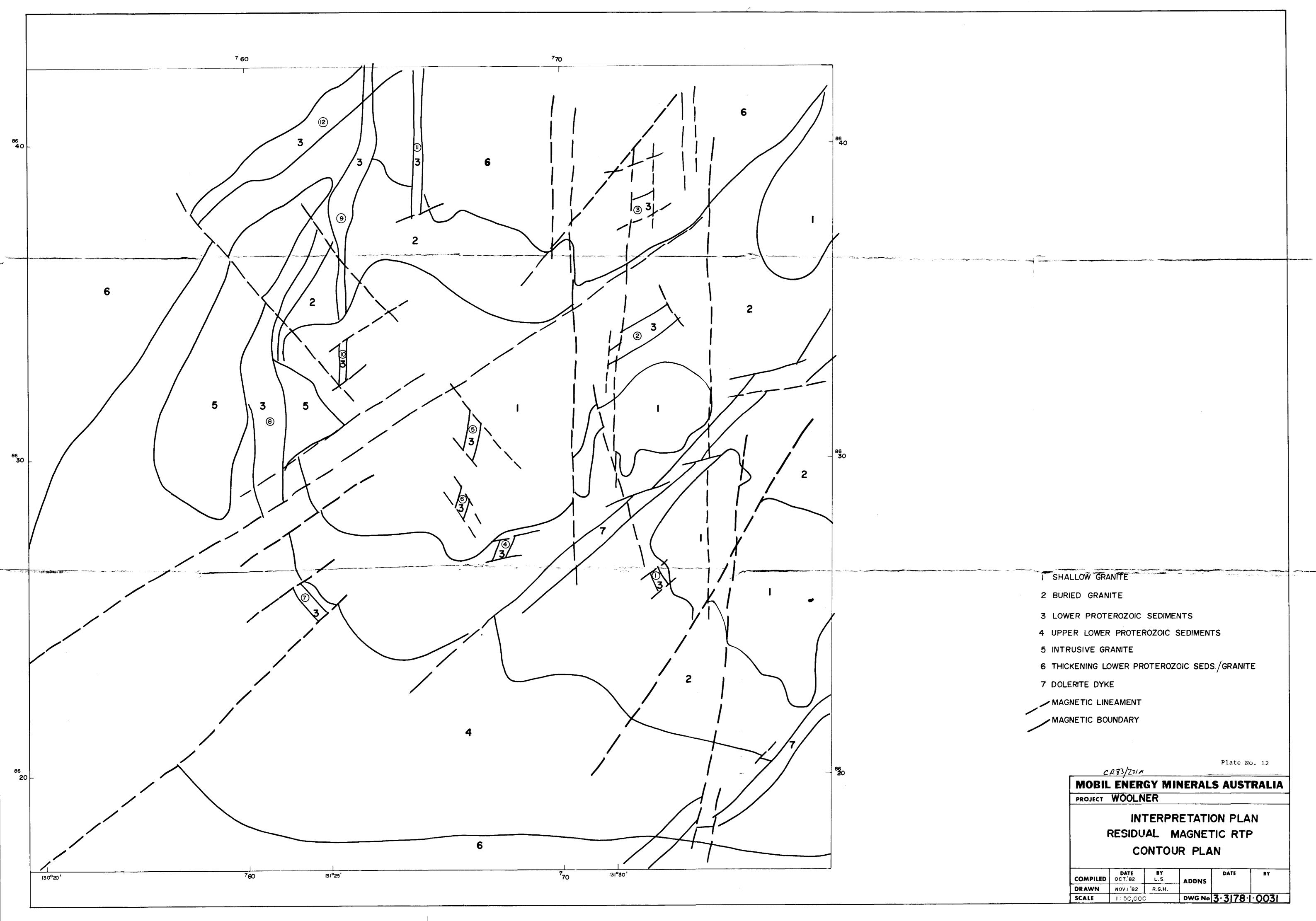


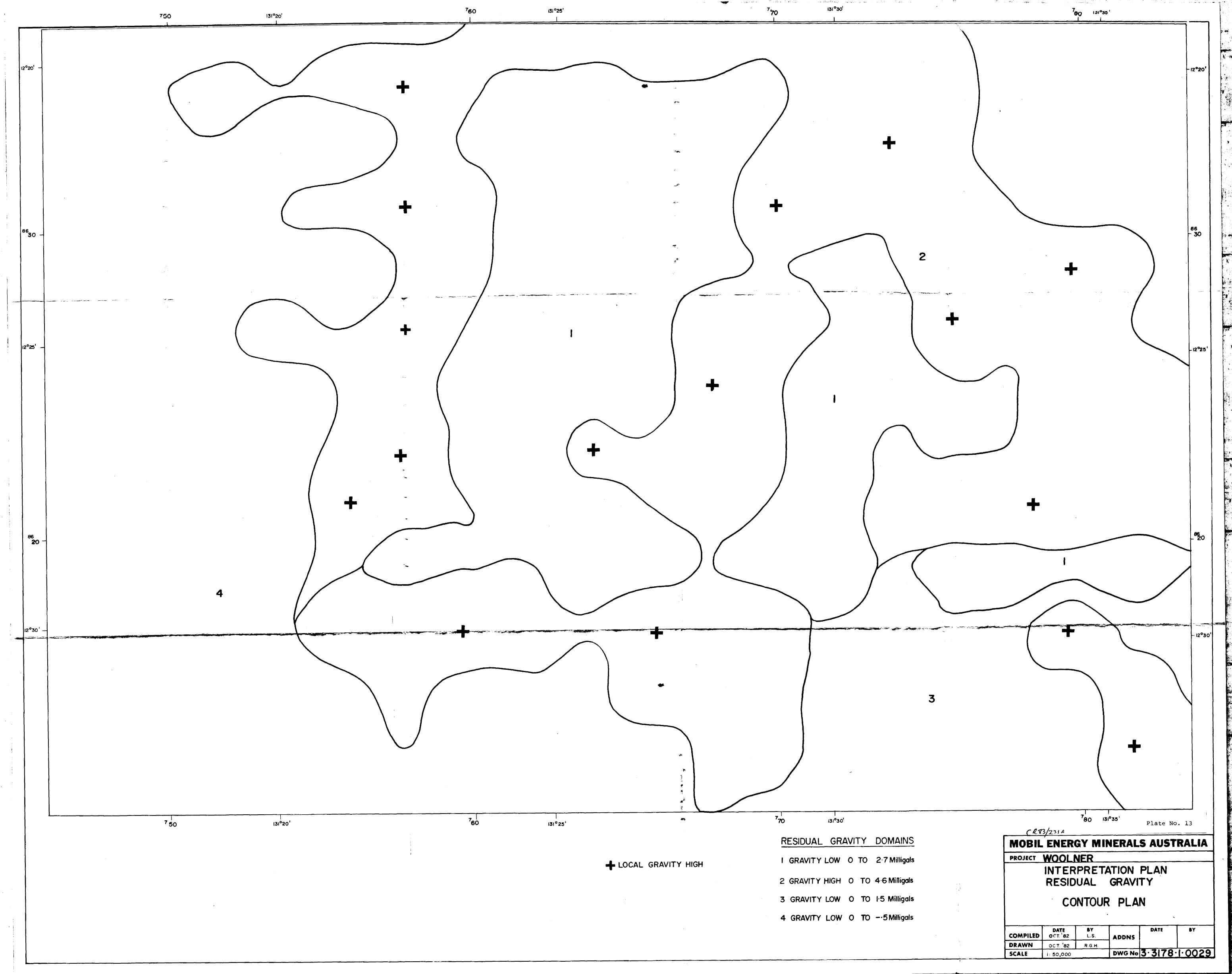












APPENDIX 1

GROUND GEOPHYSICAL SURVEYS AT WOOLNER (EL 3478)

DURING 1982

BIBLIOGRAPHIC DATA

TITLE:

Ground Geophysical Sruveys at Woolner (EL 3478) During

1982

AUTHOR:

T. Von Strokirch

PAGES:

18

FIGURES:

1

APPENDICES: 1

PLATES:

34

KEYWORDS:

Geophysics, GEM-8, Graphitic conductors, EM, Seismic,

Gravity, Magnetics, Woolner.

SUMMARY

During the 1982 field season ground geophysical surveys were performed in the Woolner area (EL 3478) of the Northern Territory. EM, gravity, magnetic and seismic traverses were made to determine the relative effectivity of the above techniques in this type of terrain and to locate the edges of the granite domes where graphitic Lower Proterozoic sediments are expected.

The GEM-8 EM method proved capable of penetrating the conductive overburden and effectively located the Fish Creek Schists, a graphitic and pyritic unit on the western edge of the West Woolner Dome. Work elsewhere in the area located a small number of conductors which were also tested by ground magnetic surveys. Line B/C which lies between the two granite domes has relatively conductive zones near both granite edges. This line was tested with gravity and seismic techniques. The results warrant further testing by drilling on this line and as well as along other portions of the edges of the granites.

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INTRODUCTION

An orientation GEM-8 survey was conducted in August 1982. It was designed to pick the granite/sediment contacts at the edges of the granite domes using the GEM-8 electromagnetic system and thus to observe the effectivity of the technique in this environment. Following this survey a number of further traverses were planned to further outline the contact and any graphitic units that might be present. Gravity and seismic traverses were also completed along Line B/C in order to test these techniques as well and to gain a better understanding of the area. These surveys were also preceded by drilling along Line B/C on a number of EM and magnetic targets.

In November 1982, 13 lines were pegged across areas of interest and the EM program was completed along these lines. Ground Magnetics were begun but not all the lines were finished due to the beginning of the wet season.

PRIOR GEOPHYSICAL SURVEYS

In 1978-79 CRAE carried out a limited ground magnetic survey however, it was not until the ground was farmed out to a Geopeko-E.Z. joint venture that a basic program of geophysics was undertaken.

Airborne magnetics were flown over the area at a height of 120 metres with a flight line spacing of 300 metres. Regional gravity was completed at 2 kilometre centres with localised infilling to 1 kilometre. A further 150 kilometres of gravity profiling was completed along pegged lines at 100 metre spacing with lines 200 m apart. Ground magnetic surveys were completed over 300 kilometres of pegged grid with readings taken every 25 metres.

Sirotem was attempted across the Fish Creek Schists in the north of the E.L. however after two days the survey was terminated as the highly conductive surface materials tidal movement made it impossible to obtain a clear bedrock response. The E.M. work was not formally written up.

METHOD OF SURVEY

Topographic Survey Positioning the lines proved to be fairly difficult in some cases since the featureless nature of the mud flats made air photo location difficult. The regular summer inundation of the greater portion of the area also removed the majority of the pegs from the Geopeko gridding program. Line B was located using obvious road junctions, fence lines and Lake Finnis. It should be accurately positioned and tied into the Geopeko grid. Lines A and E were located by travelling across country using compass and topofil. No pegs or drill holes were found on these traverses even though considerable drilling was completed along them in prior surveys. The cross correlation of the boundaries located with the EM on these lines should be recognised as being approximate and the lines if necessary could be surveyed at a later date.

Electromagnetic Survey Horizontal, vertical and large loop configurations were used in the test to find which would be the most suitable for the area. Horizontal loop surveys were completed with the transmitter 200 m west of the receiver. Vertical loop surveys were used in two configurations with the transmitter at 200 m and 300 m away to the east from the receiver.

The later work entailed a further 13 lines (Table 1). As the lines were widely separated a baseline was not established. The lines were positioned individually from air photos with the northern or westernmost peg being the one located. From this peg the lines were extended using compass and topofil. Most lines were located on black soil plains where pegging was easy. Lines around the edge of the floodplain passed through thick scrub and required grading. Line 64.2 N is located over 2 km out from the edge of the floodplain and could not be accurately located on air photographs and may therefore contain some positioning error. It was located using compass and topofil from the western end of line 24.0 north.

Lines F, 73.5 E and 28.5 N lie on a section of floodplain (a lake in the wet season) which lies in the east of the area and has rather difficult access. Any follow up drilling would require construction of a bulldozed road around the south of Lake Finnis. The eastern lines between northings 24 and 31 run through open forests and light scrub growing on lateritic soil. Some laterite is visible on the edges of the floodplain. The remainder of the lines are on the black soil plains which have a high conductivity.

15.5 kilometres was tested with the GEM-8 initially and a further 19.1 kilometres of survey was completed in the follow up program.

The horizontal loop In Line method was chosen as the technique which would give rapid progress and reasonable penetration of the conductive material overlying the Proterozoic. A 200 metre receiver transmitter separation gives a reasonable penetration whilst allowing the frequencies up to frequency six (2642 Hz) to be received. A spacing of 100 metres between stations was used as it was felt that this should be sufficient to pick conductivity changes in the basement.

Gravity Survey

Detailed gravity readings at 50 metre spacing were undertaken along traverses B & C.

Seismic Survey

A trial seismic survey was undertaken along traverses B & C with geophone spacing 10 metres and with a spread of 110 metres. The offset shot distance was 3 to 120 metres.

Magnetic Survey

Ground magnetics was completed at 10 metre spacing on line B-C to give maximum accuracy when correlating with airborne work. On other lines a 25 metre spacing was used. The survey was not completed due to the onset of the wet season.

INSTRUMENTATION

GEM-8

The GEM-8 is an eight frequency ground electromagnetic system. The frequencies range from 41Hz to 5248 Hz which gives it a useful capability as an electromagnetic sounding tool.

It has three major physical components:

- 1. The vehicle based transmitter consisting of the transmitter console, generator, transformer and tuning box.
- 2. The transmitting loop which is attached to the transmitter console. It consists of a ten metre multicored cable.
- 3. The receiver which is distant from and not connected to the transmitter. An HP 41CV calculator is used to derive Tilt Angle and Ellipticity information during receiver operation.

Seismic

The 12 channel Oyo "McSeis 1500" system was used to record data both on paper trace and diskette. The energy source was a MkIII Betsy Gun MCM-3, 40 Hz. Marshphones were spaced at 10 m intervals. Shotpoints were at 20 m intervals resulting in a six fold C.D.P. shock.

Gravity

Lacoste and Romberg gravity meter (G-473) was read at 50 m intervals. The line was levelled with a GDD electronic level. The gravity values are tied to A.G.G. by a reading on the Woolner airstrip (A.G.G. station No. A2023). The levelling was not tied to A.H.D. due to distance from bench marks and an elevation of 5 m was assumed for the base level. Accuracy of the gravity meter is ± 0.01 mgal and accuracy of the GDD is about ± 0.1 m (i.e. ± 0.03 mgal). This results in a relative error of ± 0.04 mgal though the absolute error may be as high as ± 4 mgal.

Magnetics

A geometrics Proton Precession magnetometer was used to take readings at 25 metre intervals except on Line B/C where readings were taken at 10 metre intervals.

REDUCTION OF GEOPHYSICAL DATA

Gravity

Bouguer anomalies were calculated using a density of $2.67~{\rm Tm}^{-3}$. A figure of $2.2~{\rm Tm}^{-3}$ might be more realistic as the younger sediments covering the dolomite are of low density but since the area has a height range of only 3-5 metres, no relative error will result.

Magnetics

The raw data formed the main basis for interpretation. The Adept computer program which was used to process the airborne magnetic information was also used on line B/C.

E.M.

Tilt Angles and Ellipticities were calculated from the raw data and these were plotted as profiles and pseudosections. Conductivities were derived from the ellipticities and tilt angles in some cases and these were used to determine approximate depths of penetration.

Table 1.

GEM-8 Survey at Woolner in November 1982 at 100 metre spacings.

Line	F	Runs 1.2 km North-East from FO	1.2km
Line	36	69.1E to 70.6E	1.6 km
Line	28.5N	74.7E to 75.2E	0.7 km
Line	73.5E	30.5N to 31.9N	1.5 km
Line	27.0N	64.6E to 67.4E	2.9 km
Line	26.0N	66.5E to 67.3E	1.9 km
Line	25.0N	66.3E to 67.2E	0.9 km
Line	24.0N	66.2E to 67.2E	1.1 km
Line	64.2E	23.0N to 24.1N	1.2 km
Line	31.0N	68.3E to 69.9N	0.8 km
Line	29.0N	66.4E to 68.4E	2.1 km
Line	26.0N	69.4E to 71.4E	2.1 km
Line	24.0N	69.3E to 70.3D	1.1 km
			-
Line	A		3.5 km
Line	B/C	North of Lake	4.5 km
Line	D	Along Road	5.0 km
Line	E	Near Coast	2.5 km
			34.6 km total

RESULTS

Original Surveys.

Line A and Line E were placed across the Fish Creek Schists in which graphitic schist bands have been clearly located by drilling and ground magnetics. Horizontal loop surveys were performed on both lines, and vertical and large loop modes were also used on line A in an attempt to obtain finer resolution of the anomaly. The results are as follow:

Line A

3.5 kilometres of Line were traversed using the horizontal loop system. Low signal strength combined with a very high noise level to make it difficult to obtain stable repeatable readings, especially in the low and high frequencies. Short period and long period noise meant that a long time was required to obtain accurate readings at each station. These problems were assumed to be caused by near surface change inconductivity of the black soils in the area (encrusted salt could frequently be seen in the area as well as the occasional seaweed) and the tidal effects along Fish Creek itself. Two full days were taken to complete the original traverse, with only thirty-four stations being surveyed.

A conductive zone was found on the line but no pronounced narrow conductors were visible (Plate 1). From A16.5 to A0 the bedrock conductivity increases relative to the surface conductivity. This effect is most pronounced in the tilt angle data at frequency HIHZ. This indicates that the bedrock is at the limit of penetration of the system and that the conductive response is due to poor (disseminated) conductor.

The eastern boundary is evident at Al6.5 coincident with the boundary as found by drilling. It is represented by a rise in conductivity over the schists from the granites to the east. Drill hole data located the western edge of the Fish Creek Schists at All.5. It is expected that the crystalline dolomites would be more resistive than the graphitic and pyritic schists however this does not seem to be the case. A increase in conductivity is found in the west with the contact at approximately Al2.0. The depth to the basement is estimate at 90 metres which is slightly in excess of the drill hole depth.

Vertical loop surveys were completed across the eastern contact to define more closely the boundary between the West Woolner Granite and the Fish Creek Schists (Plates 2,3). 200 metre and 300 metre spacings between receiver and transmitter were used. The noise problem was greatly reduced with this EM mode and both traverses (20 readings total) were completed in a morning. The coupling with these horizontal layers was too great and we did not get any definite response. The 200 metre spacing test was totally unresponsive, to bedrock and the 300 metre spacing data a small variation was measured at HIHZ. The tilt angle and ellipticity agree poorly indicating 3D effects however this is not the case with the Fish Creek Schists. It would appear that there is bedrock response allbeit poorly defined. The vertical loop method is unsuccessful on this line.

A large loop system (similar to TURAM) was tested to obtain deeper penetration (Plate 4). Noise effects were extreme and great difficulty was encountered in laying out the loop as the rough country making driving very slow and the featureless nature of the terrain making accurate sighting and positioning of the loop difficult. The transmitter loops were laid out as large rectangles 1.6 kilometre by 600 metres. It was only possible to get five readings on the line

from each loop up to a maximum of seven hundred metres away. An increase in apparent conductivity was observed at the eastern end of the section of line tested in both cases. This would appear to be a parametric responce rather than any true bedrock response. The large loop system was not successful in penetrating the surface material.

Line E

This line was close to the coast and its entire length was covered with seaweed. A high near surface variable conductivity produced a high noise level however deep responses were detected on the tilt angle (Plate 5). The ellipticity profile showed little variation even at 41 HZ. The tilt angle showed two contacts at E19 and E7 as well as a slight positive anomaly at E18. Drill hole information shows these contacts to be at E19 and E7, the boundaries of the Fish Creek Schists. The small conductive response is presumed to be due to the dolerite found at the eastern contact. The zone to the west of E19 which is overlain by an increasing thickness of sediments is due either to porous Coomalie dolomite or to an older dolomitic rock on the edge of the Fish Creek Schists.

Line B and C

These two lines join with peg B30 being equivalent to CO so the two lines are plotted as one profit along a line meandering from the West Woolner Dome to the East Woolner Dome. The traverse along this line was fairly rapidly done using the horizontal loop (average 30 readings per day). The results are quite interesting (Plate 6). Drill hole information indicates a thickening of the overburden in the middle of the line. This is seen on the conductivity plot (Fig. 2) for frequency 41 HZ as a regional increase across the middle of the profile reflecting the increased amount of conductive overburden. Two zones of substantially higher conductivity are located on the EM profile between stations B2 and B4 in zone of conductive material

which from drill hole information would appear to fall within the boundaries of the West Woolner Dome. Some pyrite was found in cross cutting veins in a nearby drill hole and it may be that this mineralisation is more extensive in this area than has been realized. The other possibility is that we have had rock movement in the area so that a portion of the granite has been faulted out of place and that the conductor is due to graphitic material faulted into the granite. The eastern conductive feature is more pronounced and the resistivity of the East Woolner Dome appears higher though this is partially due to decrease in the thickness of the cover. This conductor is interpreted to be due to graphitic material along the boundary with the granite.

Line D

This line is located southeast from the homestead water bore. It has no EM anomalies pronounced in its entire length so it is presumed to not have sufficient northern extent to reach the granite contact. As granite was encountered in the homestead bore the contact must be located to the south between there and the first reading which was taken two hundred metres away to be well clear of the buildings. No rise in conductivity was noted at the end point of the line so if any conductive unit exists at the granite boundary at this point it must be very narrow.

Line F

Conductive overburden is evident increasing in thickness to the south. At station F8 there is a high in the tilt angle response which is associated with an early turnover in the ellipticity. This could be indicative of a sedimentary channel in the Cretaceous sediments filled with more recent sediments and saline fluids.

Line 73.5E

The EM data approaches a three layer model when compared with the type curves although a four layer curve might be more representative. The upper layers are extremely conductive. There is a slight increase in conductivity at depth around 31.8E which indicates basement response with little variation, presumably a granite.

Line 26.0N (from 69.4E - 71.4E)

EM coupling is pronounced. The small changes in conductivity near the surface are due to variation in the thickness of the upper layer, namely the recent cover. The narrow conductor at 69.7E is also a shallow response. A feature at 71.0E is more persistent at depth and is interpreted to be related to bedrock, perhaps to sediments of Proterozoic age which are other than dolomite. The response is somewhat narrower than that on Line C which appears to be at an equivalent position in the sequence.

Line 26.0N (from 66.5E - 67.3E)

This data does not match our standard curves well. A stronger conductive layer is to be found at a depth of about 20 m. It wedges out to the west at 66.8E. The surface conductivity is particularly high at the edge of a flood plain.

Line 25.0N

Ground magnetics were completed on this line and showed up no anomalous features. The EM response indicates that the conductive overburden thickens sharply to the west of 66.8E.

Line 24.0N (from 69.3E to 70.3E)

At this line there is relatively shallow bedrock with conductive overburden generally less than 40 m in thickness. A narrow dyke with conductive features has been located 70.1E. The feature is strongly

pronounced on both tilt angle and ellipticity. West of 70.0E the overburden thickens rapidly. A ground magnetics traverse indicates that the line can be divided into two zones east and west of 70.03. To the west the magnetic response is featureless. To the east it is sharply disturbed. This appears to be the contact between the dolomite in the west and laterite on top of granite or Proterozoic sediments in contact with a granite. In this area further ground magnetics would be useful to better define the extent of these two zones.

Line 24.0N (from 66.2E to 67.2E)

A ground magnetic traverse indicates a 500 gamma feature at 67.13E. This anomaly is interpreted as being caused by a dolerite dyke which dips steeply to the west. The magnetics is otherwise featureless indicating that the bedrock is covered by a thick layer of dolomite with a depth of at least 65 m. The EM also indicates that the thickness of overburden thickens to the west. A vertical feature occurs at 66.9E. It does not extend to depth and must be located in the Cretaceous sediments and not easily explained. A steeply dipping conductor is located at 66.6E.

Line 27.0N

The conductive overburden increase in thickness to the west from 66.6E. There is a west dipping dislocation in the overburden at 65.5E. A change in bedrock occurs at 66.87E with as contact dipping to the west at a small angle. The ground magnetics has no anomalies apart from small (10 - 15 gamma) highs which are at pegged stations and are probably due to readings being taken too close to the steel fence droppers which were used for gridding. A poor conductive response is present in frequency 41HZ at 64.9E. This is close to the granite contact as outlined by Peko. It is probably indicate a fault within the granite as the contact is at 66.9E.

<u>Line 28.5N</u>

This line is only 600 m long as difficulty was experienced in clearing the bush at either end of the line. A fault is located at 74.8E with a small change of conductivity across it. This could be the contact between two different types of granite. The overburden thickness is approximately 40 m. A ground magnetic profile over this line would indicate if we have a narrow dolerite present but the wet season set in before it was possible to survey this part of the area.

Line 29.0N

Magnetic contacts occur at 66.45E and 67.40E. Narrow dykes are indicated at 66.63E and 66.76E. The EM data indicates a basement contact dipping at a slight angle to the west from 66.8E. At approximately 20 m depth on the eastern portion of the line a highly conductive band shallows over possible granite.

Line 31.0N

Traverse is 800 m limited in length by thick bush. The surface material has a lower conductivity to that found on the flood plains. No conductive basement features are evident. A dyke is visible on the magnetic profit at 68.47E which is probably produced by a narrow dolerite dyke. A contact is located on the magnetic profile at 67.57. Proterozoic sediments occur on the west with dolomite occuring to the east.

Line 36.0N

No strong bedrock responses are evident though a change in frequency 41 HZ at 70.1E is probably caused by a contact between granite and sediments. At 70.2 to 70.3E a surface conductor indicates a possible sedimentary channel. A second channel not so pronounced could occur at 69.2E.

Line 64.2E

Very high conductivity in the surface layer is located on the line below the surface layer a layer of not quite so high conductivity but still high conductivity has been mapped. The magnetic along the line is quite featureless indicating dolomite bedrock. Noticeable variation produced by variation in thickness of the surface layer is shown by the tilt angle response.

Line B/C Seismic and Gravity

These surveys were conducted as a test for seismic and gravity survey methods in the highly weathered areas as well as to define more clearly the anomalous features detected with the EM along line B/C. Gravity readings at 50 m intervals defined the edges of a broad gravity high (1 mgal on amplitude) which lies between the two major granite domes. This represents a broad slab of flat lying Coomalie Dolomite which overlies granite and probable earlier Proterozoic sediments. Two types of granitic basement are evident on the gravity and also by initial perusal of the seismic data. One granite is a bouyant intrusive granite with low density of 2.61 gms/cc and the other is a somewhat denser granite gneiss of density 2.67 gms/cc. The Fish Creek Schists are widely distributed while the other metasediments occur only in a narrow band distinguished by slightly higher density and velocity and by lower magnetic response.

Except for the area around B9, the dolomite is up to 70 m thick therefore it is quite costly to drill, but this is required to ensure proper economic understanding of this area.

CONCLUSIONS AND RECOMMENDATIONS

The schists which have been proposed to exist below the Coomalie Dolomite between the two granite domes are an exploration target which it will be necessary to test by drilling Drill holes are recommended at Line 27N 65.7E to test rock type change in basement. Line 29N 66.85E to test basement conductor near granite contact. Line C 6E to test basement conductor. Bedrock should be encountered at 160m at the outside.

Elsewhere the edges of the granites are usually associated with some conductive response (except where the amount of cover is too great for the GEM-8 to penetrate). This indicates that the granite has upfaulted conductive Proterozoic metasediments, some of which possibly contain graphite from lower sections of the Fish Creek Schist. The sediments on the western end (adjacent to the eastern margin of the western dome) of the GEM-8 traverse in the embayment between the two domes are more conductive than those on the eastern end of the traverse (western margin of the eastern dome) Therefore the sediments on the western end of the traverse are more prospective for further study if drilling determines that there are promising graphitic schist horizons below the Coomalie Dolomite in the embayment between the two granites.

APPENDIX 1

Frequencies of GEM-8 channels.

Frequency	0	41	HZ
Frequency	1	82	ΗZ
Frequency	2	164	HZ
Frequency	3	328	HZ
Frequency	4	656	HZ
Frequency	5	1312	HZ
Frequency	6	2624	HZ
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16.0

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34.8 36.3 39.2 40.8

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53.9 52.3 54.8 \$3,0 55.1 56.6 53.9

44.6 42.8

36,5 36.6 38.7 / 33.0 32.8 37.4 37.7 38.3

2624Hz

FREQUENCY

LINE F

FREQUENCY RANGE LEGEND STATION Nº 41Hz 6 82Hz 164 Hz 328Hz 656Hz 1312Hz

MOBIL ENERGY MINERALS AUSTRALIA

WOOLNER PROJECT

> **ELLIPTICITY PSEUDOSECTIONS** LINE F (200 M SEPARATION)

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LINE A (200m SPACING)
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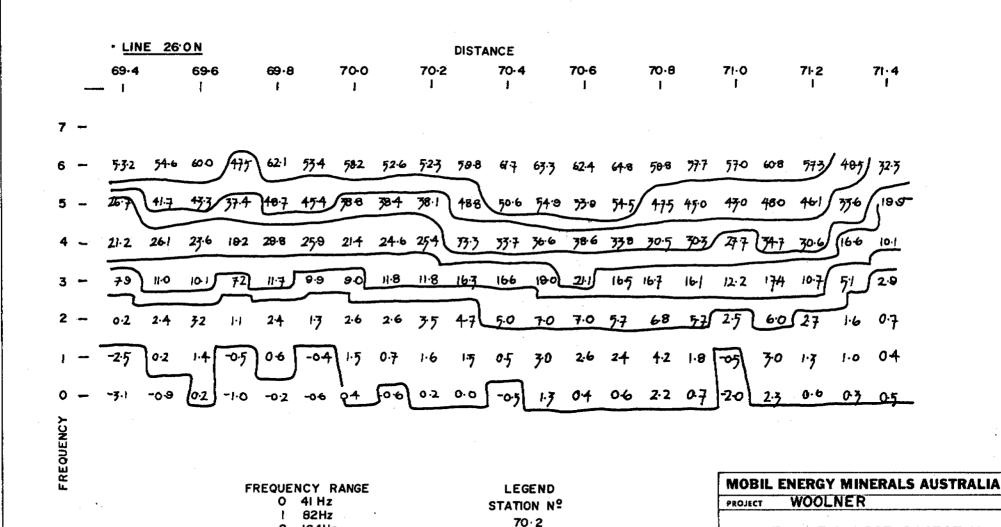
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FREQUENCY
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FREQUE	ENCY RANGE	LEGEND	
0	41Hz	STATION N	!
1	82Hz	1	
2	16 4 Hz	•	
3	328Hz		
4	656Hz		
5	1312Hz		
6	2624Hz	Tx IN	I WES

WORIT	. ENER	GY MI	NERAL	S AUST	KALIA
PROJECT	WOOL	NER			
TIL	T ANG	LE P	SEUD	SECT	ION
	LINE	24.0	N		
	WEST	ERN S	SECTIO	N	
******	DATE	67		DATE	BY

DWG No 3

TINE 25.0 N

66.4

66.6

66.8

67.0

67.2

6 — 12.5 3.6

-8.8

20.3

27.7

27.8

25.1

26.3

24.9

5 — -8.7

-8.5

-467

-25.2

42.2

29.8

35.4

35.1

17.1

20.6

3 — 51.0

48.3

44.2

50.5

50.3

41.2

38.5

44.7

45.0

62.5

2 — 49.9

151.3

49.3

62.6

35.5

37.6

23.6

23.6

23.6

23.6

23.6

23.6

23.6

23.7

24.4

25.6

25.7

26.4

26.4

26.2

27.8

26.4

27.8

28.5

29.8

35.4

35.4

35.1

17.1

20.6

3 — 51.0

48.3

49.3

62.6

35.5

37.6

26.3

30.6

29.1

33.3

1 — 51.6

53.7

51.9

51.7

23.6

23.6

23.6

17.2

19.3

22.4

23.2

0 — 40.9

40.1

33.9

28.5

12.1

11.6

10.4

12.6

10.0

5.8

Ş	FREQUENCY RANGE	LEGEND
鱼	O 4IHz	
FREQUENCY	I 82Hz	STATION Nº
E.	2 164Hz	66 ⋅8
	3 328Hz	1
•	4 656Hz	
	5 (3)2Hz	
	6 2624Hz	, Tx IN WEST

PROJECT WOOLNER TILT ANGLE PSEUDOSECTION

LINE 25.0N

i						
	COMPILED	DEC. 82	L.J.s.	ADDNS	DATE	SY
	DRAWN	7 FEB. 83	R.9, H.	1		
	SCALE	1:10,0	00	DWG No	3.3178	· 1-0101

Plate

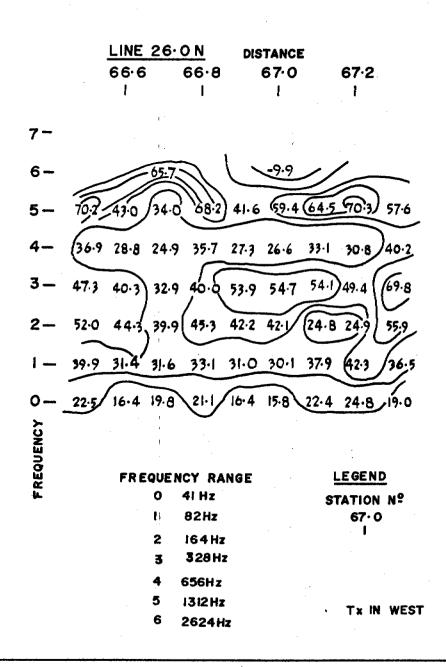
No.

12

MOBIL ENERGY MINERALS AUSTRALIA

COMPILED

DRAWN



DATE

DWG No 3:3178-1-0103

ADDNS

27

DATE

7FEB. 63 R.G.H.

1:10,000

COMPILED DEC. 82

DRAWN

SCALE

LINE 26 ON DISTANCE 69.4 69.6 70.0 70.2 70.4 71-0 70.6 71.2 71.4 7_ 37.0 31.6 75.1 26.0 37.0 39.6 38.5 36.0 37.9 25.2 22.2 20.9 29.2 24.9 29.6/33.0 29.0/30.5 41.4 42.8 41.4 46.9 46.7 44.0 46.9 48.5 40.5 39.9 41.4 39.2 39.3 35.7 40.8 41.2 40.1 42.7 39.9 39.9 47.3 45.1 46.3 43.7 49.9 48.9 45.8 41.8 40.6 47.4 46.0 47.6 45.8 46.8 43.7 42.6 44.9 44.5 47.5 42.2 130.7 35.5 37.7 35.3 33.2 40.9/36.7 32.4 34.7 34.4 41.9 42.4 44.1 45.8 41.0 /37.6 38.4 38.9 22.5 24.0 21.1 20.1 24.9 23.5 20.5 23.5 23.4 28.8 30.2 31.1 33.4 28.3 26.5 26.8 25.7 /14.5 10.2 11.1 10.5 13.2 12.2 11.0 12.3 12.8 16.1 16.2 16.6 18.2 | 4.8 | 14.7 | 15.1 | 13.6 6.4 5.7. 6.6 6.2 6.4 6.5 8.4 9.2 9.7 7.3 10.5 8.3 9.2 7.5 FREQUENCY LEGEND FREQUENCY RANGE 41Hz STATION Nº 82Hz 69.8 164 Hz **MOBIL ENERGY MINERALS AUSTRALIA** 328 Hz PROJECT WOOLNER 656Hz TILT ANGLE PSEUDOSECTION 1312Hz LINE 26.0 N 2624Hz EASTERN SECTION Tx IN WEST

```
LINE 28.5 N
                             DISTANCE
      74.7
                74.9
                                      75.3
                                                75.5
                           75.1
7 -
      44 4/52 2 50 7 55 0 52 8 49 3
      36 8 21 · 2 B5 · 6 33 · 2 34 · 7 35 · 3
FREQUENCY
                                                         MOBIL ENERGY MINERALS AUSTRALIA
        FREQUENCY RANGE
                                      LEGEND
                                                         PROJECT WOOLNER
                 41Hz
                                     STATION Nº
                                                            TILT ANGLE PSEUDOSECTION
                 82Hz
                                       75-1
                                                            LINE 28.5N
                 164 Hz
                 328 Hz
                 656 Hz
                                                         COMPILED DEC. 82 L.J.S.
                                                                                           BY
                 1312 Hz
                 2624Hz
                                        Tx IN WEST
                                                         DRAWN
                                                               3 FEB.83 R.G.H.
                                                                            DWG No 3.3178-1.0105
                                                                1:10.000
```

	LINE 29.0	N			E	DISTANCE						
	66.4	66 ₋ 6	66·8	67.0	67.2		676	67-	в е	90	68,2	68:4
7 -						•	•	•		•	•	'
·											÷	
6 -	46.4	7.3 5.4	78-0//47-8	48.9 17.8	9.2	6.6 -43.1	2.8) 13.4	II-8 13-I	14-4	17.8	12·8 i	7.5 17.4
5 ~	31.4 36.6	7.8 10.7	24.8 49.5	48-1 20-8	10.9	12.8 21.8	18.3 (21.8	23.6 23.4	18.3	22.5	10-1 21-8	18-3 /22-9
4 -	36.0 37.6	19.8 28.7	32:2 39:3	37.6) (19.9	7.2	29.5/16.7	15.5 16.8	19 16 5	16.5	20.5	14.6 16.7 1	6.9 17.3
3-	40.9 45.0	50.6 51.0	54.3) 29.2	18.2 23.0	21.7	21-1 25-2	23-8 24-2	25.3 22.7	28.7	29.5	24.0 23.6	27-2 24-7
2 -	78.0 45.6	3 43.4 30.8	27.8 11.4	8.7 12.0	12-1	10.5 12.4	13.0 10.9	11-7 10-8	14.6	17-3	13-1 12-3 1	5.2 12.9
1 -	40.7 37.1	31-1 /14-5	15.8 5.4	3.8 4.0	6.7	5-1 6-4	6.4 6.0	6.8 6.3	6.7	8.0	6.4 6.6	7.0 5.8
0 –	30-3/24-0	16.2 8.1	9.3 3.2	3.4 2.7	-0.2	2.2 3.3	3.5/-4.3	-4.8 -5.8	3.0	5.1	2.5 0.0	5.7 1.4
≻							·					

FREQUENC

FREQUENCY RANGE

LEGEND

O 41Hz

82 Hz

2 164Hz

STATION Nº

66 . 8

Tx IN WEST

328 Hz

656Hz

1312Hz

2624Hz

No.

Plate

BY

DWG No 3-3178-1-0108

COMPILED DATE BY
DEC. 92 L.J.S.
DRAWN 3FEB. 83 R.G.H.
SCALE [:]0,000

DRAWN

ADDNS

	LINE 36 N		DIST	ANCE				
_	69·1 69·3	69.5	69.7	699	7 0·I	70.3	70·5	
•	1 1	1	!	1	l l	1	1	
7 -								
6 -	85.5/ 11.8/1/51.2/	36-5 31-4 38-3	54.4) 30.9	/ 8 1 = 4.5	1 10.4 0.0	2	33.5 10.3	
5 -	34.4 13.3 32.6 3	18-6 (18-9 17-6)	39.8 28.0	29.9 8.4) _{15.7} // _{33.1}	2 25-7 2	4.2) 35.4 23.3	
4 -	26.7 (7.9) 27.9)	11-3 (18-7 / 31-5	37.3 (27.4	29.9 24.2	31.1 51.8	27.8 49	0.2 55.3 51.2	
3 -	47.6—45. 45.4 4	6.4 50.0 48.0	47.8 41.5	43.4 43.4	49.2 56.3	44.8 56	5.7 55.9 55.6	
2 -	55.5 55.9 54.8 5	2.4 53.0 53.1	54.4 48.6	46.8 48.2	56.9 50.	4 41.3 55	·8 51·7 51·3	
l -	52.8 52.2 51.9 49	9.4 49.6 51.6	52.4 52.6	50.5) 49.1	54.5 45.3	42.0 47	4 42.8 43.8	
0 _	44.6 44.8 44.9 4	3.0 43.5 43.6	42.2 42.8	41.2				
					:			
FREQUENCY		FREQUENCY		LEGEN	<u>ID</u>			-
EQL		0 4IHz		STATION	Nº			
G.		1 82H 2 164i		69-7			MOBIL ENERGY	MINERALS AUSTRALIA
		3 328	-	ľ			PROJECT WOOLNE	
		4 656					TILT ANGL	E PSEUDOSECTION
		5 13121		•			LINE 36 N	
		6 262	4Hz					
		**			Tx IN WE	ST		

`. 73

```
LINE 64.2 E
                                  DISTANCE
      23.0
               23.2
                           23.4
                                    23.6
                                              23.8
                                                         240
7 -
5 -
                 14.9 130 / 20.5 23.3 , 15.9 51.1
                                                          10.0 /20.5
                                               -8-1 16-1
                                               49.7 54.7 55.9
               55.0 56.2
3 -
           61.2 58.4 55.5 59.4 59.0 58.7 (60.6 58.5 58.8 57.9
2 -
                               57.5 59.2 55.0 56.0 53.5 52.7
FREQUENCY
                   FREQUENCY RANGE
                                                    LEGEND
                          41Hz
                                                  STATION Nº
                           82Hz
                                                      23.4
                           164Hz
                           328Hz
                           656Hz
                           1312 Hz
                           2624Hz
                                                         Tx IN WEST
```

MOBIL	. ENER	GY M	NERAL	S AUST	RALIA
PROJECT	WOOL	NER			
TILT	ANG	E PS	SEUDO	SECTION	NC
	LINE	64-2	E		
COMPILED	DATE DEC. 82	L.J.S.	ADDNS	DATE	87
	3 FEB. 83	ROH	7		1
DRAWN	U 1 LU UU				

```
LINE 73.5 E
                                         DISTANCE
                     30.8
           30.6
                                          31.2
                               31.0
                                                     31.4
                                                               316
                                                                          318
7-
            6.2 12.6 14.7 19.2 15.9 25.5 34.6/14.2 18.5 15.4 17.5 16.9, 27.0 19.4
      3.8
           1.0
                      20.6 25.8 20.0
                                     15.4 19.B
                                                               14.2 13.4 13.0 4.4
     29.9
     48-1 46-9 48-0 48-7
                           46.4/53.8 59.2 56.2 54.2 54.8 52.5 56.1
                                                                     52.2 51.1 / 49.7
                54.5 50.6 49.0 49.8 49.5
                    47.0 44.7 43.1 40.6 43.0 41.8 /39.1
     36.5
                                                          43.2 46.8 40.9 39.2 42.5
     24-1 27-8
                32.9 37.5 34.5 32.2
FREQUENCY
                                              LEGEND
                    FREQUENCY RANGE
                            4lHz
                                                                     MOBIL ENERGY MINERALS AUSTRALIA
                                             STATION Nº
                             82Hz
                                                                            WOOLNER
                                                                     PROJECT
                                                31.2
                             164Hz
                                                                      TILT ANGLE PSEUDOSECTION
                             328Hz
                             656Hz
                                                                            LINE 73.5 E
                             1312 Hz
                             2624Hz
                                                    Tx IN WEST
                                                                     COMPILED DEC. 82
                                                                                               DATE
                                                                                                       BY
                                                                                        ADDNS
                                                                           3 FEB. 88 R.G.H.
                                                                    SCALE
                                                                           1:10,000
                                                                                        DWG No 3.3178-1-0110
```

LINE 31-0N

7_

$$0 - 8.4 78 / 2.1 4.0 9.7 (4.5 3.1 4.8)$$

FREQUENCY

FREQUENCY RANGE

1312 Hz

LEGEND

0	4lHz	STATION Nº
1	82Hz	SIMING N-
•	022	68.5
2	164 Hz	1
_		
3	328Hz	•
4	656 Hz	

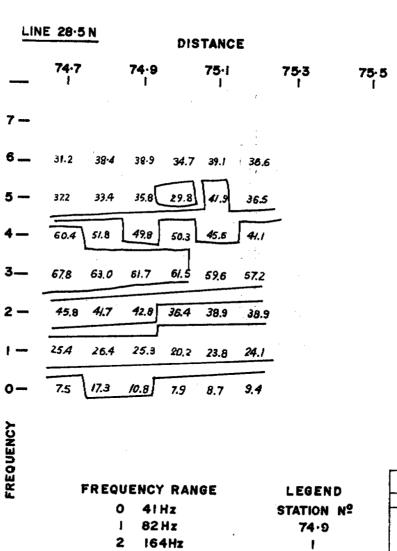
2624Hz Tx IN WEST

MOBIL ENERGY MINERALS AUSTRALIA

PROJECT WOOLNER

TILT ANGLE PSEUDOSECTION LINE 31.0 N

COMPILED	DATE DEC.82	L.J.S	ADDNS	DATE	BY
DRAWN	378983				
 SCALE	1:10,00	0	DWG No	3.3178	1010



328 Hz

656Hz 1312Hz 2624Hz

TX IN WEST

MOBIL ENERGY MINERALS AUSTRALIA
PROJECT WOOLNER

ELLIPTICITY PSEUDOSECTIONS LINE 28-5N

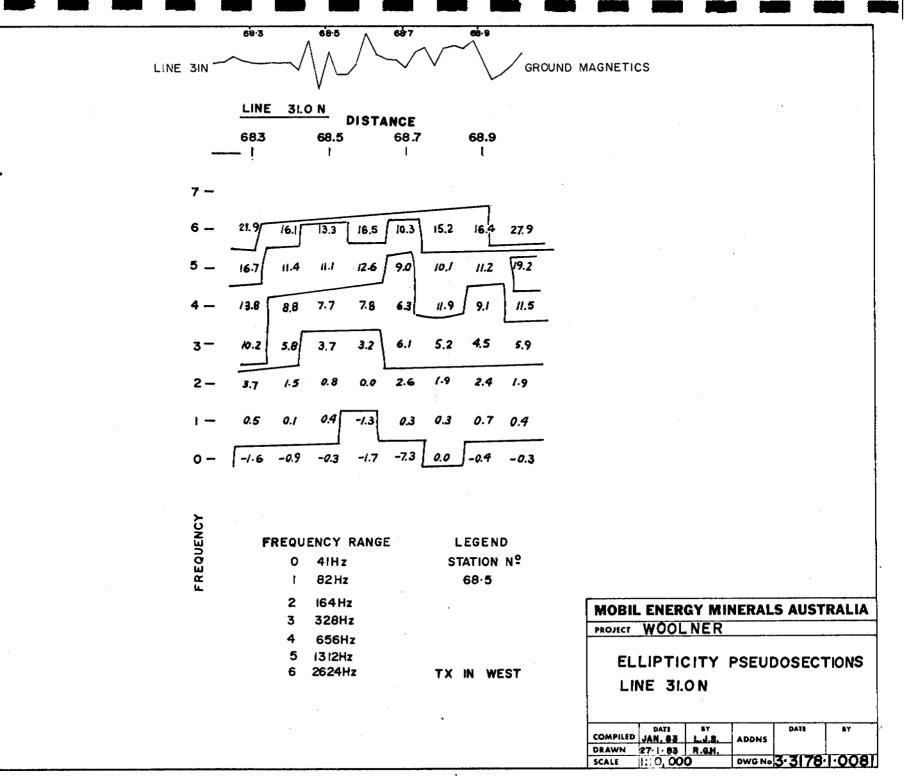
COMPILED	DATE JAN. 83	L.J.S.	ADDNS	DATE	87
	27-1-83		1		
SCALE	1:10,00	0	DWG No	3.3178	1.007

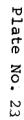
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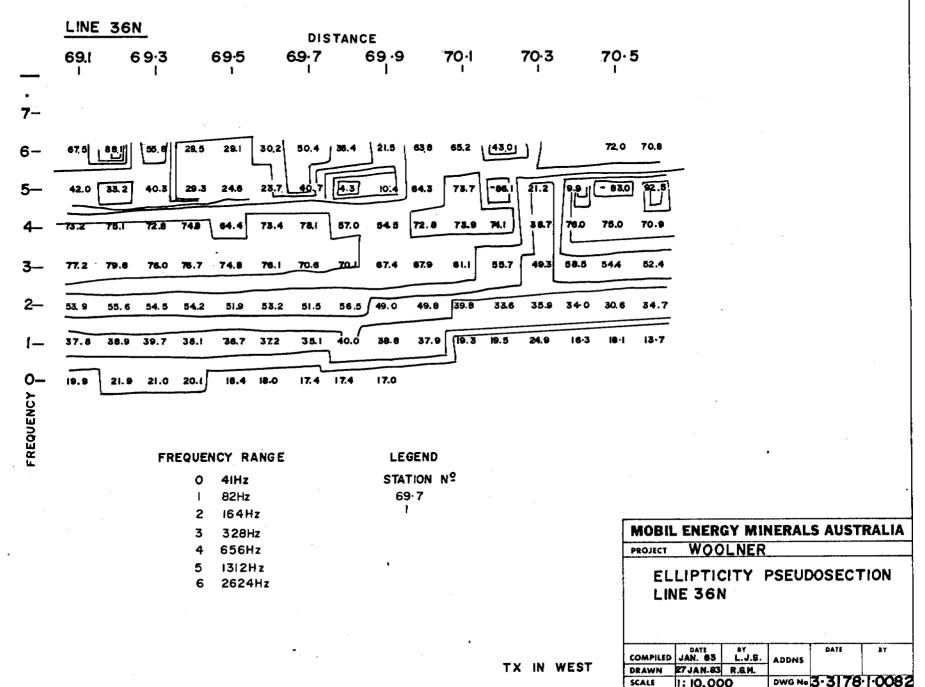
DWG No 3.31781.0080

DRAWN 28-1-88 R.G.H.

1:10,000







LINE 64.2E LINE 64.2 E DISTANCE 23-0 23.2 23.4 23.6 23 . 8 24.0 7 -5 _ 61.0 59.4 63.7 58.1 66.3 62.6 85.0 -83.3 -75.6 (-89-3 3--50.6 47.9 2 — 56.6 57.8 57.5 50.1 33./ 35.6 32.2 36.7 36.7 32.8 37.7 31.9 32.2 19.5 17.0 17.5 17.0 19.3 18.6 14.3 14.6 10.7 25.3 / 16.1 FREQUENCY RANGE LEGEND 41 H z STATION Nº 82Hz 23.4 164 Hz 328Hz 65 6Hz 1312 Hz 2624Hz TX IN WEST

236

23.4

23.8

24.0

GROUND MAGNETICS

23.0

23.2

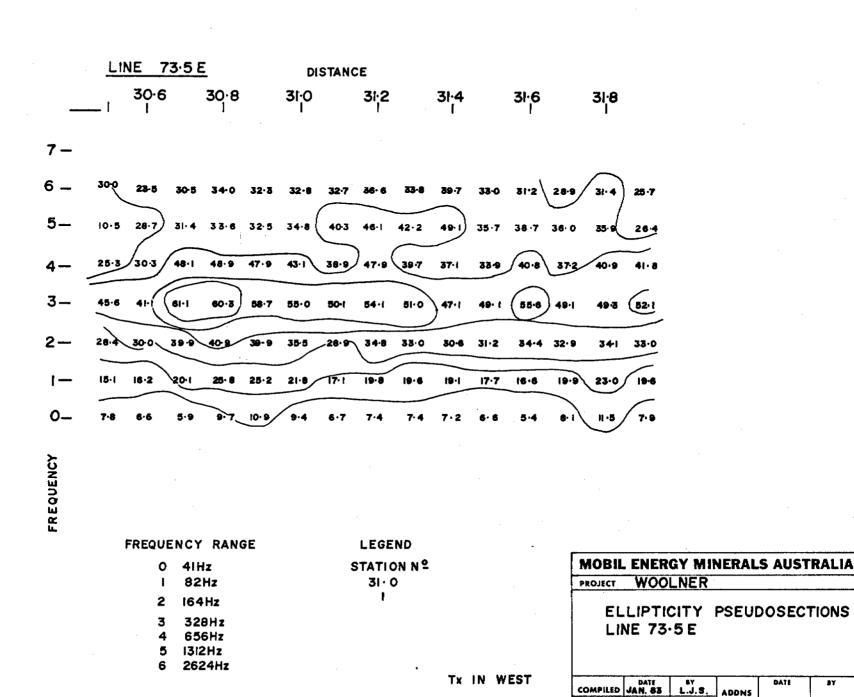
Plate No. 24

äY

DWG No 3

DRAWN 28 JAN 63 R.G.H.

1:10000



Pla

N

DWG No 3.3178 - 1.0075

1:10000

LINE 25:0N GROUND MAGNETICS LINE 25.0N LINE 26.0 N DISTANCE DISTANCE 66.6 66·8 67.0 67.2 66.6 67.2 66.8 670 7-7 3+Z 19-0 5 19.7 16.1 206 242 20.8 16.9 207 21-0 26-2 258 280 274 26-7 25-7 3 265 2 46.1 46.7 40 (51 0 **3-2** 3.0 FREQUENCY FREQUENCY FREQUENCY RANGE LEGEND STATION Nº 41Hz **MOBIL ENERGY MINERALS AUSTRALIA** 82Hz 66.8 PROJECT 164Hz 328Hz **ELLIPTICITY PSEUDOSECTIONS** 656Hz

1312 Hz 2624Hz

66.4

66-6

67.0

67.2

Tx IN WEST

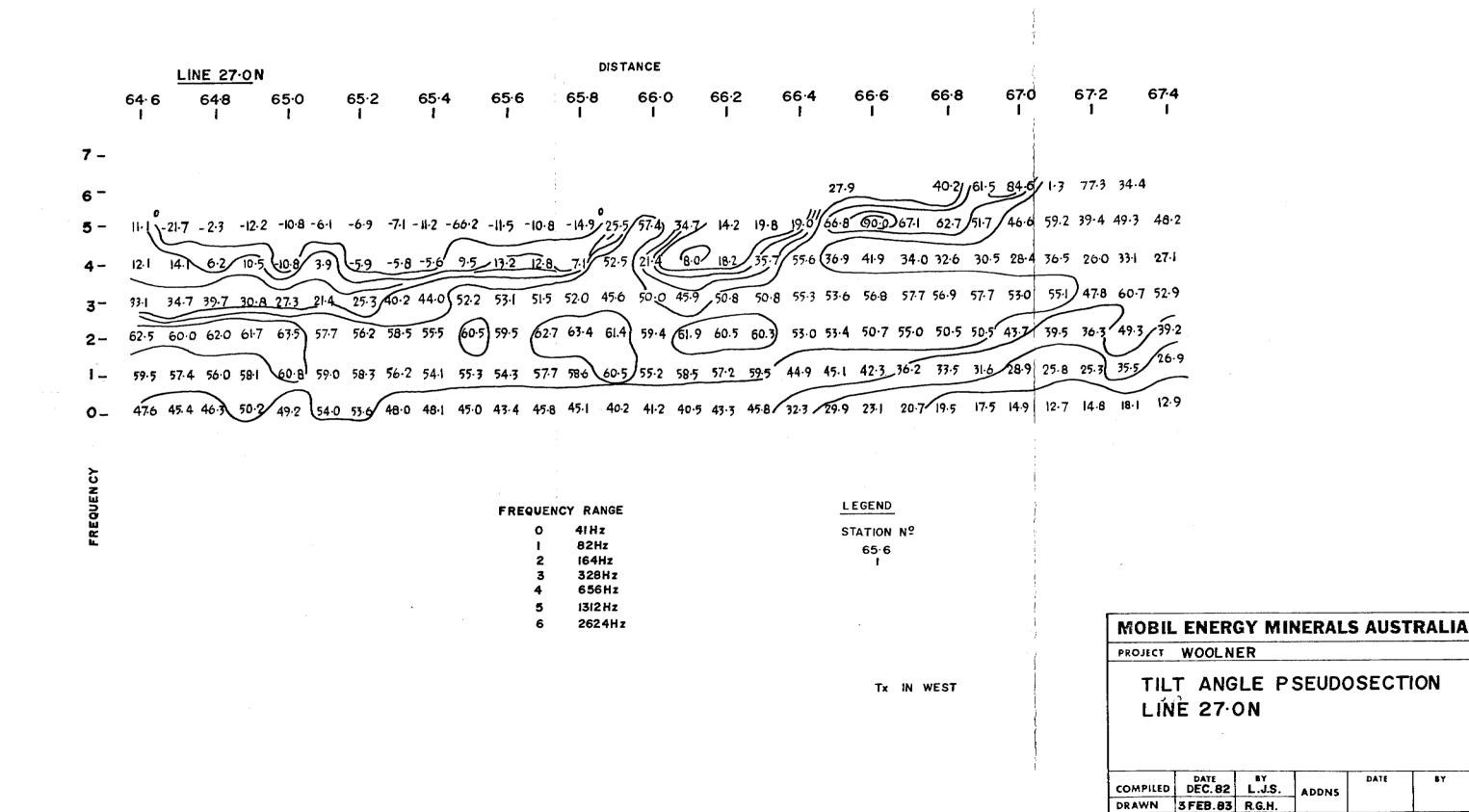
COMPILED

DRAWN

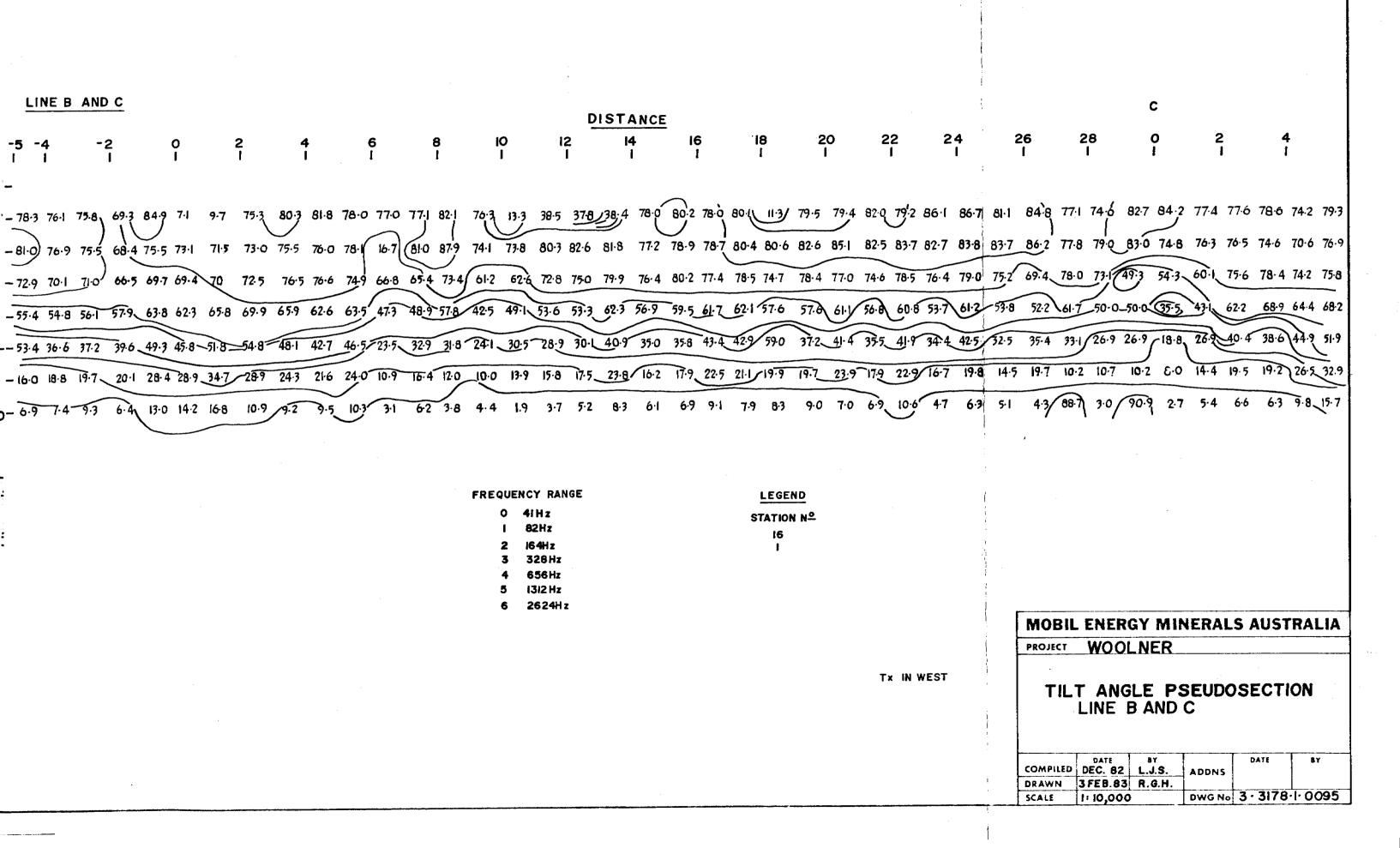
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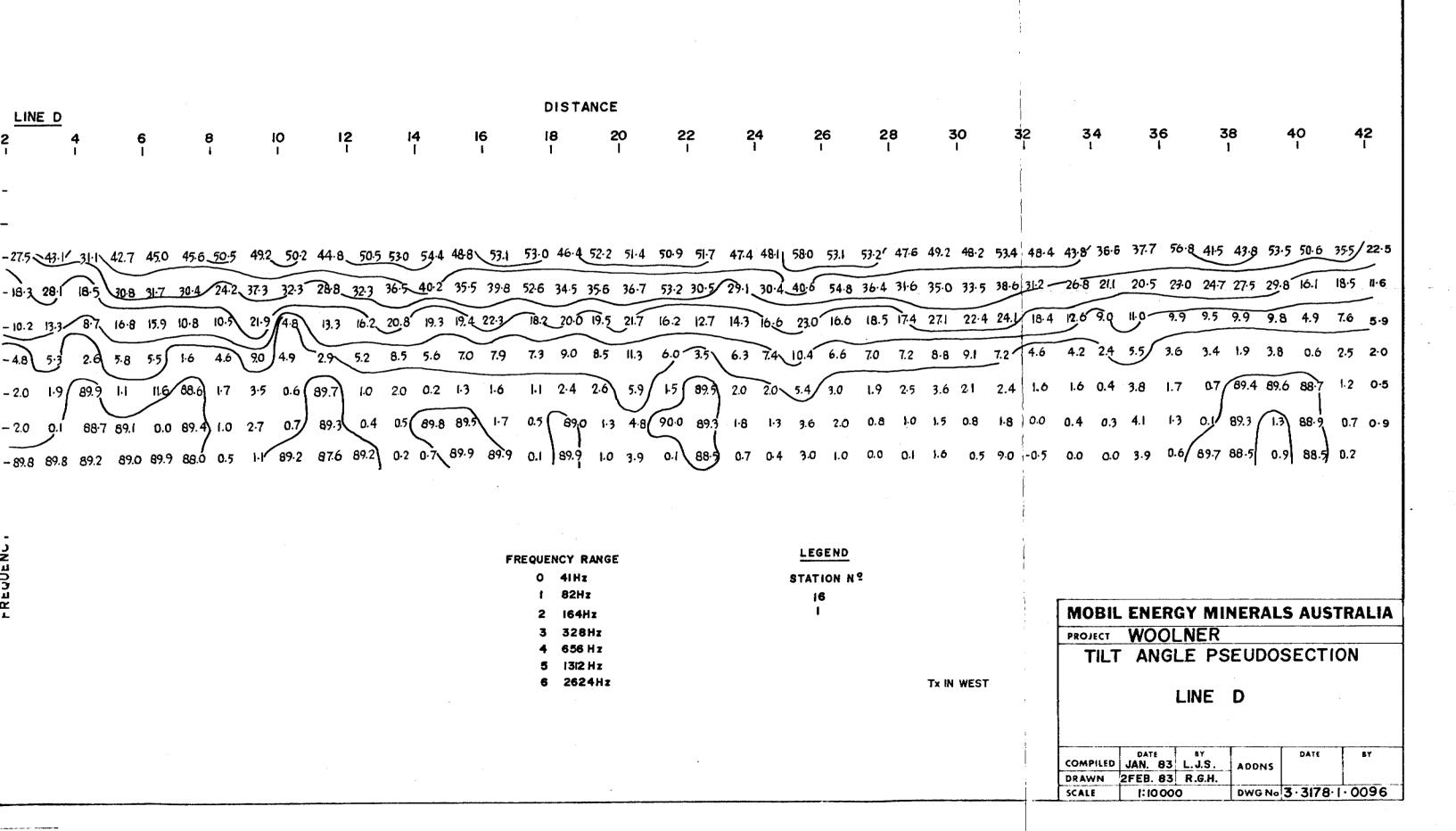
DWG No 3.3178 · 1 · 0104

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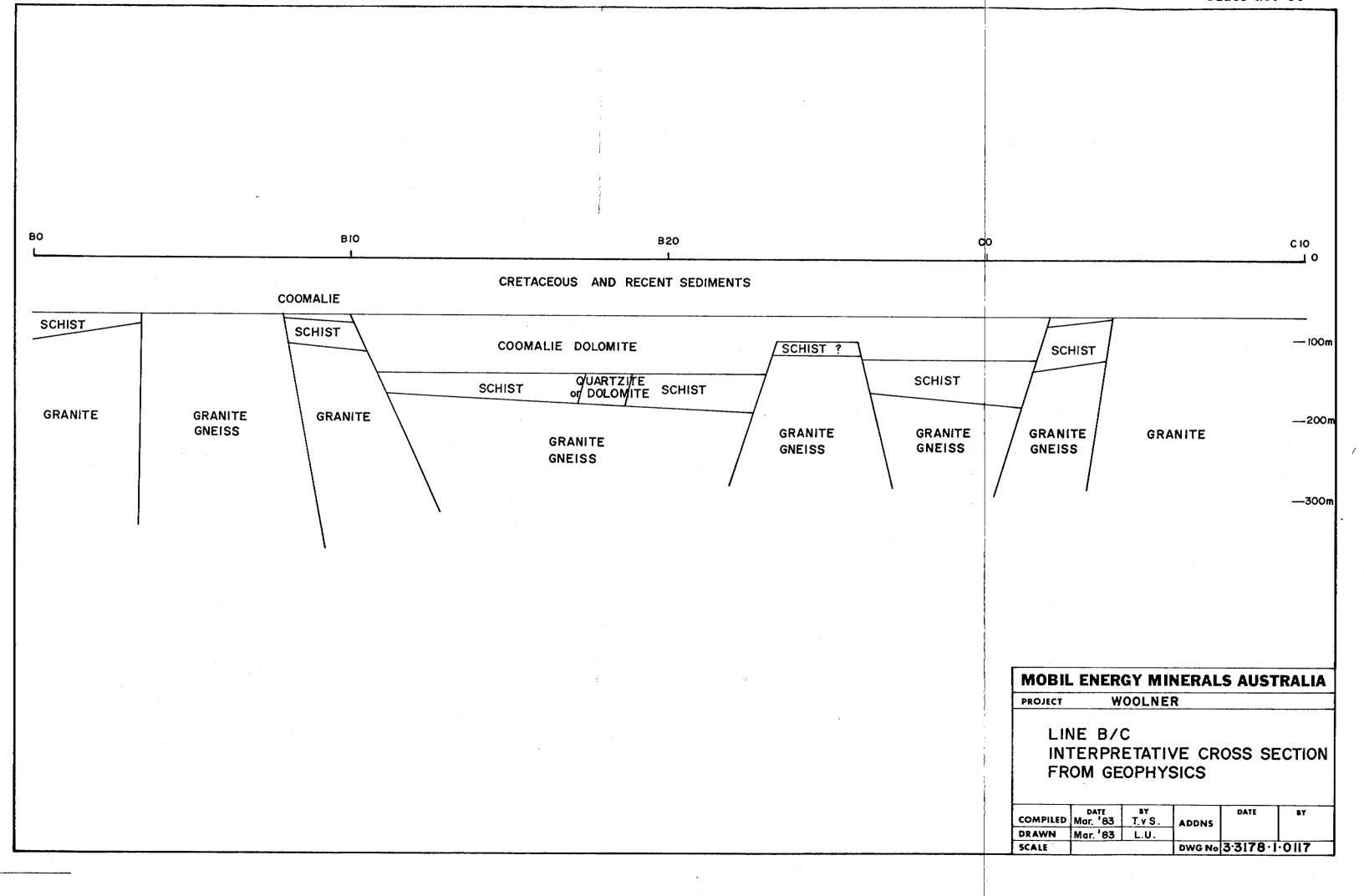
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	•	54 6	64-8		650	65'2	65-4	65.6	65.8	66 0	66 -2	66 -4	66.6	66-9	67·0	67.2	67· 4	
														-				
-																		
							·						•					
LINE 27.ON																		GROUND MAGNETICS
	LI	NE 27.	ON						DIS	STANCE						1		
		64.6	64.8		65.0	65/2	65∤4	656	65.8	66 -0	66-2	66 _: 4	66.6	66-8	67.0	67;2	67:4	
		į.	'		ı	,	ı	1	,	•	. •	•	•				• .	
7	7-															1)		
. 6	5		,										<i>13.7/</i> l		70.8 -77.2	$A \cap M$		
:	5-	4.6	6.4	2.8	4.5 3.1	2.8	3.4 6.2	6.0 4.0	4.0 2.5	85.A 77.3	84.4 76.2	78.8 84.6)	57.6 33.9	41.1 34-6	8-5 24.7 2	19.3	41.1 / (18-8	
4	4	11.4 17.	2 11.1	/3.7)	4.5 -3.7	12.6	((1.1	8.7 89.3	-81.5	83-2 89.5	87.8 87.9	85.1 (43.0	36.4 27.9	27.4 27.5	25.6 23.8	9.2 17.8	29.8	
:	3 —	-87-5 -87	7.9 -86-6	-86-2	-84.7 -83.	6 -845//6	0-2 87.2 8	0.9 79-4	81-7 81-6	78-2 80-6	93·9 83·2	78.5	54-2 46-2	40.9 38-3	36.0 31.0)	27.1	26.6	
:	<i>=</i> 2	69.3 63.	.9 62.2	68.7	69.8 77.5	5 74.9 6	53.3 63.6	9.4 53.5	54.1 54.2	50.9 53.8	56.3 55.3	55.4 35.0	34.2 \(25.4 \)	16.9 /8.0	3.7 /3.8	H.4 11.7	14.8 11.0	
	1-								35.0 34.4			/			6.9 5.7		7.3 4.1	
	•	21.2 20.	8 23.6	24,9	21.3 28.1	275 2	3.4 25.8 (18	3.7 /9.3	15.6 13.6	14.1 17.3	16.5 /5./	19.6 7.9	6.8 3.3	, 2.7	2.5 2.4	1.1 9.6	2.4 1.3	
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	EQUE																	
	F.								•			·						
				•					FREQUENCY I					LECEND				
									0 41Hz 1 82H:				S	LEGEND TATION N	<u>.</u>	.	MOBI	L ENERGY MINERALS AUSTRALIA
									2 164H 3 328H					65.4			PROJECT	WOOLNER
						·			4 656 5 1312	Hz				•				ELLIPTICITY PSEUDOSECTIONS LINE 27.0N
								,	6 262	4Hz							-	
															тх	IN WEST		
						•					-						COMPILE	DATE BY DATE BY JAN. 83 L.J.S. ADDNS 27JAN.83 R.G.H.
																	SCALE	1:10,000 DWG No 3:3178:1:0078





								DISTANCE	· • :									
Į	,	ιή	,	<u>ი</u>	=	in)	<u>i</u>	~	<u>0</u>	2	2	-25	27	-29	1	ю Ю	in H	
ı			1	•	•	·												APPARENT DEPTH
· · · · · · · · · · · · · · · · · · ·																		12 metres
	20	. A 76.7	10 4.2	86·3 -0·5	50 10 5 5 82 01 1-5	60 78·5 64·9/	0 50 0.7 ₁ 85.6	5 87 -2 86-8	85.9 83.8	11.7) 84.4	81·7 85·7\	8-6 3-5	81-2 84-2	82·9 15/3 6	36-3	84-6 2 5	į	17
į į	}		1		!!!	25.9 30.5	33.9 37.9	36-8 ¹ 80-4	82.7 86.1	83.8 82.0	81-1 83-0	73.3 84.8	81.2 83.5	80-8 Rn-8 c	3.3	81.6 2.1	0 5.5	24
	\smile		1 11			22 279	427 027	7 022 90.5	92.0//	/	85.4 86.2	63.4 05.2	0-1.2 01.6	00.6 11.0				34
	0440	22 267	25.2 -2.1	83-4 83	-2 83-8 84-6	81.2 82.3	81.3 83.5	82.7 89.2	9jiá 1.†//	90.0 88.3	88.0 0.9	88.9 89.2	87.4/-0-4	18/4/12				48
<u> </u>				0.2 0.3	2) 00 5 (07)	80.2 89.9	F6/ ⁶⁰ \86-6	6 85.0 81.6	79-8 (83-6	801/ 78.6	78.6 83.4	81.0 80.2	/ 18-3	2003 107				68 96
			66:5 65.0	572 5	9.2 69.2 73.1	67-1 65-6	64.4 63.	1 63.9 57.2	56-5 58-8	59-2 54-6	53.9 56.6	55-2 57-8	53.9 55.5	36.3 47.0	30'37 46.2	43.4	.9 46.4	136
412	46:3	507 46.8	44.3 46.9	45.2 46	67 47.6 52.5	46.7 42.2	41.2 42	1 42.9 35.6	36-2 37-7	37.4 33.5	33.7 33.3	33-7 38-7	7 3 1·5 33·	3 32.2/ 28.4	31:1 24.6	23-7 30	4 280	130
1 34.2	100								:									
									FRE	QUENCY RA	ANGE		LEGEND					
										0 41 H			STATION 15	Nō		MOBIL	ENERGY	MINERALS AUSTRA
										2 1641	Hz		i			PROJECT	WOOLNE	R
									v	3 328 4 6\$6						L	NE A HO	ORIZONTAL 200p
		.*								5 1312						T		E PSEUDOSECTION
									:	6 2624	PT14			Tx IN WEST				IN LINE MODE MITTER
							٠											BY DATE B

								:						Į						
<u> </u>	INE A	(<u>200 m </u> \$	SEPARATIO	<u>N</u>)				. :	DISTAN	CE				!						
_	2 I	4	6	8 I	10 1	12 1	14 1	16	18 1	20 	22 	24 	26 	28 i	30 I	32 I	34 I			
7-								1										FREQUENCY	APPARE DEPTH	INT H
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5-	31.7 (15.1	-1.6) (28.6) II	2.9 14.7 7.8	8 28.8 2	29.1 22.4	9 4.4 3.	3 4.9 3.	3 0.8 67	7 4.0 6.7	4.2 8.	8.5	0.5 12.6	7.2 8·8 U	.4/ 9.3	7.1 10.1	10.6 4.9	0.7 - 0.3	1312		103
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3-	10.3 (-1.4)	4.8 5.5 6	1.5 6.7 0.	7 11.4	11.4 11.8 10.	2 12.6 8	1.7/13.2 8	.5 6.4 12	2.9 12.0 (9.8.	_6.6 15.2	2 14.0	9.77.3	9.27.17	.4_7.3	22.4 27.4	25.8 36.7	30.1 34.5	328 164		!!! 2
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APPENDIX 2

FINANCIAL STATEMENT

Mobil Energy Minerals Australia Inc.

(INCORPORATED IN DELAWARE, U.S.A.; LIMITED LIABILITY)
(REGISTERED AS A FOREIGN COMPANY IN THE STATE OF VICTORIA)

BIN FLOOR, 31 QUEEN ST., MELBOURNE, VIC., 3000.

POST OFFICE BOX 4507, MELBOURNE, VIC., 3001.

TELEPHONE: 620 191 CABLE ADDRESS: "MOBILEMA" TELEX: AA 37000

STATEMENT OF EXPENDITURE FOR PERIOD JUNE 18, 1982 TO JUNE 17, 1983 EXPLORATION LICENCE 3478 - WOOLNER

Salaries and Wages (including on costs)	21,094
Professional Fees of Geologists and Consultants	27,430
Travel and Accommodation	6,180
Freight	2,176
Field Supplies	. 3,619
Communications	29 5
Equipment and Facilities	232
Hire of Field Equipment	15,503
Hire of Computer Equipment	1,9 53
Motor Vehicle Hire	742
Vehicle Operating Costs	2,627
Air Photographs	488
Airborne Surveys	1,875
Drilling Contractors	24,359
Laboratory Services	132
Contract Drafting Services	3,379
Other Contract Services	1,162
EDP Bureau Charges	9,284
Maintenance and Repairs	890
Power Generation	82
Motor Registration	297
Losses on Insurance	220
Miscellaneous Expenses	1,243
	125,262
Plus Head Office Overheads: 10%	12,526
	\$ 137,78 8
	•

J.H. Burnett

Acting Manager Accounting

APPENDIX 3

AGE DETERMINATION REPORT

Ion microprobe U-Pb dating of zircons from granitoids recovered in core from drill holes P4/1D, P11/1, P12/11 and P14/1, Woolner, Northern Territory

Prepared for

Mobil Energy Minerals (Aust.)

23 June 1983

I.S. Williams and W. Compston Research School of Earth Sciences Australian National University.

Ion microprobe U-Pb dating of zircons from granitoids recovered in core from drill holes P4/1D, P11/1, P12/11 and P14/1,

Woolner, Northern Territory.

I.S. Williams and W. Compston

Summary

This report documents ion microprobe U-Pb ages determined on zircons from four samples of granitoid recovered in drill core from Woolner, N.T. The isotopic analyses were made by the authors using the ion microprobe SHRTMP at the Research School of Earth Sciences, Australian National University, in co-operation with, and using heavy mineral concentrates prepared by, J. McAndrew (CSIRO, Sydney).

Two of the samples, P12/11 and P14/1, contained zircon with U-Pb compositions suitable for precise age measurements. These gave ages of 2660 \pm 20 and 2690 \pm 20 Ma, respectively. The compositions of the zircons from the other two samples were not suitable for measuring an age independently but were wholly consistent with these ages. As there is no evidence for a significant difference in age between the samples, the best estimate of the age of all is the mean of the above, 2675 \pm 14 Ma.

Analytical instrumentation and methods

a) Principles of ion microprobe analysis.

An ion microprobe is a specialized mass spectrometer used for determining the isotopic composition of materials on a micron scale. The sample is prepared as for an electron probe, as a polished thin section or grain mount. In the ion microprobe, it is bombarded by a beam of high-energy primary ions (oxygen in the present case) focussed to a spot less than 30 µm in diameter, thereby producing secondary ions of the sample from that spot for isotopic analysis.

Until relatively recently, with the exception of pioneering work by Andersen and Hinthorne of Applied Research Laboratories, California, ion microprobes have been little used for geological studies. The reason is that geological samples are generally chemically complex, making it technically difficult to distinguish the ions to be analysed from interfering molecular ions of virtually the same mass. SHRIMP is one of very few intruments with sufficient mass-resolving power to make this distinction, and is presently the only instrument that can do so without a great reduction in ion beam intensity and therefore in analytical precision.

b) Zircon U-Pb dating using SHRIMP.

U-Pb dating of zircon populations by mass spectrometric isotope dilution is well established as one of the most reliable methods of determining the age of rocks that have had a complex geologic history. Zircon is very resistant physically and chemically and even granulite facies metamorphism or host rock partial melting rarely completely resets its U-Pb age. The zircons in a single rock, however, commonly are not all the same age and single crystals often show zoning representing different episodes of growth. Single crystals and zones can be dated individually by ion microprobe, giving both the rock's age and some of its metamorphic history. SHRIMP, developed and constructed at the Australian National University, is presently the only ion microprobe capable of rapid, precise zircon U-Pb dating.

To date zircon, it is necessary to measure both the isotopic composition of the Pb and the elemental ratio Pb/U. On SHRIMP, the ratio of the ion beams produced by the isotopes of Pb is within 0.2% of the ratio of those isotopes in the target zircon, so Pb isotopic composition can be determined directly. Such is not the case for the ratio of ion

beams of different elements however, as different elements ionize with different efficiencies. It is necessary to correct the observed ratio $^{206}\text{Pb}^+/^{238}\text{U}^+$ to obtain the true ratio $^{206}\text{Pb}/^{238}\text{U}$ by relating the analyses of each unknown zircon to analyses of a standard zircon. By this method, Pb/U can be measured with a precision of better than 5%. Similar reference to a standard is used to determine U/Zr, for estimation of the U concentration, and Th/U.

The standard used for the present analyses was a piece of a Sri Lanka zircon, SLZ G3 Cp, in which the ratio $^{206}\text{Pb}/^{238}\text{U}$ has been determined accurately by isotope dilution as 0.0894. The ratio Pb/U in the Woolner zircons is referenced to this value.

Analyses of the Woolner zircons

The zircons from Woolner were analysed during one 15 hr working day. The data collected comprised analyses of the standard Sri Lanka zircon at the beginning and end of this period, and analyses of a total of fifteen spots on ten zircons from the four rock samples. The analyses are listed in Table 1 and plotted on a conventional Concordia diagram in Figure 1.

Because of time constraints, the analytical procedure was varied from that normally followed. Generally, we spend about 45 minutes in the U-Pb analysis of any one spot. Such an analysis consists of three determinations of all the isotopic ratios required, each determination being the mean of ten estimates of each ratio. This analysis in triplicate makes it possible to assess both the homogeneity of the spot and any systematic change in ion beam composition as the hole in the zircon is excavated. The five analyses in Table 1 which are normal triplicate determinations are the most accurate of the data listed. The remainder of the analyses were abbreviated to 15 or 30 minutes data collection, either because the

composition of the particular spot was not suitable for age measurement or beacuse an extended analysis was not considered necessary.

The zircons from Woolner (Plate 1) are mostly fine-grained (prism diameters generally <50 µm). The exception is sample Pl4/1, in which the zircons are euhedral crystals up to several hundred microns in length. In all samples, euhedral growth zoning is strongly developed, and cores (often unzoned) surrounded by thin zoned rims are common. Except for the large zircons in Pl4/1 it was not possible, with the 30 µm spot size of the ion microprobe, to analyse the zircon cores and rims separately. In the one case where separate analyses were made of a core and rim, Pl4/1 grain 12, no significant difference in age between the two was found.

The analyses listed in Table 1 are the means of from ten to thirty determinations of the isotopic ratios at any one spot. The concentrations of U and Th are calculated by reference to the standard zircon and assuming that the unknown zircons are stoichiometric $ZrSiO_4$. While the concentrations listed are approximate only, the ratios of Th/U are accurate to within 1%. The concentration data show the zircons to be generally richer in U than the average granitoid zircon (commonly ~ 500 ppm U), especially so for samples P4/1D and P11/1.

The remaining parameters listed depend upon a correct estimation of the zircons' initial Pb contents and compositions. The zircons from P4/1D are unique in our experience with respect to initial Pb content, up to 74% of the 206Pb in the zircon not being accountable for by radioactive decay of the grains' presently contained U. In normal zircons this value would usually be less than 0.1%. The initial Pb content of the zircons from P11/1 also is unusually high. The presence of large amounts of initial Pb makes estimation of the quantity and composition of the accumulated radiogenic Pb extremely difficult. Therefore, these zircons cannot be used to make an independent estimate of the age of these two samples.

The initial Pb contents of the zircons from samples P12/11 and P14/1 are considerably lower, though still above average. We believe that the high initial Pb contents of the zircons are an intrinsic feature and not surface contamination of the crystals during sample preparation. The reasons are, first, the common Pb content of the zircons, as indicated by the measured ²⁰⁴Pb/²⁰⁶Pb and ²⁰⁸Pb/²⁰⁶Pb ratios, showed no tendency to diminish as each analysis proceeded, that is, as zircon below the polished surface was exposed by the ion sputtering. Secondly, repolishing and recoating of mount P4/1D at ANU made no obvious difference to the observed common Pb content. This can be seen from the analyses of two spots on zircon 15 of that mount, spot 1 being analysed before cleaning and having 9.6% initial ²⁰⁶Pb, spot 2 being analysed after cleaning and having 14.6% initial ²⁰⁶Pb.

An important additional feature of the initial Pb in the Woolner zircons that is rare in our experience is that the initial Pb, in all except sample Pl4/1, includes a significant radiogenic component. Evidence for this is the constant excess of ²⁰⁸Pb in the zircons over that which can be accounted for by the decay of the Th presently observed, plus common ²⁰⁸Pb. Either these zircons incorporated radiogenic initial Pb generated in a high Th/U source when they first crystallized, or they have taken in a mixture of radiogenic and common Pb at some later time.

We lack the necessary information to calculate the exact composition of the radiogenic initial Pb and therefore to determine its origin. To estimate the composition of the Pb present when the zircons first crystallized it would be necessary to determine the composition of the Pb in minerals with high Pb/U ratios that formed at that time, for example, sulphides or potassium feldspar. To estimate the possible composition of labile Pb that may have entered the zircons in the past, but significantly later than they first crystallized, it would be necessary to know more of the metamorphic history of the granitoids and to search for other minerals that may have

incorporated the labile Pb at the same time. Possibilities would be epidote, micas, or secondary sulphides.

Even though we cannot determine the composition of the initial Pb uniquely, we can impose some limits on it. This can be done most readily in the case of sample P4/1D, which has the highest initial Pb content, although the composition range as determined applies only to that sample. Figure 2 shows the isotopic compositions of the Pb in the different P4/1D zircons as measured, plotted as 204Pb/206Pb and 207Pb/206Pb. both ratios which are very sensitive to initial Pb but independent of Th/U. The analyses form a nearly linear array with an extrapolated intersection on the 207Pb/206Pb axis of 0.185. The latter we infer to be the composition of the Pb produced in situ by U decay, which gives an apparent age of 2700 Ma. This is consistent with the ages measured on samples P12/11 and P14/1. The array does not extrapolate upwards to the composition of common Pb at 2700 Ma however, indicating that a third, radiogenic component is present. From the analysis of grain 6, we infer by extrapolation of line A that the 207Pb/206Pb value of the additional radiogenic component is at least as high as 0.31. This in itself places a limit on that component's origin: it must be older than 2200 Ma. even if it was separated from its parent U immediately that it formed. As it is more likely that the Pb evolved in contact with its parent U for a significant period of time before becoming incorporated in the zircon, a two-stage evolution concept is more realistic. The Pb composition does not define this evolution uniquely, but unless the 207Pb/206Pb of the radiogenic component exceeds 0.43, it does imply that the Pb was not incorporated in the zircon when it first crystallized. Assuming 0.31 and using a two-stage model, the age of the Pb depends upon the time it is assumed to have been separated from its parent U and vice versa. the Pb evolution stopped yesterday, then it started 3500 Ma ago. If it started 2700 Ma ago, namely when the granitoids crystallized, then it

stopped about 1800 Ma ago. Considering what is known of the history of the area, we consider the latter to be likely and have assumed such an evolution in calculating the P4/1D zircons' radiogenic Pb contents.

No matter what the initial Pb content or composition of the P4/1D zircons is, the isotopic analyses show that some of the zircons, 4 and 6 especially, are extremely deficient in the radiogenic Pb produced since crystallization. They contain only about 10% of the radiogenic Pb that their present U would have generated in 2700 Ma. Given general experience with zircon, we attribute this deficiency to a major loss of radiogenic Pb rather than a recent gain of U. The pattern of discordance in Figure 1 shows that this loss of Pb occurred comparatively recently.

The balance between the zircons' U and radiogenic Pb contents is well illustrated on the Concordia diagram, Figure 1. The graduated curve represents the present day Pb-U (daughter-parent) ratios of closed isotopic systems of different ages. It is general experience that zircon analyses may fall slightly below the curve, that is, they show a small radiogenic Pb deficiency, in which case the zircons' age is determined by extrapolating the line defined by the analyses (the discordance line) to the Concordia. The zircons from samples P12/11 and P14/1 have low initial Pb contents and plot close to the Concordia, giving relatively precise intersection ages of 2660 ± 20 (2 σ) Ma and 2690 ± 20 Ma. samples from P11/1 and most from P4/1D have high initial Pb contents and plot far below the Concordia, making the intersection age extremely uncertain. Data for these samples do, however, plot near to the extrapolation of the discordance line for the P12/11 and P14/1 zircons and we have no reason to suggest that they differ in age. This conclusion is strongly supported by P4/1D analysis 15-1, which plots close to the P12/11 and P14/1 data.

We conclude, therefore, that all the samples analysed are, within our uncertainties, the same age. Our best estimate of that age is

2675 ± 14 Ma. Three of the four sets of zircons contain radiogenic initial Pb, which is unusual in our experience. In some of the zircons, the initial Pb contents are extremely high. Evaluation of sample P4/1D shows that, in that sample, the initial Pb is older than 2200 Ma. It is likely that it was incorporated into the zircon at or after that time.

Table 1. Ion microprobe age determinations on zircons from the Woolner area (N.T.), submitted by Dr. J. McAndrew on behalf of Mobil Energy Minerals (Aust.).

Rock	Grain-spot	ppm U	Th ppm	Percent initial Pb	206 _{Pb} *	207 _{Pb} *	207 _{Pb} *	207 _{Pb} * 206 _{Pb} * age(Ma)
70479b,	_	311	135	8.9	0.271	6.67	0.1787	2640
P12/11	4 - 25	348	242	0.6	0.499	12.62	0.1833	2680
	3 - 1 [§]	578	245	1.9	0.527	13.25	0.1822	2670
	3 - 2 [§]	475	354	3.4	0.473	11.69	0.1793	2650
70471,	4 - 1 [†]	2749	570	71.9	0.0475	0.685	(0.105)	(1710)
P4/1D	6 - 1 [†]	2412	319	73.9	0.0456	1.88	(0.299)	(3460)
	15 - 1 [§]	392	275	9.6	0.432	11.38	0.1911	2750
	15 - 2 [†]	1099	494	14.6	0.287	6.76	0.1711	2570
	16 - 1 [†]	2776	1554	49.6	0.0793	2.64	(0.241)	(3130)
70478,	5 - 1 [†]	1445	729	17.8	0.197	4.76	0.175	2610
P11/1	10 - 1 [¶]	1544	1873	5.8	0.181	4.72	0.1891	2730
70492,	11 - 1	370	149	2.3	0.498	12.52	0.1822	2670
P14/1	11 - 2	311	122	0.5	0.465	11.68	0.1823	2670
	12 - 1 core	956	389	0.1	0.469	11.96	0.1848	2700
	12 - 2 [¶] rim	365	122	0.8	0.476	12.24	0.1867	2710

Decay constants: $\lambda^{235} = 0.98485 \times 10^{-9} \text{ yr}^{-1}$, $\lambda^{238} = 0.155125 \times 10^{-9} \text{ yr}^{-1}$

ANALYSTS: I.S. Williams,

W. Compston

Australian National University

s normal triplicate analysis

[¶] duplicate analysis

[†] single analysis

^{*} corrected for initial Pb

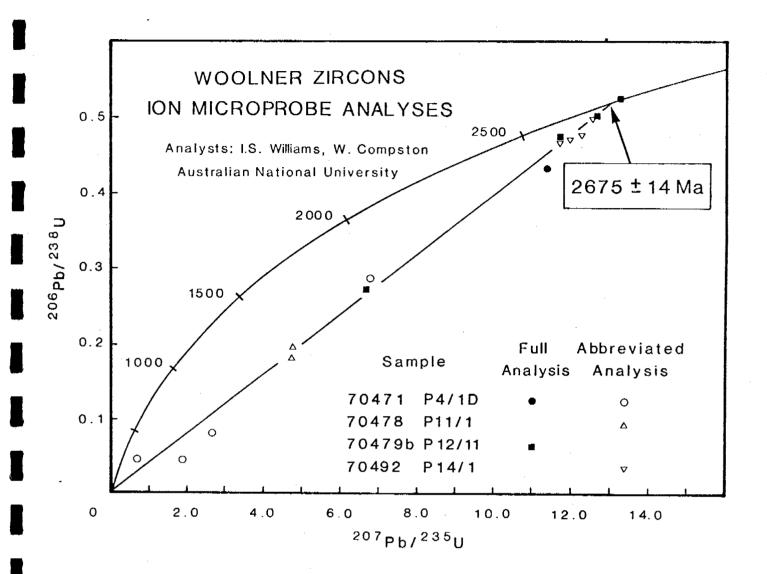


Figure 1.

Concordia diagram showing ion microprobe U-Pb analyses of zircons from granitoids intersected in four drill holes at Woolner, N.T. Only the zircons from holes P12/11 and P14/1 are suitable for precise age determinations. The mean age of these is 2675 ± 14 Ma.

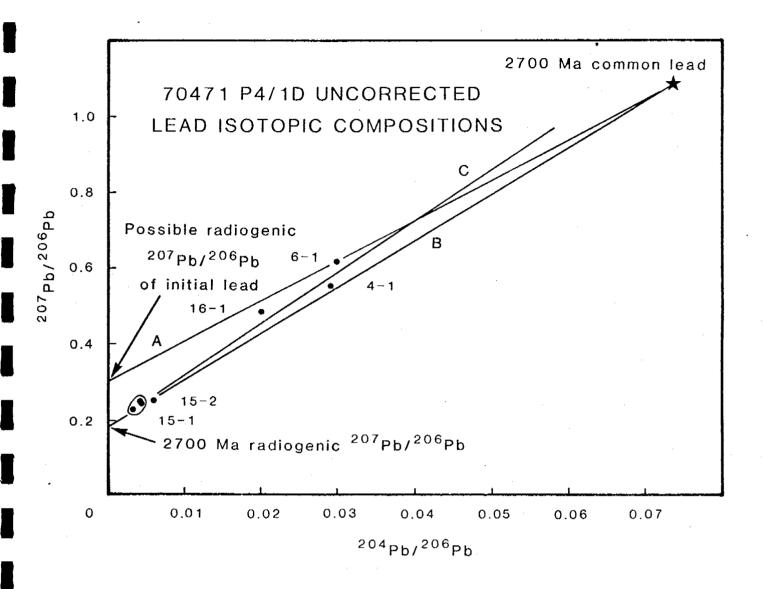
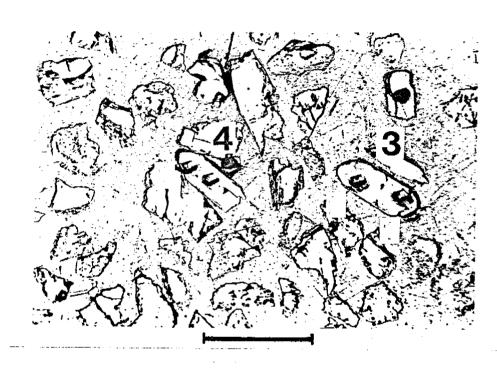


Figure 2.

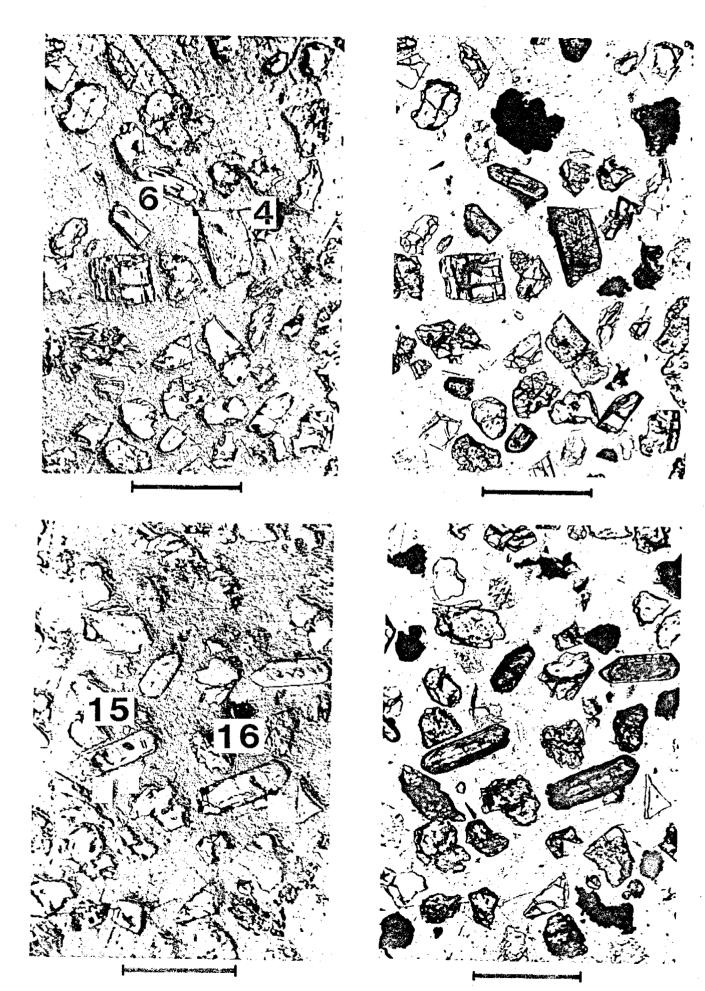
Pb compositions for zircons from sample 70471 (P4/1D). The mixing line defined by the analyses (C) deviates significantly from the mixing line between 2700 Ma common and radiogenic leads (B). The 207 Pb/ 206 Pb of the initial radiogenic Pb must be at least as high as indicated by the y-intercept of a line linking the composition of grain 6 and common Pb (A).

Plate 1. Photomicrographs of analysed grains. The photographs are in pairs; one transmitted light, one reflected light with Nomarski differential interference contrast. The numbers on the reflected light photographs identify the analysed grains (red) and spots (blue). Scale bars 200 $\mu m.$



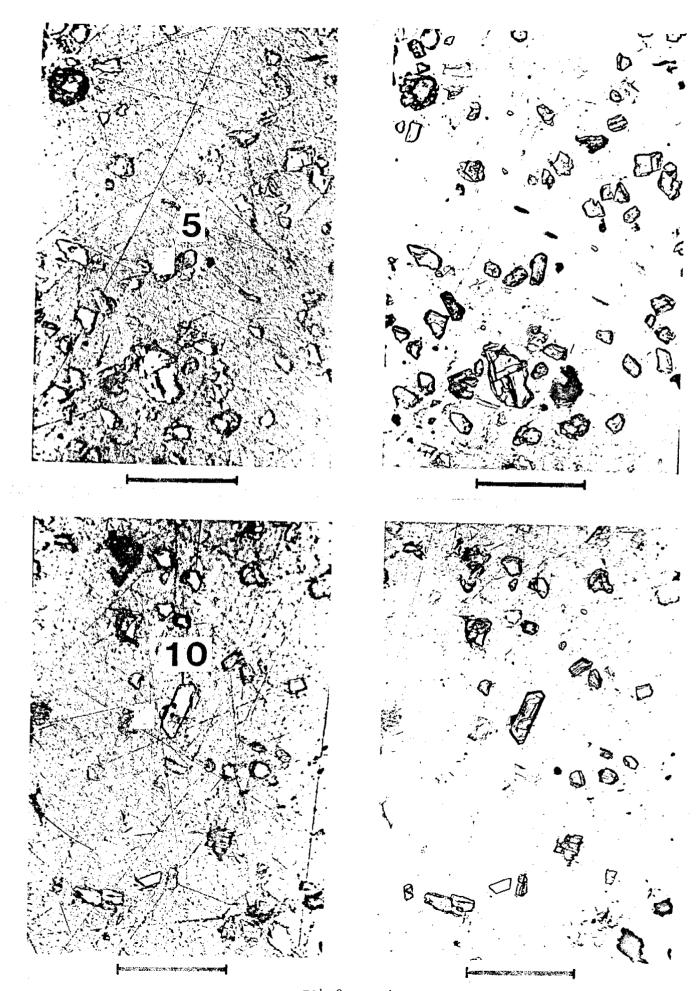


70479b, Pl2/11



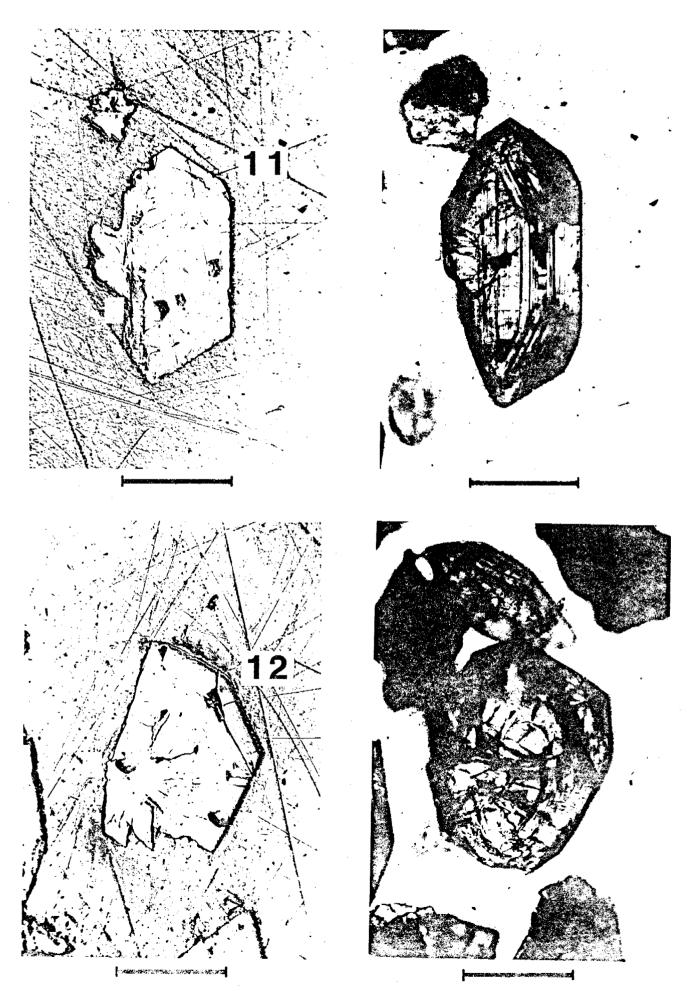
70471, P4/1D

Plate 1 cont.



70478, Pl1/1

Plate 1 cont.



70492, P14/1

APPENDIX 4

DRILL LOGS

DRILLING LOG

HOLE NUMBER WD1

				LOCAL B19 HOLE No. 170.30E/33.68N HOLE No.	WD1	. 7
- PHOJE	T NATE:	WO	OLN	ER COLLAR R.L	Aercica	<u>.</u>
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HTYESO	Length Drilled	L/Ý	.x	See footnotes for information required.	MUMEER & TYPE	respuic
1	BL.			Yellow brown med. grn. sand.		W
2	D E D		<u> </u>	Yellow sand with laterite pebbles		7000y
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4				Grey to black mud	•	~ E
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8				n		₁ G
9	·			п		<u>B</u>
10				-15		~
11	-			Cream white sandy clay		اح ح
12	F)			u .		-C
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14) -0 -0			" and thin interbedded pale grey		開
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16				" /washod sample gives only med grained		Ę-
17				(washed sample gives only med. grazion		<u>'C</u>
18				" rounded white quartz)		<u>c</u>
- 19				n		├ <u></u> ं
20						
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	$G \dots MO1$			000 CORE/HOLE SIZE & LENGTH Pre-collar 0-62 CASING SIZE & LENGTH CORE 62-82.2	PACE L OF	· 5.
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22				TI .		·c					
23				11		·c					
24				11	1	C.					
25				11		C:					
26				11		ر ار					
27				"		C					
28				II		C.					
29				11		·c.					
30				11		c					
31				n							
32				n		C					
33				п		· · ·					
34			<u> </u>	"		, C					
35				11	<u> </u>	-c					
36				u	·						
37				11	<u> </u>	• <u>č</u>					
38				"		6					
39				u ´		C					
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IG TYPEA RILLER: TART:	CO:	 		CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH VATER TABLE	PACE 2 OF	.5					
	COlc incl	our &) L. miti	ROCX	NAME: degree of weathering or alteration, proportion, grain size & roundness tedding (quantitative, angle to LCA), achistosity (ditto), fractures/joints (c); ORE MINERALS (%, grain size & numes); VISUAL ESTIMATE OF CRADE.	of mineral litto), brec	compos ciatio					

	NWE/NO. IT NWE:			CHORDINATES LOCAL HOLE N R AMG ESARIM	•	
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43				11		(C
44				n e		6
45				TI .		c · c
46				11		
47				Increasing clay content with depth.		C
- 48				n .		C
49				u .		1.0
50			,	11	:	<u>ر</u>
51	· -			***		C.
52		·	·	и		C
53		·		n		C.
54				11		()
55				n .		7
56	·			11		C
57				11		10
58		-		11		· C.
59				11		C
60				II.		C
C TYPE/ ORILLER:	CO:	· • • • • • • • • • • • • • • • • • • •	•••••	CORE/HOLE SIZE & LENGTH CASIDIG SIZE & LENGTH	LOG BY	
NISH:	ICN: Cold	our &		NAME: degree of weathering or alteration proportion emin air t common	HOLE T.D.	Compare -
				bedding (quantitative, angle to LCA), achistosity (ditto), fractures/joints (c; ORE MINERALS (X, grain size & names); VISUAL ESTIMATE OF CRADE.	litta), br e c	ciation.

TITLE	۸۳۳۷۱۱۰۰.	:	· • • • • •	CO-CHDINATES LOCAL B19 HOLE No	WD1	
PROJEC	T NAC:	WOO	DLNI	ER AMC BEARING	TION	
אדיכ₃נו	Longth	RECO	VERY	DESCRIPTION:	SAGLE	Fraght
	Drilled	L/X	z	See footnotes for information required.*	MUMBER & TYPE	Log
51	O OF POE ECOIPA					C
52 	3 2 2		-	Weathered silicified dolomite-clay and qtz frags.		
63 				Weathered and silicified cream dolomite- broken		
			<u>0</u>	core - qtz fragments after silic. dolomite	•	
54.2			9.0	Minor kaolin and dolomite		
				Pale creamy yellow dolomite with interbedded talc		
				chlorite schist. Dolomite silicified in part,	·	
	NQ			generally along conformable regular shaped bands		
				2cm to 10cm wide (20% silicified). Stylolites.		1
				Pale green finely foliated talc-chlorite schists		
	-			67.60 - 67.90m, 5° to core axis.		片
69.30				Pale green carbonate-chlorite-talc schist,		
	-		-	foliated, gold sheen on foliation planes, finely		
	-		%0	disseminated sulphides along foliation and pale		
			ŏ	pink hematitic alteration after 69.60.		
9.90	·					
				Grey white dirty dolomite with thin schist		
70.5				interbeds		17
6.60		_		-		-
TYPE A TULER:	10DEL:	· · · · · · · · · · · · · · · · · · ·	•	CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH WATER TABLE		5.
	CN: Coto incl	Ar & S - matr	ROCK I	NAME: degree of weathering or alteration, proportion, grain size & noundress of pedding (quantitative, angle to LCA), achistosity (ditto), fractures/joints (ditto); ORE MINERALS (X, grain size & names); VISUAL ESTIMATE OF GRAPE.		_

				CO-CROINATES LOCAL HOLE N R AND COLLAR B. 1	o.: WD1	
PROJEC	J Nƌ:	WOO	LNE	R AMC BEARING ODLLAR R.L. INCLING	T	-
אדיבשו	Length	RECO	VERY		SAMPLE	brapt
<u>-</u>	Drilled	LN	<u> </u>	See footnotes for information required.	MU-BER L TYPE	Log
76.6				Silicified dolomite - massive white qtz		-
78.9				fractured in part		
			%00	Pinkish dolomite - hematitic alteration along		
				fractures - 2cm green chlorite schist band		
				79.15m - partial silicification throughout		
82.2						
				END OF HOLE		<u> </u>
				Pvc tubing (32mm) inserted to bottom of hole.		
				Tubing not split.		ļ
					<u> </u>	_
			·			-
	-					-
		_				
						-
						_
					1	-
						-
EUT DK (l ∞:				LOG BY	<u> </u>
TYPEAR RILLER:	10DEL:	• • • • • • • • • • • • • • • • • • •	·	CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH	:	
ISH:		•	•	VERTICAL SCALE	HOLE T.D.	

IRIPTION: Colour & ROCK NAME: degree of venthering or alteration, proportion, grain size & roundress of mineral compone incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), breclation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF CRADE.

DRILLING LOG

HOLE NUMBER WD2

DEPTH	T NAME:	RECO		COLLAR R.L. INCLINA DESCRIPTION:	SAMPLE	Graph1
	Drilled	L/V	.*	See footnotes for information required.	MUMBER & TYPE	Log
1				Black soil clays (edge of lake)		W
2				Red-yellow-white mottled sands		000
3				Grey clayey sand and laterite		000
4				Grey to white sandy clay	÷	ودر
5				11		c.
6				Yellow and white sandy clay - increasing	·	YC.
7	-			clay content with depth		Yc.
1						· Yc.
22						·4c.
1				Yellow clayey sand-coarse qtz grains-rounded-		· ·
27		_		decreasing clay content with depth		نزر.
↓ :				Med. grn. rounded qtz grains in grey clay		ΔΔ.
31				(lithified?) matrix hard drilling (kaolinite)		Δ.Δ Δ.Δ
1				Pale yellow fine-med. grn sand minor clay		·Δ. Δ.Δ
41		_		throughout		;
				White amorphous carbonate		
49						
				Carbonate plus minor pale green to grey sandy		
53				clay: increasing clay content to 53m		9c.
						gc. gc
	ο. Gad			1000 CORE/HOLE SIZE & LENGTH Pre-collar 57m	LOG BY B.	Ŗ.

PERPOSE: EM Anomaly RESULT: Dolomite - negative

*DESCRIPTION: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral component, incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

חחדוו	* ヘヘ	TO CO DOM
DRILL	LUG	ruru

TITLE	NAME/NO.		<i>-</i>	CO-CRDINATES LOCAL HOUSE	WD2	
				NED AMG	KG KG KATION	•••••
DEPTH	Length	RECCO	VERY	DESCRIPTION:	SAMPLE	- Craphi
	Drilled	L/V	7	See footnotes for information required.	MUMBER & TYPE	Log
<u> </u>						1.
54	4			Coarse angular qtz chips minor white clay -		Π
1	g '		, , ,			1
57	PRE			from silicified dolomite fragments		
	NQ			Fine white qtz sand with fragments of weathered	*	1
60.3	Core			silicified dolomite		
				Yellowish 'dirty' dolomite - finely laminated	·	1
-				in part, lamination 10° to axis, pale green		
				clay and fuchsite along fractures and lamellae	·	
69.5				- minor silicification		
				Brecciated dolomite - coarse breccia (fragments		
<u>-</u>				up to 3cm) with hematitic alteration along		1
				matrix lines. Green clay and fuchsite, also	. :	
71	-			minor silicification		
	-			Creamy massive dolomite - fine lamination,		1
				stylolites - cross cutting silicification		
81.3				around 71.3 and 79.3m		1
вон						
				9m of casing and shoe bit lost down hole		
-				(48-57m)		
				Pvc tubing placed in hole		
BUITING (æ:	· · · · · ·			LOG BYB.R	•
IG TYPE/N DRILLER: START:	· • • • • • • •	•		CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH WATER TABLE	PAGE 2. OF	2
TNISH:					HOLE T.D	

SCRIPTION: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral components. Incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

DRILLING LOG

HOLE NUMBER WD3

			•			
	.•			DRILL LOG FORM		
	NAME/NO. CT NAME:	- •	OLNI	CO-CROUNTEES LOCAL C./ AMG. 71.82E/33.72N HOLE N EEARLY COLLAR R.L.	ю	
	1		VERY	DESCRIPTION:		
DEPTH	Length Drilled		T	See footnotes for information required.*	SAMPLE MUMBER & TYPE	LOE
1				Grey soil, in. white and yellow sand		(F)
₩				Mottled red-white-yellow clayey		000
3				sand		0
↓				White clayey sand - fn grn rounded qtz		CG
22						.c
23	·			Laterite material		000
2 4		,		White and yellow clayey sand		· 40
				Grey to white sandstone with hard grey		AZ
Ē				(kaolinite) matrix (hard drilling -		in
34	Collar			small flaky chips)		ν·τ ·V.
	EF.			White and pale yellow med to coarse grn qtz		
Y .				sands - clay rich in part		
47						
	_			Pale green clay and coarse qtz sand		9c :
				(rounded grains)		gc
52		_]:9c
-				Coarse qtz sand - angular chips (silicified	<u> </u>	
£ ↓	<u>.</u>			dolomite fragments)		广
<u> </u>						
56						
	œ: Gạ¢			******	LOG BY.B.	R
RIG TYPE/	MODEL War			CASING SIZE & LENGTH NO Core 56-75.3m	PACE 1 OF	
START:	17	/11/	8.2	VATER TABLE 3m VERTICAL SCALE RESULT: Negative - dolomite	HOLE T.D.7	5.3m
BURPOS	ON: Colo incl (qua	millo ur & ; . mati ntita	ROCK rix; tive)	y. RESULT: Negative - dolomite NAME: degree of weathering or alteration, proportion, grain size & roundness bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (d ; ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF CRADE.	of mineral c litto), brecc	:Ompone :iation

				CO-CRDINATES LOCAL C7. HOLE No.	WD3	
PROJE	CT NAME:			R COLLAR R.L. BEARING INCLINAT	ION	
DEPTH	Length Drilled		VERY	DESCRIPTION: See footnotes for information required.*	SAMPLE NUMBER ATYPE	Craph!
56				Dolomite - generally grey with yellow-orange		H
<u> </u>				alteration in part (weathering)	•	ac
K R			100	%) Minor silicification. Lamination at low angle		
Cok				to axis 10-30°	÷	1
Ø				60.6-70m - pale grey green clay (weathered		H
		,		schist)	•	
75.3						1
ЕОН				Pvc inserted with bottom lm split and wrapped	<u>.</u> .	
		-		in alfoil.		
· 						
	:-					
				•		-
			_			
	<u>.</u>	_	_			
			_			
TYPE/	DDEL:		•	CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH P	OG BY	
	•••••••				DLE T.D	

*DESCRIPTION: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral components incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

DRILLING LOG

HOLE NUMBER WD4

DRILLING LOG

HOLE NUMBER WD5

	E NAME/NO. ECT NAME:	τ.		LOCAL B10 LOCAL 69.45E/33.55N HOLE NO. NER COLLAR R.L. INCLINAT		
DEPTH	1	1	VERY	DESCRIPTION:	SAMPLE NUMBER	Craph Log
	Drilled	157	*	See footnotes for information required.	& TYPE	+-
1		<u> </u>		Grey sand		(4)
2				Yellow sand with minor gravel pebbles		9,3
3				(dune deposit)		-/4
4			•	Coarse laterite pebbles & grey-black clayey sand	1	00
- 5				Grey sand, coarse laterite pebbles and shell		0.00
6				fragments		:0.
7	-7			Grey sand and black vegetable material		₽́_
8	:					P-
<u> 9</u>				Purple red sands, red clay, fine laterite pisolit	.es	0
10		•		n		0 6
<u> </u>				Laterite - laterite chips, small pisolites		0 0
12				Laterite & mottled clays (mottled zone)		C
13			·	White-yellow-orange mottled sandy clays (mottled	-	8
1				zone	, -	SE
21				Silcrete zone - mottled lithified zone - of med		(. \(\D \)
<u> </u>				grn qtz grains in a mottled siliceous matrix		. A Al.
<u> </u>				- iron rich in part, breccia fragments, chert		β. Δ. Δ. Λ
				fragments, minor mottled clay		D.C
30				Grey matrix- supported sandstone- med grn rounded		Ÿ.
<u> </u>				qtz grains in a grey hard (siliceous) clay (kaolin) matrix-(hand drilling)		Δ.
DRILLING	co: Gaç	len. arma	Dri n 1	lling 000 core/HOLE SIZE & LENGTH NO 83.2-102.4	ьс вуВ	R
DRILLER:	G. MOI	rıs	; 	CASING SIZE & LENGTH -0-83m - Precollar WATER TABLE 4m	PAGE 1 OF	.3.
ĐƯƯƯ	22/11 SE: EM	Ano Xur &	ntal Rock i	VERTICAL SCALE Y RESULT: Negative-dolomite VANE: degree of weathering or alteration proportion and other to the state of		
				proportion, grain size & roundness of codding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (dit ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.	mineral o	:ompone :iation

	FORM

				CO-CRDINATES LOCAL B10 HOLE NO. AMC BEARING VER COLLAR R.L.		,
	1		VERY	INCLINA		
DEPTH	Length Drilled		1	•	SAMPLE NUMBER 4 TYPE	Craph Log
31			Ĉ	Grey matrix supported sandstone		. Ψ.
1						Δ
42					# · · · · · · · · · · · · · · · · · · ·	Δ.
₩				Pale yellow and white sandy clay	Ŷ	./c
49				(white med to coarse sand and minor white		:::
52				(clay	\	
				White & yellow sandy clays, laterite pebbles		:0
58				(possible contamination)		10 10
V				Med. to coarse white sands & minor hard kaolin		
63				chips - (uncertain if kaolin or magnesite)		
	-w. 1.27			White sandy clay		ر ر
66'						0.
				White to pale grey hard kaolin? (magnesite?) -	;	$\Delta \Delta$
70				amorphous, minor sand and clay		$\nabla \nabla$
		_		White sandstone - med grn with white kaolin		
82		_ :		cement - granular texture - decreasing cement with		
	A A		·	depth	1.	
	e.103			Sandstone & brown to black chert chips, silicified		1
83.2	RE			dolomite chips		Δ - Δ
						1 23
IC TYPE/	CO:		• • • • •		OG BY	
TÄRT: TNISH:		- 3		VATER TABLE		

(10N: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral component incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

				CO-CROINATES LOCAL BLO HOLE NO AMG BEARING		
PROJEC	T NAME:	W	ΩOLI	NER COLLAR R.L. INCLINA	TION	
DEPTH	Length Drilled		VERY	DESCRIPTION: See footnotes for information required.	SAMPLE NUMBER & TYPE	Graph Log
83.2		0.4	•	Brown silicified dolomite - broken core and		1
84.4				kaolin		
	3	0.2	5	Kaolinite & fragments of silicified dolomite		
87.4				showing lamination		
	3	0.3		Black grey and brown silicified dolomite		
90.4				fragments - distinctive carbonate texture		
				White and grey kaolinite - generally		
	3	2.5		massive - some minor brecciation		-
93.4	427.3					
		:		White kaolinite, strongly brecciated with		
Ē	-			sandy breccia matrix - 15cm of		
96:4				silicified dolomite		-
E	3	1.3		White kaolin and small silicified dolomite		
99.4	·			fragments mixed together		
<u> </u>		0.4!	5.	Kaolin and silic. dolomite fragments - large		
102.4				grey silicified dolomite fragments - scattered		
ЕОН						
				102m of Pvc tubing and with bottom m split		
E				and wrapped in copper wire.		
E						
DRILLING					LOG BY	
RIG TYPE				•••••••		
DRILLER:	· · · · · · · · · · · · · · · · · · ·			CASING SIZE & LENGTH	PACE 3 OF	3.
FINISH:					HOLE T.D	

*DESCRIPTION: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral componing. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

DRILLING LOG

HOLE NUMBER WD5

~~	TTT	T 00	FORM
112	1 1 . 1 .	1.6 16	

				CO-CRDINATES LOCAL B9 LOCAL B9 AMC LNER COLLAR R.L.* HOLE NA BEARING INCLINA	;	
		RECO	VERY	DESCRIPTION:	SAMPLE	<u></u>
DEPTH	Length Drilled		T	See footnotes for information required.	NUMBER & TYPE	Craphic Log
1				Grey sandy soil		W
3				Yellow sand with laterite pebbles		0.0
4				Black mud with humic material		<u> १२२</u>
6				Black sandy mud with abund. lat. pebbles & minor	shells	·2.
V				red-orange-yellow white mottled sandy clays		0
20						D _O
1		<u> </u>		Pale grey sandstone - med grn qtz with grey clay		
30				(lithified) matrix		Δ.Δ
-				White sandy clay		.C.
61			·			ن ن ن
<u> </u>	.		-	White clayey (kaolin)sand - coarse qtz grains,		
77		:		grain-supported sandstone with kaolin cement	į	
V				coarse qtz sand - sub angular fragments		
82	· · · · · · · · · · · · · · · · · · ·					
<u> </u>				Pale grey to white lithified kaolinite, minor		Δ
93				coarse sand		ΔΔ
,				Weathered granitic material - coarse angular		+
				qtz and kaolin	_	+-
				118-first fragments weathered green schist- trace		+
133			-			+
RILLER:. TÄRT:	MODEL:Wa G. 22/11	rma Mor /82	r.is 	CASING SIZE & LENGTH . NO. COLE	PAGE (. OF	2.
TNISH: PURPOS ESCRIPTI		bur & L. mat	ROCK	VERTICAL SCALE aphy RESULT: Granite NAME: degree of weathering or alteration, proportion, grain size & roundness of bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto) ORE HINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.		

DRILL	TOO	TO CATALOG A
11K11.1.	1.1 11 -	F C J K IVI

TITLE	NAME/NO.	:		CO-ORDINATES LOCAL HOLE No		
PROJEC	T NAME:	• • • • •	WQQ.	LNER COLLAR R.L. BEARING INCLINA	TION	
DEPTH	Length Drilled	RECCO L.√V	_	DESCRIPTION: See footnotes for information required.*	SAMPLE MUMBER & TYPE	Craphi Log
134						+
	Bladed =>			Weathered granite - coarse qtz - pink feldspar		++
	Bla		,	trace mica and green schist - small granite		+
<u> </u>				fragments		1
150						+
-	ЕОН			Hole bottomed in weathered granite -		
		1		not suitable for coring.		
				Pvc tubing inserted to 150m with bottom lm	·	
				split and entwined in copper wire.		
						-
			1			
				<u> </u>		
			-			
-						<u>.</u>
		-			-	
PRILLING (_				LOG BY B. R	\
RIG TYPEA DRILLER: TART:	ODEL:			CORE/HOLE SIZE & LENGTH CASING SIZE & LENGTH WATER TABLE	PACE 2 OF2	
INISH:					HOLE T.D	*****

DESCRIPTION: Colour & ROCK NAME: degree of weathering or alteration, proportion, grain size & roundness of mineral componer incl. matrix; bedding (quantitative, angle to LCA), schistosity (ditto), fractures/joints (ditto), brecciation (quantitative); ORE MINERALS (%, grain size & names); VISUAL ESTIMATE OF GRADE.

